



Solar radiance modelling and assimilation

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Outline 1. Motivation

- 2. Synthetic satellite images for solar channels Fast 1D RT, 3D effects, cloud overlap
- 3. Applications data assimilation, model evaluation







5 JUNE 2016

N O R T H S E A

Brussels

images from NASA WorldView

MODIS 11µm thermal (window) channel, 1km resolution

180K 340K

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20 m







5 JUNE 2016

low clouds can be distinguished from ground / sea

cloud surface structure

Sensitivity to cloud phase, particle radii

high resolution

images from NASA WorldView

cloud

shadows

MODIS $0.6 \mu m$ / $0.8 \mu m$ / $1.6 \mu m$ solar channels, 250m / 250m / 500m resolution

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Strategy for fast radiative transfer method MFASIS



Simplifications

- Simplified Equation:

Method for Fast Satellite Image Synthesis

 $3D RT \rightarrow 1D RT$ (tilted independent columns) Computational effort for a SEVIRI image of Germany: CPU-days (3D Monte Carlo) \rightarrow CPU-hours (1D DISORT)

- Simplified vertical structure:

Cloud water and ice can be separated to form two two homogeneous clouds at fixed heights without changing reflectance significantly

- \rightarrow only 4 parameters (optical depth, particle size)
- + 3 angles, albedo \rightarrow 8 parameters per column

Reduction of computational effort

Compute **reflectance look-up table (LUT)** with discrete ordinate method (DISORT) for all parameter combinations \rightarrow effort for looking up reflectances: CPU-minutes

Problem: Table is huge! O(10GB) \rightarrow not suitable for online operator, slow interpolation \rightarrow **compress table to 20MB** using truncated Fourier series \rightarrow CPU-seconds





Look-up table compression in MFASIS

- **Problem:** $R(\theta, \theta_0, \Phi \Phi_0)$ contains a lot of rainbow-related small-scale features
- **Solution:** Consider $R(\theta, \theta_0, \alpha)$ instead : smooth function for constant scattering angle α
 - \rightarrow approximate by 2D Fourier series, obtain Fourier coefficients by fit to DISORT results



We need to store only 18 coefficients C_{kl} , S_{kl} instead of O(1000) reflectance values (for each combination of the remaining 6 parameters) \rightarrow compression by a factor of ~O(100)





Accuracy and computational effort

Error of MFASIS (8 parameters/pixel) with respect to DISORT (full profiles available) (model data: COSMO-DE fcsts for 10-28 June 2012)



Computational effort per column: DISORT (16 streams): 2.3 x 10⁻² CPUsec MFASIS (21MB table): 2.5 x 10⁻⁶ CPUsec (on Xeon E5-2650, for 51 level COSMO data)







MFASIS for aerosols?

Fourier compression:

No obvious problem, same number of Fourier coefficients should be ok for aerosols...

Main problems:

- Large number of LUT dims.: Many species (O(10)), in case of COSMO-ART also effective radii
- Vertical profiles matter: Aerosol A can overlap aerosol B and vice versa



Reflectance(SZA,VZA) for const. scatt. angle + fit residuum

Plan: Investigate how the vertical structure can be simplified without causing too large reflectance errors, test if it is sufficient to consider only the N (=2,3?) most important aerosol species in each column...





3D effects not accounted for in 1D radiative transfer







3D effects not accounted for in 1D radiative transfer



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Accounting for 3D RT effects: Cloud top inclination



Rotated frame of reference with ground-parallel cloud \rightarrow nearly a 1D problem (inclined ground is taken into account by using a modified surface albedo) \rightarrow Solve modified 1D problem, transform back to non-rotated frame.





Cloud top inclination



SEVIRI 0.6mu+0.8mu, 3 June 2016, 6UTC 3h COSMO fcst without 3D correction

Cloud top definition : optical depth 1 surface (detect tau=1 in all columns, fit plane to column and 8 neighbour columns)

Cloud top inclination correction \rightarrow Increased information content Much more cloud structure is visible, in particular for larger SZAs For instance, one can distinguish convective from stratiform clouds





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Comparison with 3D Monte Carlo RT calculations



Clean comparison (only RT errors, no model errors) based on high-res. ICON runs from the HD(CP)² project:

- RMSE is reduced
- Histogram shape is improved
- Derived empirical function to scale down 3D correction for thinner clouds

Other 3D effects are still missing (e.g. shadows, flux through cloud sides)...





