# Validation of CRTM using Observation and RTTOV



Emily Liu<sup>1</sup> and Andrew Collard<sup>2</sup>





<sup>1</sup>SRG @NOAA/NCEP/EMC <sup>2</sup>IMSG @NOAA/NCEP/EMC Members of JCSDA CRTM Team

2019 International Workshop on Radiative Transfer Models for Satellite Data Assimilation

## Outline

- Introduction and Motivation
- Overview of differences between CRTM and RTTOV
- Validation of CRTM under scattering conditions for MW sensors
  - Surface Reflectivity
  - Jacobians
- Validation of CRTM under clear-sky conditions for IR sensors
  - Surface Emissivity over Ocean
  - Input Profile Check
  - Simulation of ABI
- Summary



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## **Introduction and Motivation**

- CRTM and RTTOV are each called from the GSI data assimilation system using the same model background fields and compared both with each other and with observations (i.e., the innovation).
- This allows direct comparison between model calculations for a wide range of atmospheric scenarios.
- Differences may be:
  - Deficiencies in one or both models
  - Errors in implementation of the model
  - Opportunities for improvement.



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## **Overview of Differences CRTM and RTTOV**

Features	CRTM	RTTOV
Interface	<ul> <li>Unified interface for all sensors and conditions</li> <li>(MW-IR-VIS)</li> </ul>	<ul> <li>Separate interfaces for MW ad IR sensors under all-sky conditions (e.g. RTTOV-SCATT for MW, RTTOVCLD for IR)</li> <li>Discrete Ordinates Method (DOM) multiple-scattering solver for thermal emission and solar radiation</li> <li>MFASIS for fast visible cloud scattering parameterization</li> <li>Principal components</li> </ul>
Hydrometeor Types	<ul> <li>Six types: water, ice, rain, snow and hail and graupel</li> </ul>	<ul> <li>Four types: cloud water, ice, rain, snow (MW)</li> <li>Six types: two stratus, three cumulus, one cirrus (IR)</li> </ul>
Total Cloud Cover	<ul> <li>Four cloud overlapping schemes available</li> </ul>	<ul> <li>Hydrometeor weighted average overlap (MW)</li> <li>Maximum-random overlap (IR)</li> </ul>
Fractional Cloud	<ul> <li>Two-column radiance calculation</li> </ul>	<ul> <li>Two-column radiance calculation (MW)</li> <li>Streams method (IR) – six columns (for six cloud types)</li> </ul>
RT Solver	<ul> <li>Thermal emission for clear-sky and non- precipitating clouds</li> <li>Advanced Adding and Doubling scheme for scattering conditions</li> </ul>	<ul> <li>Delta-Eddington approximation (MW)</li> <li>Single-stream RTE with scaling of cloud optical thickness to account for the scattering effect (IR)</li> <li>Discrete ordinates method (IR/VIS)</li> </ul>

## **Overview of Differences CRTM and RTTOV (contd.)**

Features	CRTM	RTTOV	
Hydrometeor Optical Properties for MW	All hydrometeors – sphere; modified gamma	Cloud water/ice – sphere; gamma Rain – sphere; Marshall – Palmer Snow – sector; Field	
	Mie solution	Mie solution (sphere); DDA (non-spherical)	
	Constant density for each hydrometeor type LUT for bulk mass extinction and scattering coefficient and Legendre phase function coefficients, as a function of frequency, temperature, hydrometeor type and effective radius	Constant density for each hydrometeor type LUT for volume extinction coefficient, single scattering albedo and average asymmetry, as a function of frequency, temperature, hydrometeor type and hydrometeor water content	
Hydrometeor Optical Properties for IR	Similar to MW except for cloud ice Cloud ice – hexagonal columns; gamma Mie solution – sphere FDTD/GO – non-spherical	Water clouds – sphere; modified gamma Ice clouds – ensemble of shapes & PSDs Mie solution – sphere Baran database – function of IWC & temperature	

#### **Overview of CRTM and RTTOV (contd.)** Transmittance Coefficients

- Transmittance coefficients for the regression based transmittance model are solved by applying the diverse profiles of atmosphere states and the corresponding transmittances computed from the LBL model
- Due to the variability of water vapor in space and time, water vapor is more challenging than other gases to fit in a regression-based, fast RT model and produce smooth, physically representative Jacobian profiles
- Transmittance Models
  - ODPS (Optical Depth at Pressure Space) fit the effective channel transmittance using the prescribed pressure levels
  - ODAS (Optical Depth at Absorber Space) fit the effective channel transmittance with the prescribed integrated absorber amount levels
- RTTOV ODPS for all variable gases
- CRTM
  - ODAS for water vapor and OPDS for all other variable gases
  - A polynomial function is applied to the regression coefficients to improve the water vapor Jacobians (Compact OPTRAN).
- ODAS generates more accurate fast RT transmittance than ODPS for water vapor

