

Microphysics and convection in the "grey zone"

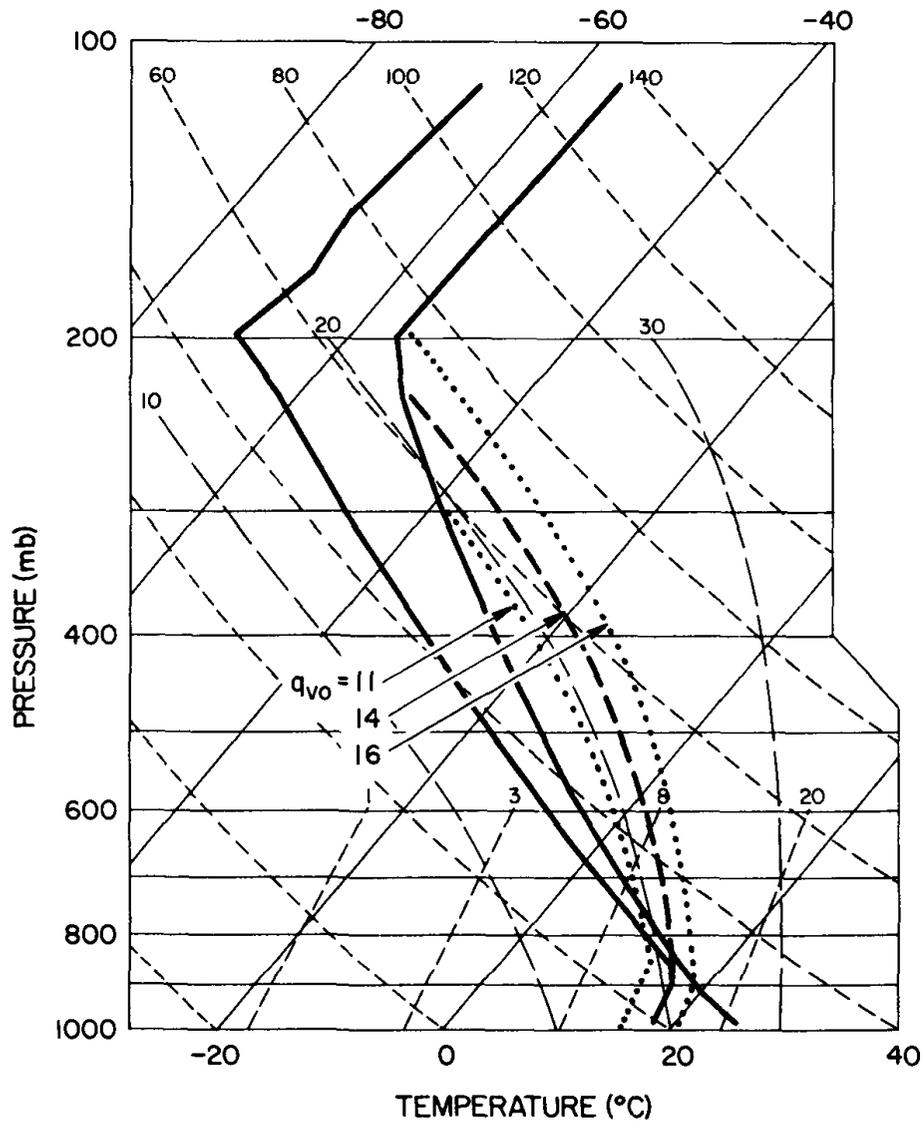
Luc Gerard

6 November 2012

(How) do we need to parameterize deep convection at resolutions of a few km ?

- Academic Study
 - No deep convection parametrization
 - Classical diagnostic scheme behaviour
 - A dedicated prognostic scheme: 3MT Model behaviour
 - Issues not addressed by 3MT
 - A complete scheme: CSU
 - The triggering problem – Model behaviour
- Real case multi-resolution behaviour
 - no conv param – diagnostic param – 3MT – CSU – Domain statistics –
- final remarks

