



Workshop on parameter estimation and inverse modelling for atmospheric composition 22 to 24 October 2013

Estimating emission rates of reacting constituents by variational inversion

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2. specific problems

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Data assimilation and Inverse Modelling: What makes the difference?

- Data assimilation: estimate
 - transient parameters:
 - prognostic parameters, state variables of prognostic systems
 - for reactive chemistry: initial values
- Inverse Modelling: estimate
 - intransient system control parameters,
 - forcing parameters,
 - diagnostic parameters
 - for reactive chemistry: <u>emission rates</u>, (deposition velocities)
 Time and spatial scales carefully to be considererd, if combined

Problem formulation (increasingly ill-posed)

- which parameter sets are to be optimized at all, and are the observations available?
- poor representativity of observations
- often only integral measurements available (NB: PM_x, NO_{x/y}, column soundings from satellites)
- reactive chemistry, and its observabiliy (with poorly observed emitted compounds)

SPECIFIC PROBLEMS

In the troposphere, for **emission rates,** the product (*paucity of knowledge*importance*) is high



The representativity problem (1): non-Gaussian distribution example: Rhine-Main area box: (Frankfurt-Mainz)

9.-10. August 1997



Incremental Formulation

- Analysis State: $egin{array}{ccc} x^a = x^b + \delta x^a \ u^a = u^b + \delta u^a \end{array}$
- New "State" Variables: $v = \mathbf{B}^{-1/2} \delta x$ $w = \mathbf{K}^{-1/2} \delta u$
- Cost Function:

$$J(\boldsymbol{v}, \boldsymbol{w}) = \frac{1}{2}\boldsymbol{v}^{T}\boldsymbol{v} + \frac{1}{2}\boldsymbol{w}^{T}\boldsymbol{w} + \frac{1}{2}\left[\mathbf{H}\delta\boldsymbol{x}_{i} - \boldsymbol{d}_{i}\right]^{T}\mathbf{R}^{-1}\left[\mathbf{H}\delta\boldsymbol{x}_{i} - \boldsymbol{d}_{i}\right]$$

• Gradient:

$$\nabla_{\boldsymbol{v}} J = \nabla_{\boldsymbol{v}} J_{IV} + \nabla_{\boldsymbol{v}} J_O = \boldsymbol{v} + \mathbf{B}^{\mathrm{T}/2} \nabla_{\delta \boldsymbol{v}} J_O$$
$$\nabla_{\boldsymbol{w}} J = \nabla_{\boldsymbol{w}} J_{EF} + \nabla_{\boldsymbol{w}} J_O = \boldsymbol{w} + \mathbf{K}^{\mathrm{T}/2} \nabla_{\delta \boldsymbol{w}} J_O.$$

The representativity problem (2):

Tesselated terrain and associated covariances

High resolution CORINE land use of the Cologne Aix-la-Chapelle area with a forest harboured station O





For high resolution, correlation functions need to bridge gaps

Integral information - Satellite data: ESA UV-VIS satellite footprints Ruhr area comparison



Integral information – Do NO2 columns profice useful information? Average OMI averaging kernel profile over model domain for July 9th 2006 OMI NO2 clm 100 200 pressure [hPa] 300 400 500 600 700 800 1888

model domain mean averaging kernel.

1.0

averaging kernel

1.5

2.0

2.5

0.0

0.5

Integral information – Exploitation of NO2 column averaging kernel information

- AK profile mostly dependent on optical properties of the atmosphere (cloud cover), rather than NO2
- typical maximal sensitivity above the boundary layer
- **does not allow** a clear distinction between PBL or lower free troposphere pollution burden

Integral information – Observation operator H

Formally an integral equation to be solved for vertical NO₂ molecule density function x (σ vertical coordinate)

$$y = \int_1^0 w(\sigma) x(\sigma) d\sigma$$
 $y = \sum_{k=1}^K h_k x_k$

At the minimum $\mathbf{x} =: \mathbf{x}_a$

$$d\mathbf{x}_a := \mathbf{x}_a - \mathbf{x}_b = (\mathbf{B}_0^{-1} + \mathbf{H}^T \mathbf{R}^{-1} \mathbf{H})^{-1} \mathbf{H}^T \mathbf{R}^{-1} \left\{ \mathbf{y}^0 - H[\mathbf{x}_b] \right\}$$
$$= \mathbf{B} \mathbf{H}^T (\mathbf{R} + \mathbf{H} \mathbf{B} \mathbf{H}^T)^{-1} \left\{ \mathbf{y}^0 - H[\mathbf{x}_b] \right\}$$

For scalar column retrieval:

$$d\mathbf{x}_{a}^{c} = \underbrace{\mathbf{B}\mathbf{h}^{T}(r+b)^{-1}}_{a} \left\{ y^{0} - h[\mathbf{x}_{b}] \right\}$$

adjoint representer

→ vertical structure function in **B** essential!