Cloud overlap schemes for synthetic satellite images

RT must take assumptions about **overlap of subgrid clouds** into account. **So far:** Schemes for 1D RT in vertical columns, > O(10km) grid, deterministic **Here:** Columns tilted towards satellite, O(km) grid, ensemble DA (no adjoint req.)

\rightarrow Questions: How important is...

...the uncertainty related to the unknown subgrid cloud distribution for DA? **Do we need a stochastic scheme?**

...this inconsistency: Overlap assumptions are valid for **vertical direction**, whereas 1D RT is performed in **columns tilted towards** the satellite

...the **cloud size distribution**? (What are we assuming at the moment?)

 \rightarrow experiments with different schemes

Not addressed: "What is the best overlap assumption?" We consider only maximum-random (as used in COSMO):

Clouds in adjacent layers overlap maximally, clouds separated by empty layers overlap randomly.



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Implementation 1: RTTOV "streams" approach

(Matricardi 2005)

• For each layer: compute total cloud fraction γ_{tot} , "right-align" single cloud

$$-\gamma_{\text{tot}}^{\text{randmax}}(n) = (1 - \gamma_1) \prod_{k=2,\dots,n} \frac{1 - \max(\gamma_k, \gamma_{k-1})}{1 - \gamma_{k-1}}$$

(can be derived under the assumption that there are no horizontal correlations)

 Subcolumn boundaries are created where needed → variable number of subcolumns ("streams"), up to 2N_z









Implementation 2: Stochastic cloud generator

Räisänen (2004), Marquart & Mayer (2001)

- Fixed number of subcolumns N
- Subcolumns are independently filled with clouds according to stochastic rules
 → correct expectation values γ, γ_{tot}
- Clouds wider than 1 subcolumn form only by chance → we assume clouds are as small as possible.
- Convergence: Average over n → ∞ realizations converges to the same value as every single realization for N → ∞
 - \rightarrow spread is only related to discretization error, vanishes for N $\rightarrow \infty$
 - $\rightarrow\,$ Physical spread requires finite cloud size







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Implementation 3: Stochastic continuous clouds

- Exactly one continuous cloud per layer (periodic boundaries)
 → clouds as large as possible
- Fixed number N of subcolumns (cloud fraction discretization error < 1/N)
- Cloud positions are limited by rand.-max. rules, but otherwise random
- Average total cloud cover for many realization converges to a value that is in general different from the equation (correlation between subcolumns)
- Spread does not vanish for N → ∞
 but converges to a finite value

Upper limit for spread that would result from a more realistic cloud size distribution?









Example: 3 clouds with constant cloud fraction 0.25 spanning several layers, bundle of 8 x 8 subcolumns

Maximum overlap holds for the vertical direction, not along the tilted column \rightarrow compensate by shift in x-direction in each layer \rightarrow nearly vertical clouds



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Results for June 2016

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based on operational 3h COSMO-DE fcsts

- All 2D maximum-random implementations lead to similar results for the mean values
- **Spread** obtained for continuous clouds is small, exceeds 0.01 in only 12% of pixels, max. spread is 0.05
 - \rightarrow probably not relevant for DA
- Most consistent implementation:
 3D maximum-random results are closer to
 2D random than to 2D max.-rand.
 (local reflectance differences up to 0.15)
 - → Taking tilted columns into account is of similar importance as the choice of the overlap assumption.
- Missing in all implementations: **Shadows...** (Not a problem for thermal channels)



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LETKF (Local Ensemble Transform Kalman Filter) Assimilation experiments



- Reference runs: Cycling with conv. obs. from June 4th, 21UTC June 5th, 18UTC
- Runs with conventional obs. + 0.6µm VIS SEVIRI channel: Branched from ref. run at 5UTC \rightarrow first analysis at 6UTC







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160

140

120

100

80

60

40

20

160

140

120

100

80

60

40

20

0



Precipitation forecast improvements

P(R>0.5) only conventional obs.