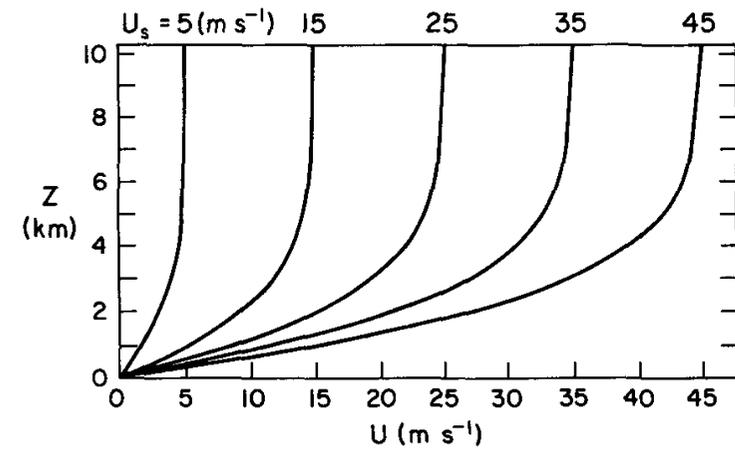
Academic test bench



Weisman & Klemp 1982: single profile with CAPE

Imposed pbl moisture

Zonal wind with vertical shear



Ellipsoidal bubble of θ perturbation

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- Cyclic domain, no Coriolis, no orography, no radiation.
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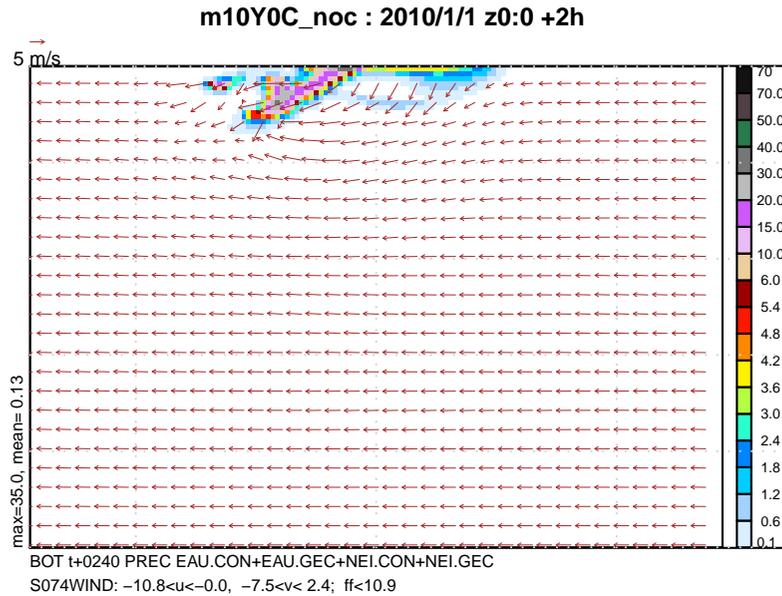
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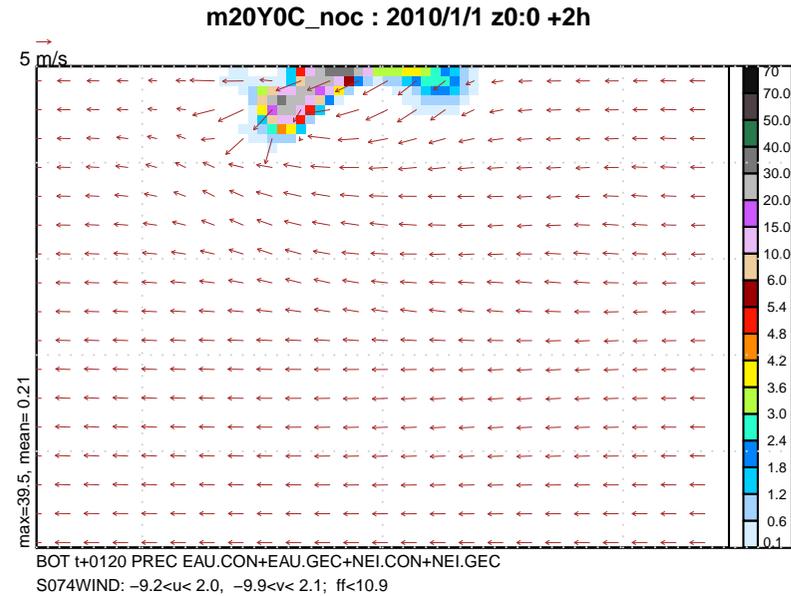
CAPE present over whole domain: very sensitive to perturbations inducing triggering

