





Model error representation in ensemble convection-permitting forecasts and ensemble data assimilation

Glen Romine NCAR (MMM/IMAGe)

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NCAR Collaborators: Craig Schwartz, Ryan Sobash, Kate Fossell, Judith Berner

Core motivation: hazardous weather prediction

Severe convection with the right combination of:

- Moisture
- Instability
- A lift mechanism
- Sufficient shear

Key forecast challenges to address:

What is the probability that convection will occur at a given location?

How intense might convection be?

What convective modes are favored (primary hazards)?



Forecast system design components:

- Ensemble want probabilistic, not deterministic predictions
- High-resolution for convective mode and intensity (convectionallowing (CAM) horizontal grid spacing)
- Computational constraint **regional model** (e.g., WRF)
- Ensemble data assimilation for initial conditions (e.g., DART)

Regional ensemble forecast system components:

- Initial condition uncertainty (e.g., ensemble DA)
- Surface and lateral boundary condition uncertainty
- Model error representation CAMs are notoriously under dispersive!



Example: NCAR Real-time ensemble



GFS + perturbations

ensemble.ucar.edu

Lower boundary: free forecast land surface, fixed sea state



NCAR ensemble – hazard prediction sample



Day 1 probability of UH > 75 m^2/s^2 w/ NWS warnings

Probability of simulated supercell thunderstorms (fill) overlain with issued severe weather warnings during the valid prediction period

See Sobash et al. (2016) for more on this effort



NCAR ensemble – skill/reliability for precipitation

Skillful across range of rainfall intensity, more skillful with larger neighborhood for verification

Modest bias at all intensities

Underdispersive*, less so For larger neighborhoods

* IC/BC perturbations only* No obs error assumption

Model error treatment?



Schwartz et al. (2015)



<u>None</u>

Rely on lateral boundary perturbations and initial condition diversity

Multi-model/multi-physics/multi-parameter

- Uncertain representations of physical processes
- Model dynamics/assumptions drive model climate
- Ensemble members may have varying skill and biases
- May be challenging to post-process (e.g. grids, variables, state size)

Stochastic methods

- Random model error process (ideally)
- Single model and physics climate
- Options available in WRF-ARW:
 - 1) Stochastic Kinetic Energy Backscatter Scheme (SKEBS)
 - 2) Stochastically Perturbed Parameterization Tendencies (SPPT)



Ensemble reliability – precipitation



Attributes diagrams @ 1 mm h⁻¹ threshold

Overconfident predictions of precipitation (no observation error)

Stochastic methods can improve reliability in longer range storm-scale forecasts, but little impact on short-range (< 12 h) prediction



Forecast bias and spread time series - temperature



Bias drift relative to Control forecast SPPT – largest bias drift, but also largest spread



Forecast verification against rawinsondes



Verification against 24 h forecasts



Practical reliability for precipitation forecasts





Observed state

Control ensemble: Estimates true evolution of the atmosphere Lacks sufficient dispersion to capture the observed evolution after short integration

Select options: Multi-XXX, calibration, perturbed boundaries, stochastic methods





For the NCAR ensemble, perturbing the lateral boundary condition improves spread somewhat, but late in the forecast. Ensemble mean is about the same. Note forecast area is far removed from true lateral boundaries.





SKEBS leads to greater dispersion, beginning earlier in the forecast, with nearly the same ensemble mean as the control and perturbed boundary ensemble.





SPPT leads to even greater dispersion, beginning much earlier in the forecast, but the ensemble mean is further from the observed state relative to the control. SPPT here requires calibration/tuning. Downside – some wild forecasts!



If only we could just IMPROVE the model!



Reduce dependence on spread to compensate for a poor model trajectory, try to **improve the forecast model** to evolve more like the real atmosphere.

Then – find structural error growth deficiencies that require model error approaches to correct.



Model error diagnostics in continuously cycled analysis

Continuous cycling is 'best practice'

First guess (**B**) for analysis is short **forecast** from prior analysis

Minimal 'spinup' needed, near the model attractor



For regional models – nearly all centers use 'partial' cycling – periodically replacing the background from another (often global) analysis, adjustment to regional model climate can take days

Bad forecast model = degraded background for the analysis and forecasts



Continuous cycled DA – model error revealed



Identify model errors through continuous cycled DA – compare analyses against observations or other (trusted) analyses (GFS above).



Observation space verification



3-km ensemble forecast verification against rawinsondes

40 forecasts (late April to early June)

Initial down-scaling, diurnal bias in mid- and lower-troposphere, drift near tropopause



Spread-Error ratio during forecast





Real-time analysis mean innovations

August 2015 mean analysis innovations for 00 UTC

Lowest model level water vapor

Lowest model level temperature





MYJ PBL scheme for analysis system

Classic cool and wet bias, but not everywhere



Real-time analysis mean innovations

December 2015 mean analysis innovations for 00 UTC

Lowest model level temperature

Lowest model level water vapor



MYJ PBL scheme for analysis system

Slight warm bias in December but regional variability



PBL physics – Surface T mean innovations – May 2015



Systematic errors in surface temperature are only weakly dependent on PBL physics. Need to test surface physics.





NCAR

WRF model challenges – surface moisture





WRF model challenges – surface moisture

Spike in surface moisture owes to decoupling PBL while latent heat flux is still positive

Most prominent in heavily vegetated areas with calm winds





Physics errors from downscaling



Ensemble 'warm start' 15-> 3 km GFS 'cold start' 0.5 -> 3km

0.60 LHFX (a) 0.50 PW difference (mm) _HFX difference (Wm⁻²) PW 16 0.40 12 0.30 0.20 0.10 0.00 12 20 22 06 08 10 12 14 16 18 00 02 04 Hourly precipitation difference (mm) (b) DM difference (mm) 0.00 -0.10 0.10 0.00 -0.10 lourly precipitation -lourly change in PW 0.20 12 16 14 18 20 22 02 08 10 12 00 04 06 700 hPa 850 hPa 925 hPa 2 m 12 16 20 22 02 06 08 10 12 14 18 00 04 Time of day (UTC)

Difference in 12Z vs. 00Z GFS initialized forecast time series (avg over 30 days)



Schwartz et al. (2015)

Future: High resolution ensemble analysis



Avoid errors from downscaling, leverage convective scale observations, improve model climate (?). Eventually, global ensemble.



Physics tendencies for further improvement

Demonstrated by S. Cavallo yesterday...





Adapted from Cavallo et al. (2016)

Summary – high resolution ensembles

Storm-scale ensemble design remains largely ad hoc:

- Stochastic methods to improve reliability
- Lots of opportunity to improve models at high-resolution prediction
- DA for high-resolution grids is still immature

Stochastic schemes are found to:

- improve ensemble dispersion characteristics
- introduce bias that may require additional spread
- difficult to verify adequate ensemble spread
- Effort is need to better target when and where additional spread is needed



Improving/diagnosing model climatology

1) Continuously cycled DA (ongoing)

- Improve model climate toward obs/trusted analysis
- 2) Careful analysis/verification of forecasts (ongoing)
 - Many examples the last few days, a few here as well

3) High resolution analysis grid (planned)

- Minimize physics and downscaling spinup errors
- Convective scale obs for analysis and verification
- 4) Physics tendency methods (planned)
 - e.g., Rodwell, Cavallo talks
 - Identify and correct error sources