There are also examples for the suppression of "false alarm" clouds with precipitation...

conv. + 0.6mu

0.9

0.8

0.7

0.6

0.5

0.4

0.3

0.2

0.1

0.0

0.9

0.8

0.7

0.6

0.5

0.4

0.3

0.2

0.1

Reflectance error evolution for different assimilation settings

RMSE is smaller than in **reference run** for all settings even after >3 hours. Bias evolution: some clouds dissolve

Full obs. density:

(~9300 obs./hour), obs. error 0.3 is better than 0.2 (corr. err.?)

Temporal thinning improves 3h fcsts

Temporal & spatial thinning: similar 3h fcst results

Impact on conventional observations

Relative change in RMSE of 3h forecasts caused by VIS assimilation: Mostly beneficial. But this is for only one day... Longer period is under investigation at DWD (Lilo Bach).

Model evaluation

HD(CP)² : ICON runs with 156m, 312m and 625m resolution for Germany.

Comparison with MODIS (250m)

Paralellized (MPI+OpenMP) offline operator based on MFASIS (still without cloud top inclination 3D correction)

Cloud size distribution: power law (down to effective model res.) reproduced

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Summary

- MFASIS: fast method for simulating solar channels (now in RTTOV)
- Cloud top inclination parameterization reduces the systematic error
- Cloud overlap: Tilted columns matter, reflectance spread is small
- Ensemble data assimilation: reflectances & precipitation are improved
- Useful tool for high resolution model evaluation

Next steps: Longer assimilation periods, optimization of assimilation settings, MFASIS for aerosols, more channels, more 3D effects...

Publications:

Scheck, Frerebeau, Buras-Schnell, Mayer (2016): A fast radiative transfer method for the simulation of visible satellite imagery, Journal of Quantitative Spectroscopy and Radiative Transfer, 175, p. 54-67.
Scheck, Hocking, Saunders (2016): A comparison of MFASIS and RTTOV-DOM, NWP-SAF visiting scientist report, http://www.nwpsaf.eu/vs_reports/nwpsaf-mo-vs-054.pdf
Heinze et al. (2017): Large-eddy simulations over Germany using ICON: a comprehensive evaluation, QJRMS, Vol. 143, Issue 702, p. 69-100
Scheck, Weissmann, Mayer (2018): Efficient methods to account for cloud top inclination and cloud overlap in synthetic visible satellite images, JTECH, Vol. 35, Issue: 3, p. 665-685

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Results for June 2016 (0.6µm SEVIRI images for 3h-COSMO-DE fcsts)

It is essential to take cloud overlap into account, setting all clouds fractions to 1 or using only grid scale clouds causes large errors.

Differences related to different assumptions or implementations are much smaller.

Good agreement with observations (no tuning!)

Cloud fraction 1 SEVIRI observation random overlap random-maximum overlap (2D stochastic continuous clouds)

grid scale clouds only

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Superobbing, Thinning and Localization

- **Superobbing:** 3×6 pixels $\rightarrow 18 \times 18$ km² in model space, O(eff. model resolution) Reflectance obs. every 15min \rightarrow 9255 reflectance superobs. per hour (> conv. obs.)
- Thinning, e.g. factors 4 in space & time \rightarrow 581 superobs. per hour (< conv. obs.)

- Different localizations (to avoid that VIS overwhelms conv. or vice versa)
 - Aim for both conv. and VIS: #obs. / grid point ~ O(ensemble size)
 - Reflectances: No vertical localization

Single observation experiments

- Model equiv. computed with nonlinear operator differ from LETKF estimate
- **Ambiguity** of VIS: LWC, IWC, RH are modified \rightarrow resolve using other channels?

Nonlinearity of the operator

Comparison of linear estimate for analysis model equivalents from LETKF and actual model equivalents obtained by applying nonlinear operator to analysis (incl. inflation, saturation adjustment): Significant differences for individual (super-)observations (blue), less impact on ensemble mean (red).