noCP: 2-h accumulated precipitation

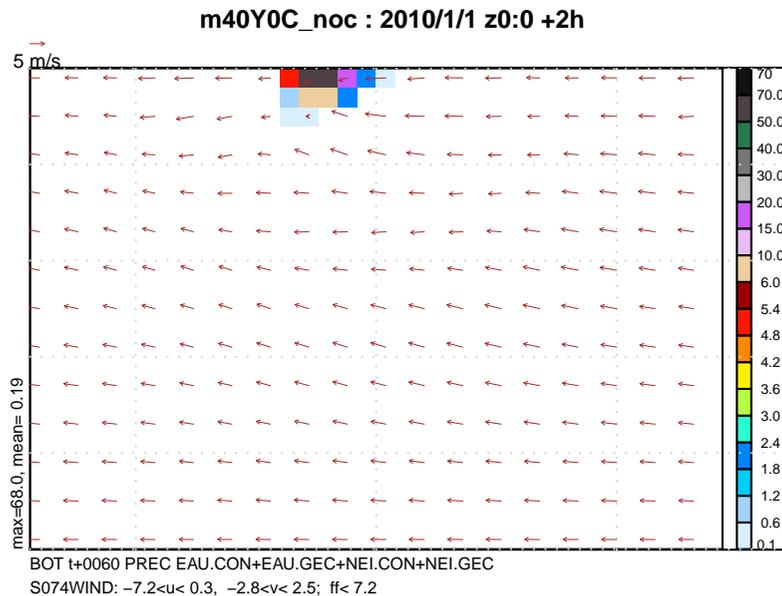
$\Delta x = 1km$



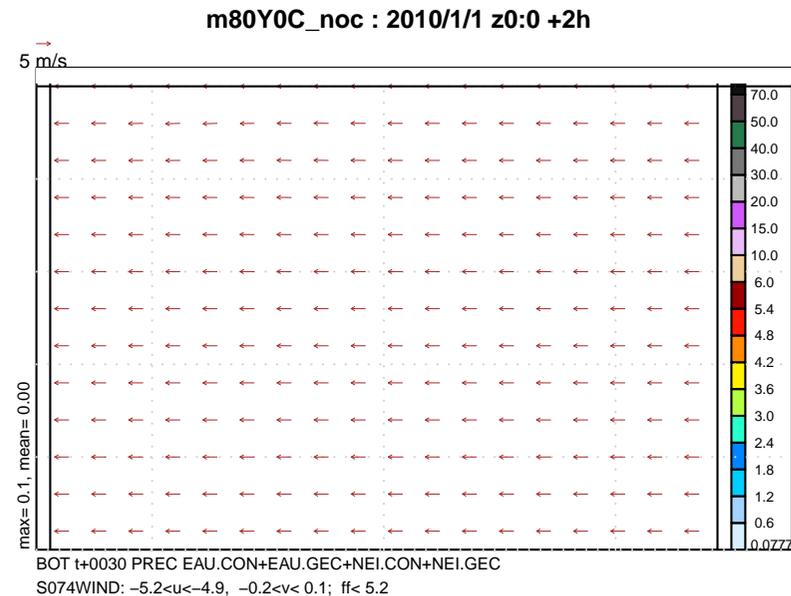
$\Delta x = 2km$



$\Delta x = 4km$

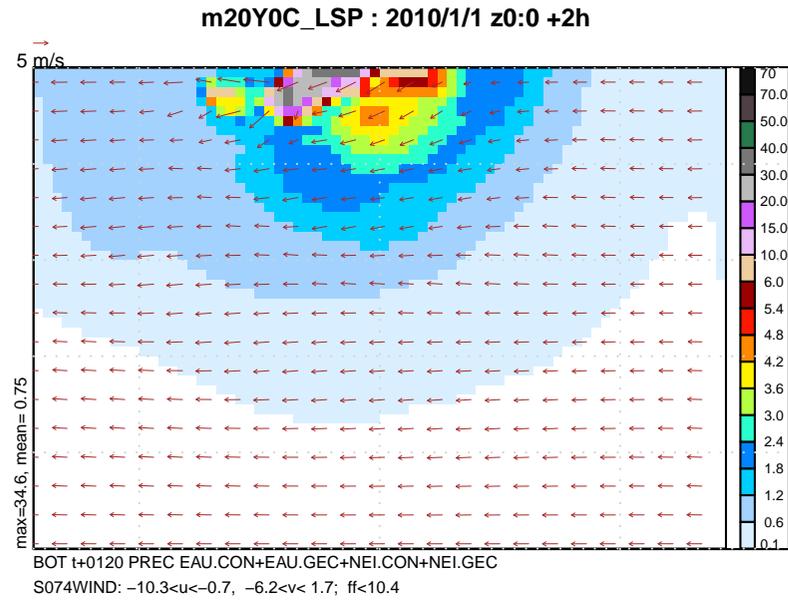
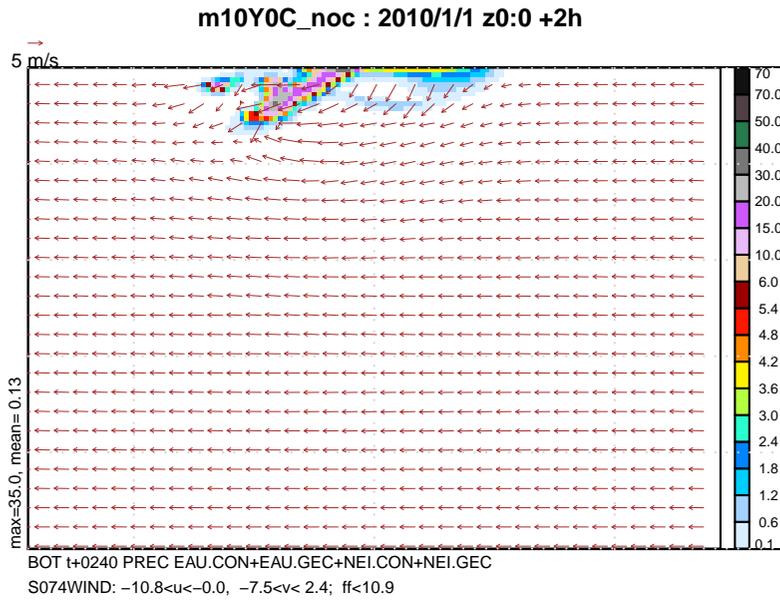


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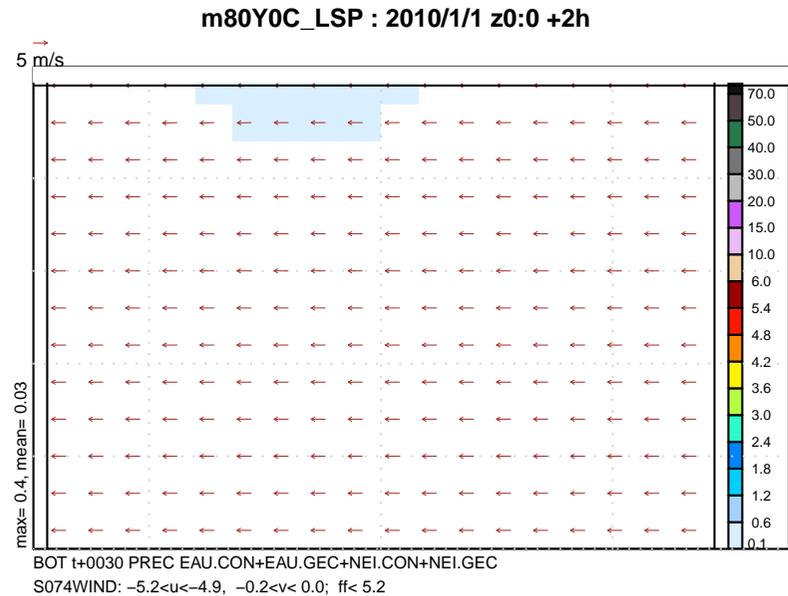
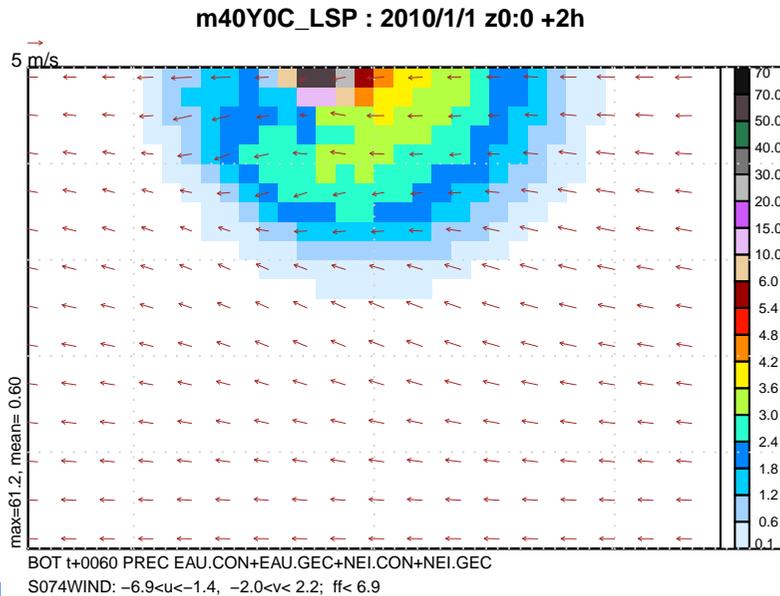
Diagn param: 2-h accumulated precipitation total