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## Validation of CRTM under scattering conditions

- The current all-sky assimilation at NCEP/EMC is for non-precipitating clouds as precipitating hydrometeor profiles are not available from the current model.
- To facilitate the move to include these extra variables (available in the soon-to-be-implemented FV3GFS model), we compare CRTM and RTTOV scattering calculations....



#### **Evaluate CRTM under Scattering Conditions** Issue



- The calculated CRTM BTs have systematic biases for surface sensitive channels (1-5, and 15) at locations where ADA solver is involved.
- The CRTM BTs are too cold in general

### **Evaluate CRTM under Scattering Condition** Enhancement



- It is found that the off-diagonal terms of the surface reflectivity matrix is zero so that there is no diffuse radiation being reflected towards the viewing direction
- A proper surface emissivity model to work with multiple scattering algorithm is necessary
- For non-scattering RT, a reflection correction is used to account for the diffuse radiation.

The work-around to reduce the biases:

- Reflection correction is included in conjunction with ADA solver and the correction is only applied to stream angles ≤ 60°
- Stream angles > 60° are taken as 60° when ADA is on.
- The bi-directional reflectance distribution function (BRDF) for MW will be developed to replace the work-around.





The resulting CRTM BTs with the work-around are comparable to RTTOV BTs for all channels in AMSU-A

#### **Enhancement of CRTM** Validation w/ Obs

**CRTM OmF** 





**AMSU-A** 



**ATMS** 



**MHS** 





#### **SSMIS**

#### **Enhancement of CRTM** Impact of Fractional Cloud Coverage

- The impact of fractional cloud coverage on BT can be significant when precipitation is involved
- For high frequency channels, the impact could be over 100 K
- Significant impact for small-scale cloud and precipitation in convective region





T <sub>b,cloudy</sub>	T <sub>b,clear</sub>	MHS 89 GHz
250K	228K	$T_{b,overcast} = 245K \longrightarrow T_b = 235K$
		$T_b = (1 - TCC) \times T_{b,clear} + TCC \times T_{b,cloudy}$
тсс		$T_b = (1 - 0.33) \times 228 + 0.33 \times 250 = 235$
0.33	0.67	

#### Surface Emissivity Jacobian CRTM vs. RTTOV



#### **Surface Emissivity Jacobian**

### **Unphysical Response and the Fix**

![](_page_21_Figure_2.jpeg)

- Unphysical response occurs under scattering condition where ADA is used (blue area)
- This was found to be an output logic error under scattering condition in CRTM and has been corrected for CRTM 2.3.1 release
- The quality of the surface emissivity Jacobian is crucial for the radiance data quality control and bias correction procedures in GSI analysis

![](_page_21_Figure_6.jpeg)

![](_page_21_Figure_7.jpeg)

![](_page_22_Figure_1.jpeg)

![](_page_23_Figure_1.jpeg)

![](_page_24_Figure_1.jpeg)

![](_page_25_Figure_1.jpeg)

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![](_page_26_Picture_11.jpeg)

#### Evaluate CRTM for Clear-sky IASI Compare with RTTOV

#### RMS ( $BT_{CRTM} - BT_{RTTOV}$ )

![](_page_27_Figure_2.jpeg)

### Evaluate CRTM for Clear-sky IASI Validation w/ Obs & RTTOV

#### $RMS(OMF_{RTTOV}) - RMS(OMF_{CRTM})$

![](_page_28_Figure_2.jpeg)

#### **Evaluate CRTM for Clear-sky IASI** Validation w/ Obs & RTTOV

![](_page_29_Figure_1.jpeg)

#### **Evaluate CRTM for Clear-sky IASI** Compare w/ RTTOV

#### Surface Emissivity Difference (RTTOV - CRTM)

![](_page_30_Figure_2.jpeg)

#### **Evaluate CRTM for Clear-sky IASI** Validation w/ Obs & RTTOV

![](_page_31_Figure_1.jpeg)

The OMF values show improvement not only in surface channels but in near surface channels as well

The IREMIS (RTTOV) model includes sea surface temperature in estimating sea surface emissivity (Newman 2005)