<u>Vertical</u> structure function: Extending the information from observation location by vertical exchange of polutants and information



Integral information – Modal representation of particles



Integral information -

Example: Aerosol Chemistry in **MADE**

Modal Aerosol Dynamics for EURAD/Europe (Ackerman et al., 1998, Schell 2000)

dM_i^k/dt=nuk_i^k+coag_{ii}^k+coag_{ij}^k +cond_i^k+emi_i^k

M_i^k:=kth Moment of ith Mode

assimilation of aerosol By satelite retrievals: e.g. Bridge from optical to chemical properties

MERIS MODIS AATSR+SCIAMACHY...



SYNAER retrieval algorithm

Species Mapping

EURAD-IM [µg/m³]		SYNAER - AOT
SO ₄ , NH ₃ , NO ₃ , H ₂ O, SOA		WASO (WAter SOluble)
Unidentified PM		INSO (water INSOluble)
Elemental Carbon		SOOT
Sea Salt		SEAS
Mineral Dust		DUST
radiative transfer model 🛛 🔶		
🗧 adjoint radiative transfer model		

Further significant problems, not considered here

- Meteorological driver model with boundary layer, convection, and precipitation
- chemistry model errors
- deposition velocities



ISSUES ON OBSERVABILITY

Observability ad a simple forcing model

Consider two one-way coupled linear models: $A(t_i, t_0)$ atmospheric model, $E(t_i, t_0)$ emission model, and C the coupling between both.

$$\begin{pmatrix} H(t_i), & 0 \end{pmatrix} \begin{pmatrix} x(t_i) \\ e(t_i) \end{pmatrix} = \begin{pmatrix} H(t_i, t_0), & 0 \end{pmatrix} \underbrace{\begin{pmatrix} A(t_i, t_0), & C \\ 0, & E(t_i, t_0) \end{pmatrix}}_{:=M(t_i, t_0)} \begin{pmatrix} x(t_0) \\ e(t_0) \end{pmatrix}$$

 $M(t_i, t_0)$ model integration from time t_0 to t_i $H(t_i)$ observation operator at time t_i x(t) atmospheric model state, e(t) emission parameter.

Observability

 $H(t_i)$ observation operator at time t_i

$$O_{v} := \begin{bmatrix} H(t_{0}) \\ H(t_{1})M(t_{1}, t_{0}) \\ H(t_{2})M(t_{2}, t_{0}) \\ \vdots \\ H(t_{v})M(t_{v}, t_{0}) \end{bmatrix}$$

Consider rank of matrix, Singular Vectors and -Values of

 $O_v^T O_v$

to be above some noise level.

How can we identify the degree of observability?

Given CTM (here RACM and EURAD-IM)
acting as tan.-lin. model operator
$$\mathcal{L}$$
: $\delta \mathbf{c}(t_F) = \mathcal{L}_{t_I,t_F} \delta \mathbf{c}(t_I), \quad \mathcal{L}_{t_I,t_F} = \frac{\partial \mathcal{M}_{t_I,t_F}}{\partial \mathbf{c}} \Big|_{\mathbf{c}(t_I)}$ **1. Berliner et al., (1998) Statistical design:**
"Minimize" the analysis error
covariance matrix A (say, via trace): $\min \mathbf{A} = \mathbf{B} - \underbrace{\mathbf{BH}^T(\mathbf{HBH}^T + \mathbf{R})^{-1}\mathbf{HB}}_{\text{to be maximized by H}}$ For this find maximal eigenvectors
as observation operators H,
which configure observations. $\mathcal{L}_{t_I,t_F}\mathbf{B}\mathcal{L}_{t_I,t_F}^T\mathbf{H}^T = \lambda\mathbf{H}^T$ **2. Palmer (1995) Singular vector analysis:**
Observe maximal SV configuration: $\max \frac{\|\delta \mathbf{c}(t_F)\|_{\mathbf{B}}^2}{\|\delta \mathbf{c}(t_I)\|_{\mathbf{B}}^2} = \max_{\delta \mathbf{c}(t_I)} \frac{\delta \mathbf{c}(t_I)^T \mathcal{L}_{t_I,t_F}^T\mathbf{B} \mathcal{L}_{t_I,t_F} \delta \mathbf{c}(t_I)}{\delta \mathbf{c}(t_I)^T\mathbf{B} \delta \mathbf{c}(t_I)},$

Empirical Kinetic Model Approach Scheme A prototype non-linearity example:

•Nitrogene oxides and numerous hydrocarbons act highly nonlinearly as precursors of ozone.