reference



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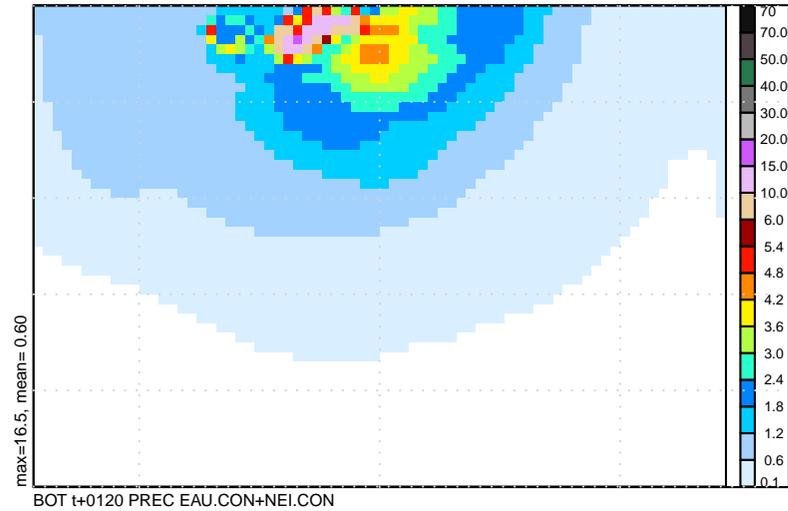


Diagn param: 2-h accumulated precipitation

subgrid part

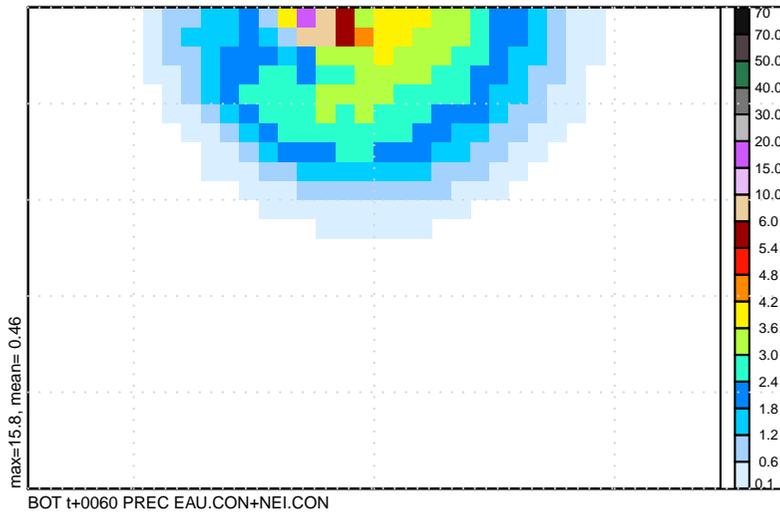
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m20Y0C_LSP : 2010/1/1 z0:0 +2h

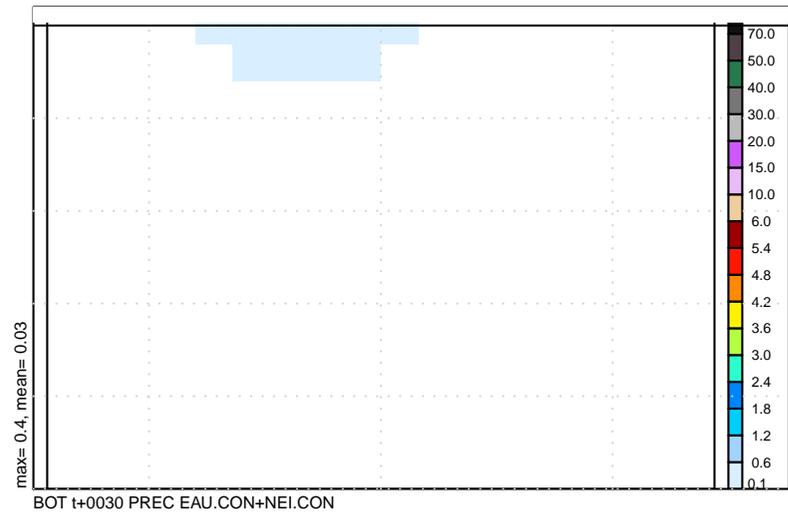


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m40Y0C_LSP : 2010/1/1 z0:0 +2h