CRTM is working on including the sensitivity of sea surface temperature to its IR ocean emissivity model

![](_page_32_Picture_1.jpeg)

#### IASI Channel 600 2561 cm<sup>-1</sup>

![](_page_32_Figure_3.jpeg)

![](_page_32_Figure_4.jpeg)

![](_page_33_Figure_1.jpeg)

**CRTM OMF** 

**Direct Solar Reflectivity** 

0.4 0.6 0.8

0.2

0.08

0.1

![](_page_34_Figure_1.jpeg)

2561.00 cm<sup>-1</sup> Mean -7.8 STD 12.4 Min -108.0 Max 27.8 Nobs 365910 -16.0

120

120 16.0

![](_page_35_Figure_1.jpeg)

#### Evaluate CRTM for MHS Surface Emissivity

#### Impact of SZA Fix

![](_page_36_Figure_2.jpeg)

#### **Clear Sky Radiances**

#### Comparison between CRTM & RTTOV

TABLE 1. Summary of the wavelengths, resolution, and sample use and heritage instrument(s) of the ABI bands. The minimum and maximum wavelength range represent the full width at half maximum (FWHM or 50%) points. [The Instantaneous Geometric Field Of View (IGFOV).]

Future GOES imager (ABI) band	Wavelength range (µm)	Central wavelength (µm)	Nominal subsatellite IGFOV (km)	Sample use	Heritage instrument(s)
I	0.45-0.49	0.47	I.	Daytime aerosol over land, coastal water mapping	MODIS
2	0.59-0.69	0.64	0.5	Daytime clouds fog, inso- lation, winds	Current GOES imager/ sounder
3	0.846-0.885	0.865	Ĩ	Daytime vegetation/burn scar and aerosol over water, winds	VIIRS, spectrally modified AVHRR
4	1.371-1.386	1.378	2	Daytime cirrus cloud	VIIRS, MODIS
5	1.58-1.64	1.61	I.	Daytime cloud-top phase and particle size, snow	VIIRS, spectrally modified AVHRR
6	2.225-2.275	2.25	2	Daytime land/cloud properties, particle size, vegetation, snow	VIIRS, similar to MODIS
7	3.80-4.00	3.90	2	Surface and cloud, fog at night, fire, winds	Current GOES imager
8	5.77-6.6	6.19	2	High-level atmospheric water vapor, winds, rainfall	Current GOES Imager
9	6.75–7.15	6.95	2	Midlevel atmospheric water vapor, winds, rainfall	Current GOES sounder
10	7.24-7.44	7.34	2	Lower-level water vapor, winds, and SO <sub>2</sub>	Spectrally modified cur- rent GOES sounder
н	8.3-8.7	8.5	2	Total water for stability, cloud phase, dust, SO <sub>2</sub> rainfall	MAS
12	9.42-9.8	9.61	2	Total ozone, turbulence, and winds	Spectrally modified cur- rent sounder
13	10.1-10.6	10.35	2	Surface and cloud	MAS
14	10.8-11.6	11.2	2	Imagery, SST, clouds, rainfall	Current GOES sounder
15	11.8-12.8	12.3	2	Total water, ash, and SST	Current GOES sounder
16	13.0-13.6	13.3	2	Air temperature, cloud heights and amounts	Current GOES sounder/ GOES-12+ imager

Source: Schmit, T.J., Gunshor, M.M., Menzel, W.P., Gurka, J.J., Li, J., Bachmeier, A.S., 2005, Introducing the Next-Generation Advanced Baseline Imager on GOES-R, Bulletin of the American Meteorological Society, v. 86, p. 1079-1096.

#### **CRTM Simulation of ABI**

#### **Room for Improvement**

![](_page_38_Figure_2.jpeg)

![](_page_38_Figure_3.jpeg)

#### **CRTM Simulation of ABI**

#### **Room for Improvement**

![](_page_39_Figure_2.jpeg)

The v8 predictor files allow variable CO<sub>2</sub> and has different treatment of H<sub>2</sub>O

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![](_page_40_Picture_11.jpeg)

## Summary

- Careful evaluation of radiative transfer capabilities is extremely important.
  - Having a large number of users evaluating the system will help catch possible errors sooner
  - Forward calculations and Jacobians (and TL/AD) should be considered.
- Intercomparison with other models can be an effective way to quickly find bugs and implementation errors
  - It can also point to possible areas where one or both models may potentially be improved.
- We will continue with these comparisons including extension into shortwave infrared and visible wavelengths and cloudy infrared.

![](_page_41_Picture_7.jpeg)