•Chemical conditions are either controlled by NOx or VOC deficit, delineating the "chemical regime".

•Both 4D-var and Kalman filter should start with the proper chemical regime.

EKMA diagram



Is $NO_x \underline{the}$ key to ozone production? And consequently, its observation the key to better forecast?



- \checkmark within a fixed time span
- \checkmark initial conventrations of NO / HCHO were varied
- \checkmark change of final concentration is given by colour
- \checkmark gradients (SVs) of maximyl ozone production given by arrows



When is the chemical regime sensitive? Leading **emission** singular vectors for VOCs (vs Nox)



TSVD of the optimal grouped relative singular vectors with respect to emission uncertainties for scenarios (a) PLUME and (b) URBAN.

Both panels depict VOC-sections of the grouped relative singular vectors.

When is the chemical regime sensitive? Leading **emission** singular vectors for olefines (left) and lower alkanes (vs Nox)



1st Projected Singular Vectors (HC3)



3D-targeted observations for flight missions

Given the need for a forecast at some final time t_F of a mission flight, what and where should be observed at the beginning t_I ?





Example observation targeting: SV optimal placement of observation sites



Initial concentrations and optimal horizontal placement of NO (left) and O3 (right) at surface level . Isopleths of the optimal horizontal placement are indicated with black lines.

(from Nadine Goris, PhD thesis 2011)

What is important to observe for O_3 prediction?

$\delta O_3 > \delta HCHO > \delta CO > \delta HONO > \delta OH$

with NO and NO₂ typically varying between $\delta O_3 > \delta NO_x > \delta HONO$ dependent on the specific chemical scenario.



Orders of magnitude of different compounds (left) and position of maximal values (right) of maximal singular vector entries per model level.

Where is it important to observe for O_3 prediction?



RESULTS

Implementation of emission rate optimisation In the troposphere, for **emission rates,** the product (*paucity of knowledge*importance*) is high

Emission Rate Optimization

minimize cost function

 $J(\mathbf{x}(t_0), \mathbf{e}) = \frac{1}{2} (\mathbf{x}^b(t_0) - \mathbf{x}(t_0))^T \mathbf{B}_0^{-1} (\mathbf{x}^b(t_0) - \mathbf{x}(t_0)) + \frac{1}{2} \int_{t_0}^{t_N} (\mathbf{e}_b(t) - \mathbf{e}(t))^T \mathbf{K}^{-1} (\mathbf{e}_b(t) - \mathbf{e}(t)) dt + \frac{1}{2} \int_{t_0}^{t_N} \left(\mathbf{y}^0(t) - H[\mathbf{x}(t)] \right)^T \mathbf{R}^{-1} (\mathbf{y}^0(t) - H[\mathbf{x}(t)]) dt$

deviations from background initial state deviations from a priori emission rates

model deviations from observations

- $\mathbf{x}^{b}(t_{0})$ background state at t = 0
- $\mathbf{x}(t)$ model state at time t
- $\mathbf{e}_b(t_0)$ background emission rate at t = 0
- $\mathbf{e}(t)$ emission rate field at time t

K emission rate error covariance matrix

H[] forward interpolator

 $\mathbf{y}^0(t)$ observation at time t

 \mathbf{B}_0 background error covariance matrix

Treatment of the inverse problem for emission rate inference



emission_treatmentid

How can the degree of freedom of emission variation be reduced? a "**strong constraint regularisation**" Normalised diurnal cycle of anthropogenic surface emissions *f(t)*

emission(t)=f(t;location,species,day) * v(location,species)

day in {working day, Saturday, Sunday} v optimization parameter



Optimisation of emission rates

amplitude optimisation for each emitted species and grid cell



Novel setup of joint emission rate-initial value optimisation after upgrade of regional emission inventory

- Preparation:
 - adaptation to extended emitted species, ("adjoint emission sources")
 - more biogenic emission