m80Y0C_LSP : 2010/1/1 z0:0 +2h

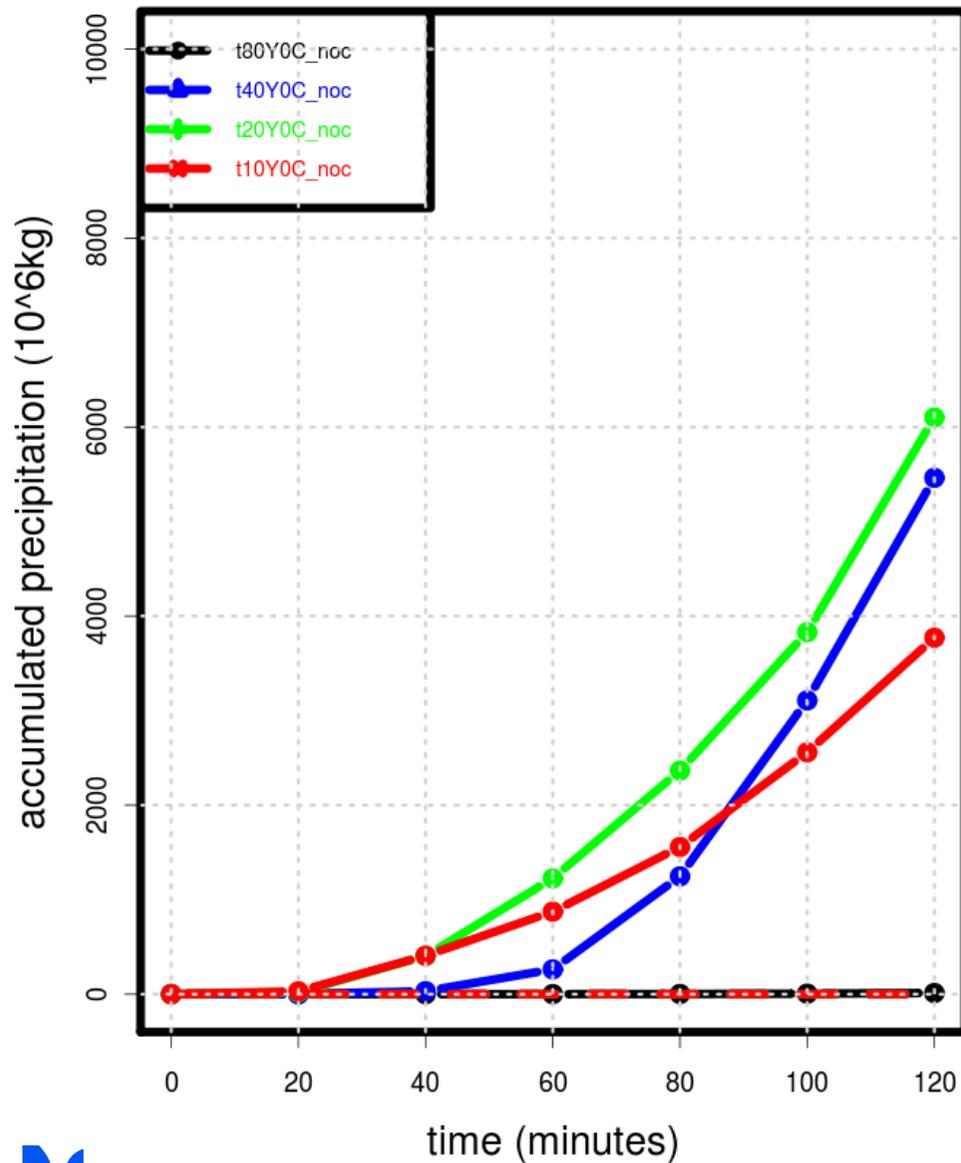


Statements

- NOCP:
 - no signal at 8km resolution
 - excessive maximum at 4km resolution
- Diagnostic parametrization:
 - Acceptable at 8km
 - Strong maximum at 4km
 - very wide extension at 4 and 2 km (also, no triggering criterion)
 - Subgrid contribution remains predominant
 - Not shown: structure and evolution in time.
 - Consequences on wind, circulation and further evolution.