Reduces effectiveness of LETKF for large increments → avoid long assimilation intervals, assume larger observation errors? Outer-loop-like strategies?

Comparison with RTTOV-DOM

(with J. Hocking, R. Saunders)

RTTOV-DOM: Implementation of discrete ordinate method by MetOffice / NWP-SAF

See http://www.nwpsaf.eu/vs_reports/nwpsaf-mo-vs-054.pdf

- Reflectances for clouds agree well!
- Backscatter glory: reduced accuracy, depends on unknown width of size distribution
- Clear sky contributions problems:
 - -In MFASIS only a constant water vapour profile is included (affects 0.8µm channel)
 - \rightarrow linear correction developed
- RTTOV-DOM: no multiple cloud clear-sky scattering processes
 → negative bias for dense clouds

MFASIS has been included in RTTOV 12.2 by DWD + MetOffice

Performance

method	LUT size	time/pixel
DISORT (128 streams)	_	4.9s
DISORT (16 streams)	_	2.3×10^{-2} s
MFASIS ($N_k = N_l = 4$)	36MB	3.3×10^{-6} s
MFASIS ($N_k = N_l = 3$)	21MB	2.5×10^{-6} s
MFASIS ($N_k = N_l = 2$)	9MB	1.8×10^{-6} s

(on Xeon E5-2650 with 20MB level 3 cache)

Scattering angle randomly chosen in [130°,140°], all other parameters chosen completely randomly Scattering angle completely randomly: Only $N_k=N_l=4$ (LUT exceeds cache size) is 20% slower.

Part of the table required for one SEVIRI image fits into cache \rightarrow high performance

 $\rightarrow\,$ If required MFASIS LUT can be extended without degrading performance

Uncompressed LUT for R(θ , θ_0 , ϕ'): 7.5GB, limited α range does not help

 \rightarrow cache misses in almost every pixel \rightarrow slow!

Parameter values in the LUT

 C_{kl} , S_{kl} stored in LUTs with dims. α , τ_w , r_w , τ_i , r_i , A Parameter values are chosen such that linear interpolation error for reflectance < 0.005 Adaptive α -grid: high resolution (2°) is required only around cloud bow \rightarrow LUT factor 3 smaller

parameter	values
A	0.0, 0.5, 1.0
$ au_{\scriptscriptstyle W}$	0, 0.25, 0.5, 1, 2, 4, 8, 16, 25, 50, 100,
	300, 1000
$ au_i$	0, 0.125, 0.25, 0.5, 1, 2, 4, 8, 16, 25,
	50, 100, 300
r_w [μ m]	2.5, 5, 10, 25 Size: 21MB
<i>r_i</i> [µm]	20, 40, 60
α [°]	40, 45, 50, 60, 70, 80, 90, 99, 109, 119,
	129, 133, 135, 137, 139, 141, 143, 145,
	147, 149, 153, 159, 165, 171

Fit and interpolation errors

As a function of max. zenith angle and scattering angle for 3 x 3 Fourier terms:

Backscattering glory

Atmospheric state from June 15, 2012, 12UTC. Sun angles from other months.

Scattering angles larger than 175° in October / March \rightarrow Backscattering glory \rightarrow from a geostationary point of view the glory is not a rare event! Not included in the LUT \rightarrow errors several times larger Glory depends on width of droplet radius distribution \rightarrow no input data available Assimilation with higher assumed observation error may still be useful.

Error decomposition

What is the contribution of the various simplifications to the total error?

 ΔR_{rad} and ΔR_{int} are the most important and compensate each other partially. Higher accuracy (e.g. for 1.6µm) requires better way to compute effective radius.

Cloud top inclination correction

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Stochastic overlap schemes: Convergence for smallest / largest cloud approach

number of subcolumns

(a) Maximum deviation, 99% percentile of the deviation and root mean square deviation in ensemble mean reflectance for the stochastic maximum-random overlap method STO-N with different numbers of streams relative to the 512 stream case computed for the June 2012 test period. Ensembles with 100 members were used. (b) Like (a), but for the STO-C implementation.