4D-var optimisation EURAD-IM 24 h window



7. July 2010, 0 UTC - 23 UTC

7. July 2010, 0 UTC - 23 UTC

NO₂ emission correction factors summer 2010

7. July

Emission Factor of NO2 (1)



x coordinate of projection (km)

Output from EURAD-IM

Range of Emission Factor of NO2: 0 to 20 1 Range of x coordinate of projection: 22.5 to 4792.5 km Range of y coordinate of projection: 22.5 to 4342.5 km Current time: 19 hours since 2010-07-07 00:00:00 Current sigma at layer midpoints: 0.9975 sigma_level File gradout_O3_epis_ivef_4dvar_eur_188.nc 16. July



x coordinate of projection (km)

Output from EURAD-IM

Range of Emission Factor of NO2: 0 to 20 1 Range of x coordinate of projection: 22.5 to 4792.5 km Range of y coordinate of projection: 22.5 to 4342.5 km Current time: 5 hours since 2010-07-16 00:00:00 Current sigma at layer midpoints: 0.9975 sigma_level File gradout_O3_epis_ivef_4dvar_eur_197.nc

analyses of NO₂ concentrations summer 2010

7. July

y coordinate of projection (km)



Range of y coordinate of projection: 22.5 to 4342.5 km Current time: 19 hours since 2010-07-07 00:00:00 Current sigma at layer midpoints: 0.9975 sigma_level File ctmout O3 epis ivef frw eur 188.nc

Range of y coordinate of projection: 22.5 to 4342.5 km Current time: 19 hours since 2010-07-16 00:00:00 Current sigma at layer midpoints: 0.9975 sigma_level File ctmout_O3_epis_ivef_frw_eur_197.nc

16. July

CO emission correction factors

Emission Factor of CO (1)



Pando of Emission Factor of CO: 0.0473919 to 8.1

<u>Analyses</u>, Example (i): Analysis of emissions by 4D-var (VERTIKO)



URAD System

Observed and analysed ozone evolution at St. Poelten Vertical bars: ozone observations with error estimates.



-----. emission factor optimisation. joint initial value and emission factor optimisation (Strunk et al., 2011)





Some BERLIOZ examples of NOx assimilation (20. \rightarrow 21. 07.1998)



Emission source estimates by inverse modelling Optimised emission factors for Nest 3



height layer ~32-~70m

surface

OMI NO2 column assimilation: Forecast validation summer (high convection)/ winter (low inversion)

July 1st to 15th 2006 inversion 16th-31st validation (by SCIA)



Cost function values of the background simulation (green) and analysis results (blue), normalised by the cost function value of the control simulation without data assimilation December 11 to 25, 2007 inversion 26^{th} to Jan. 10^{th} 2008 validation

NO2 Emission factors TNO emission inventory



Comparison of nitrogen dioxide emission factor results for the summer case study (left panel) and the winter episode (right panel). Please note the different ranges on the color scale.

SCIAMACHY O_{SCIA}mX_{OMIT} probability density functions



Control run (OmC) (no data assimilation at all,) black bold line, forecasted values (OmF) green bold line, analyses (OmA) blue bold line. For comparison: Gaussian fit to OmF pdf by mean and standard deviation given by broken purple line.

Comparison of NO2 tropospheric columns in molecules/cm2 for July 6th, 2006, 09-12 UTC.



Data assimilation result in terms of tropospheric columns for July 6th, 2006. NO2 model columns based on OMI and SCIAMACHY

assimilation within interval, 09-12 UTC.



Data assimilation result in terms of tropospheric columns for **July 7th**, 2006. NO2 model columns based on OMI and SCIAMACHY assimilation within the assimilation interval, 09-12 UTC.



Emission rate optimisation factors for NO2 after assimilation of

OMI retrieved NO2 tropospheric columns



2 x 4 days assimilation sequence. Left panel shows results after assimilation procedures from July 1.-4. 2006, right panel for July 7.-11., 2006. OMI data from KNMI

Summary

- Emission rate estimation cannot be performed isolated. Rather, other controling parameters (initial values) must be included.
- A smoother type spatio-temporal data assimilation algorithm is required (integration "backward in time and process": 4D-var)
- As ill-conditioned problem, regularisation is needed.
- Future: chemial scenario dependent covariances need to be designed
- A rigorous (control) theory based balance between initial value and emission rate weights is needed.
- Emission rates need be optimized SNAP-wise.

Some references

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