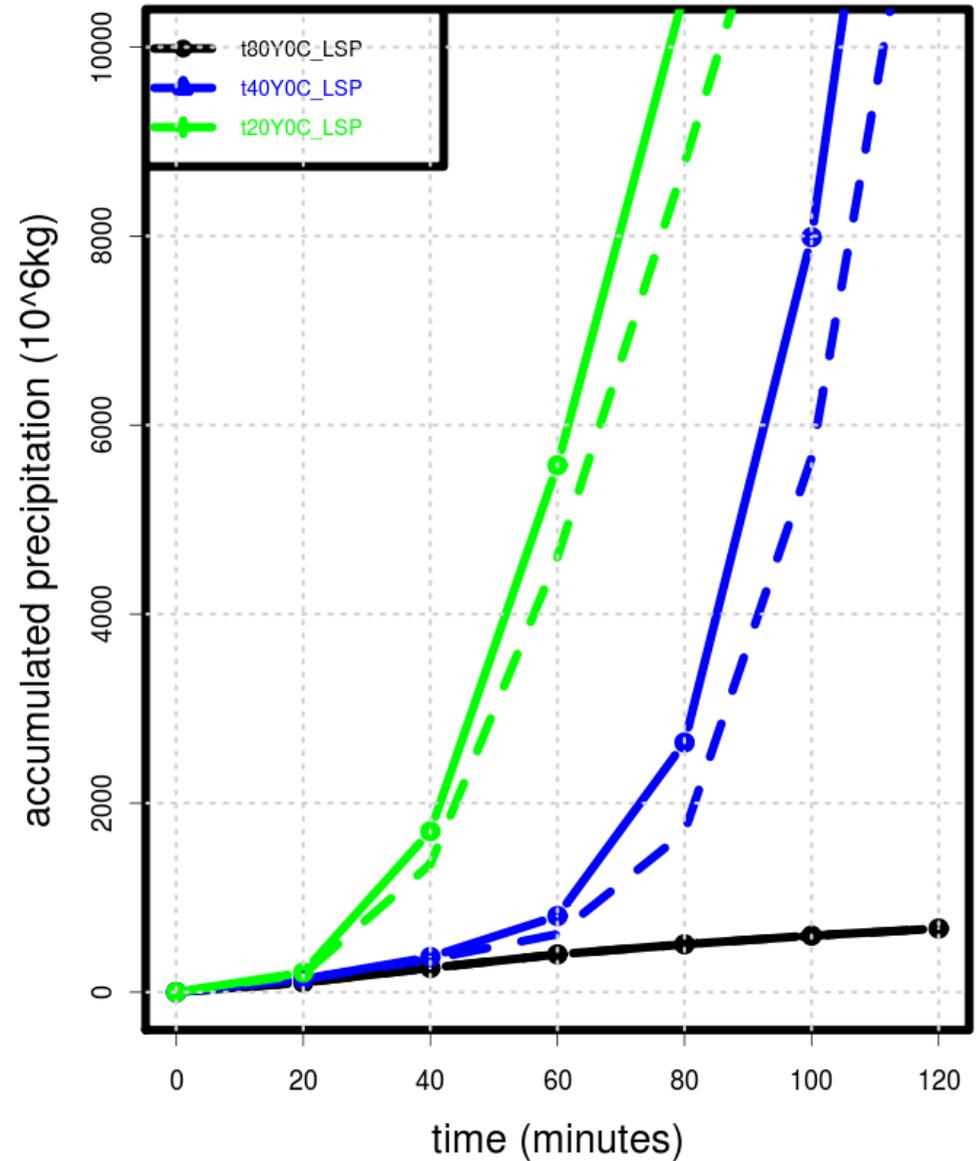
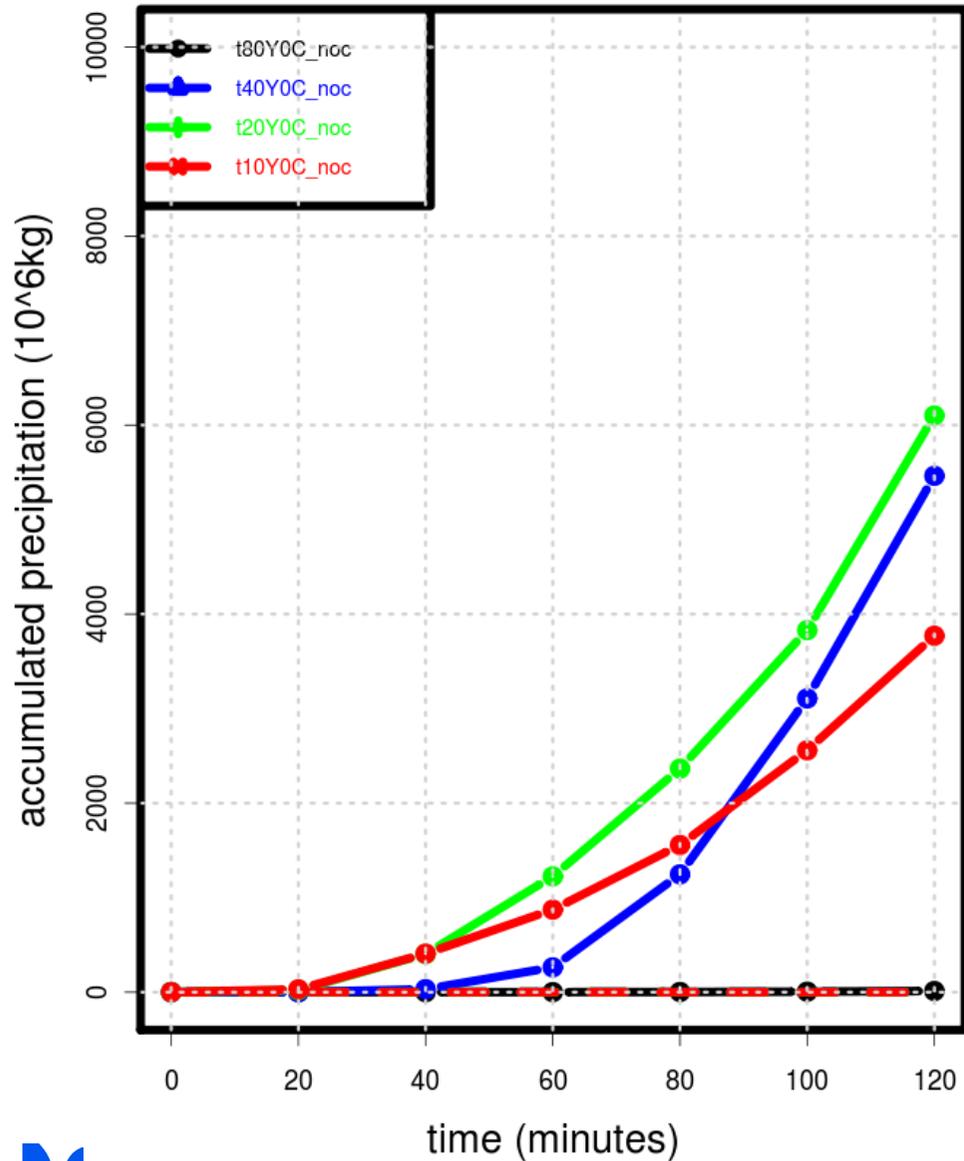
Domain-total accumulated precipitation evolution

accumulated water [10^6kg] over domain $100 \times 200 \text{km}$



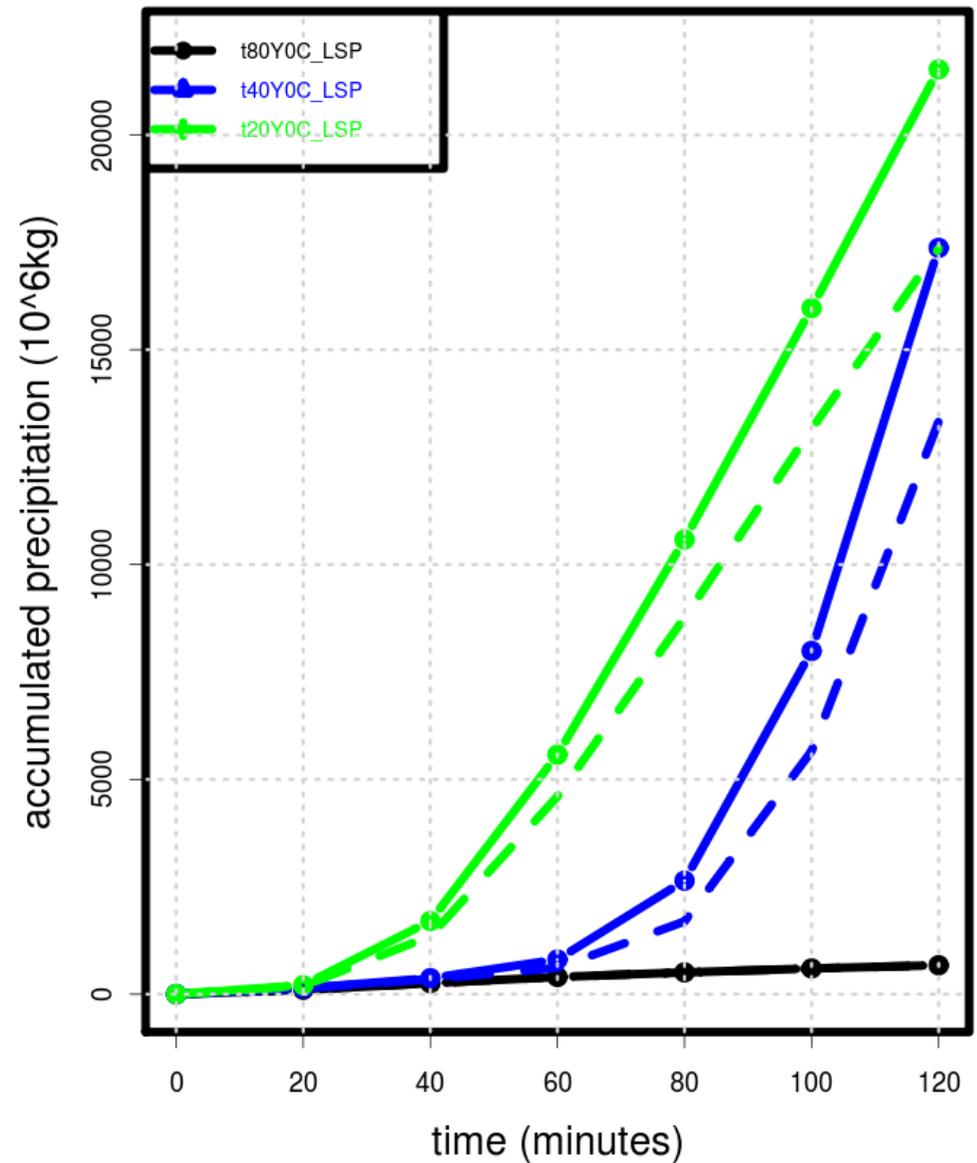
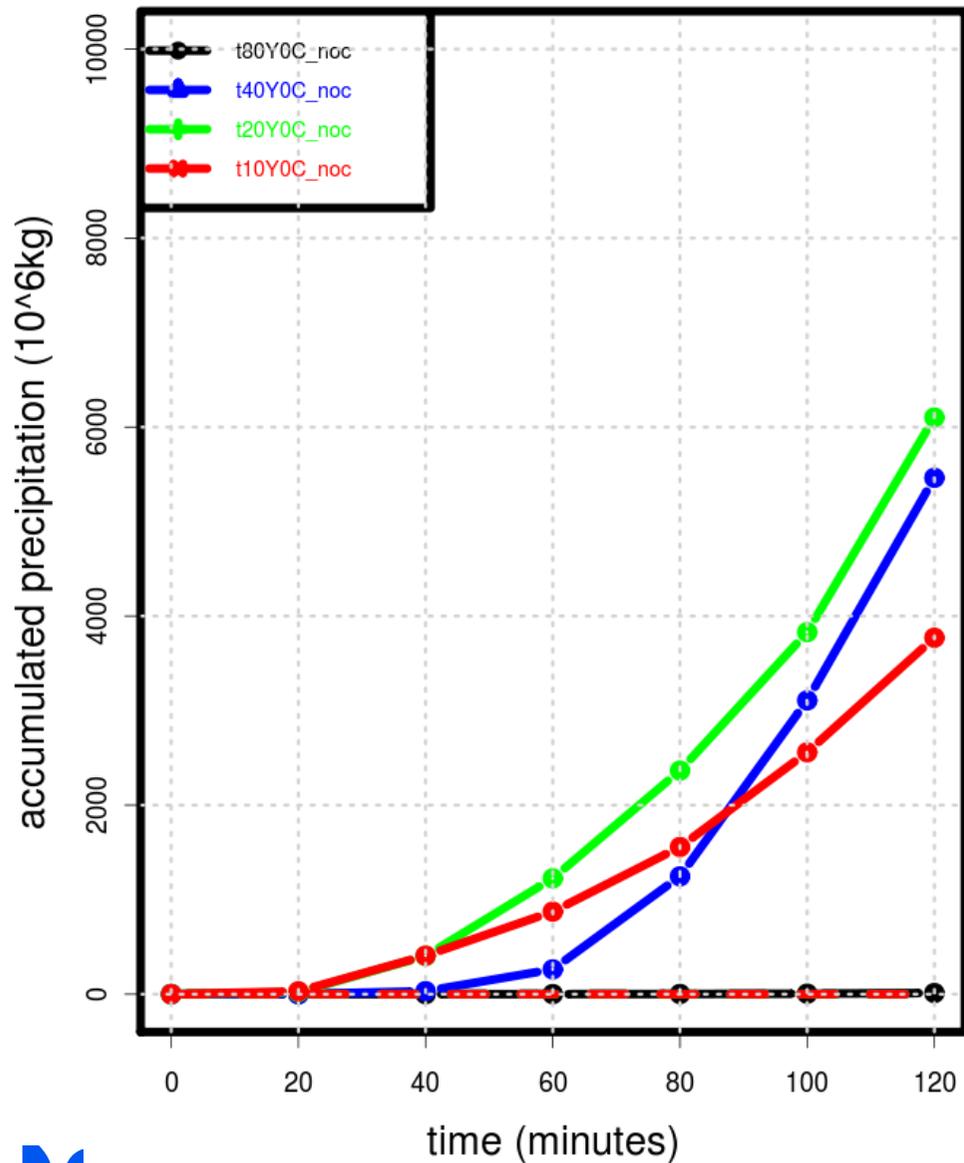
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The 3MT approach, successes and limitations

Main features:

- Sequential organization, single microphysics.
- MT concept
- Prognostic variables for updraught: ω_u, σ_u .
- cloud geometry
- Interaction between parameterizations from one time-step to the next

(Gerard *et al.* Mon. Wea. Rev. 2009.)

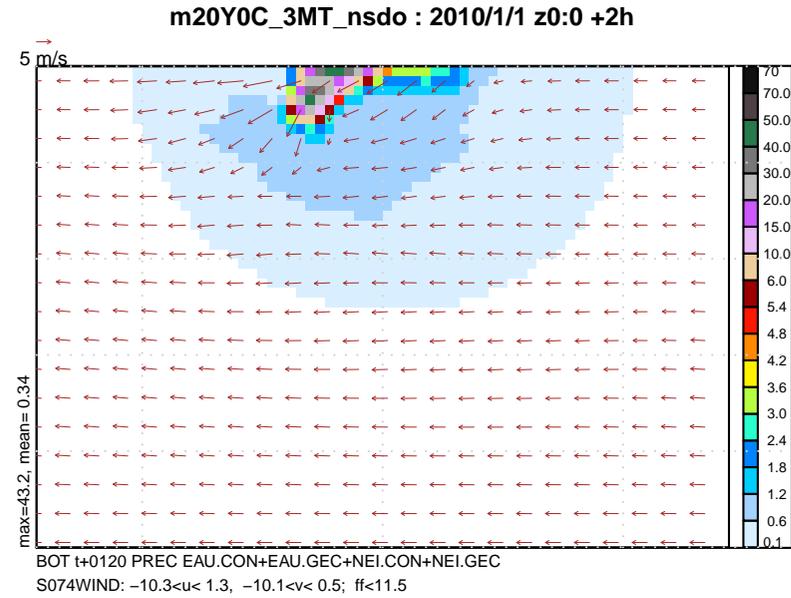
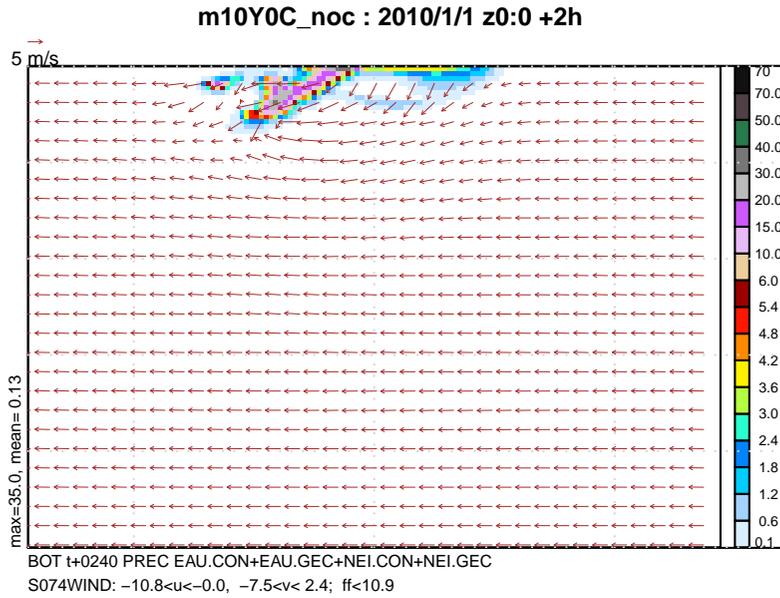
Benefits: significant model improvement, consistent forecasts at 4-km resolution.

But: while the total precipitation is kept realistic, no gradual extinction of the subgrid contribution when increasing resolution.

3MT: 2-h accumulated precipitation

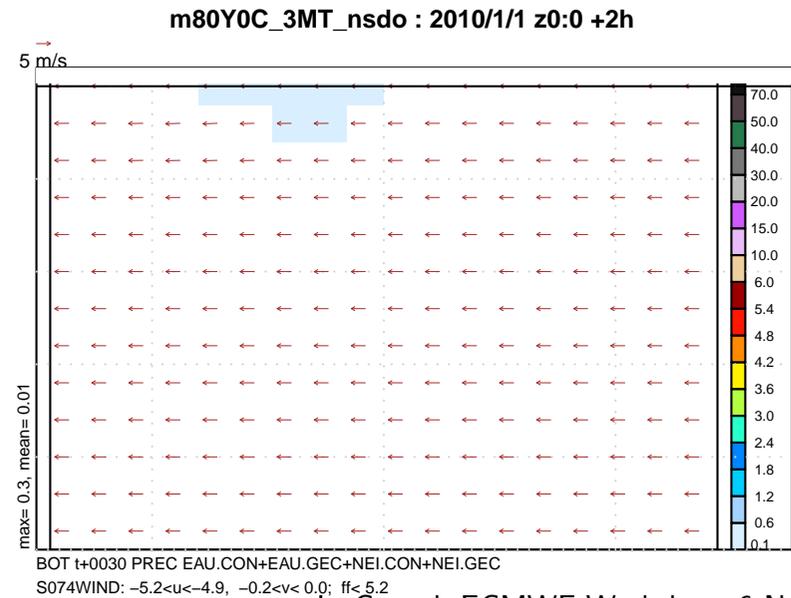
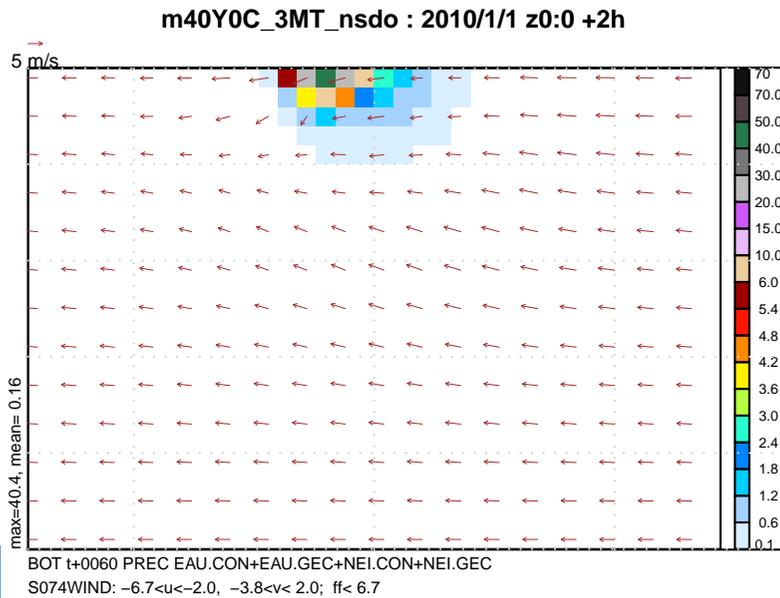
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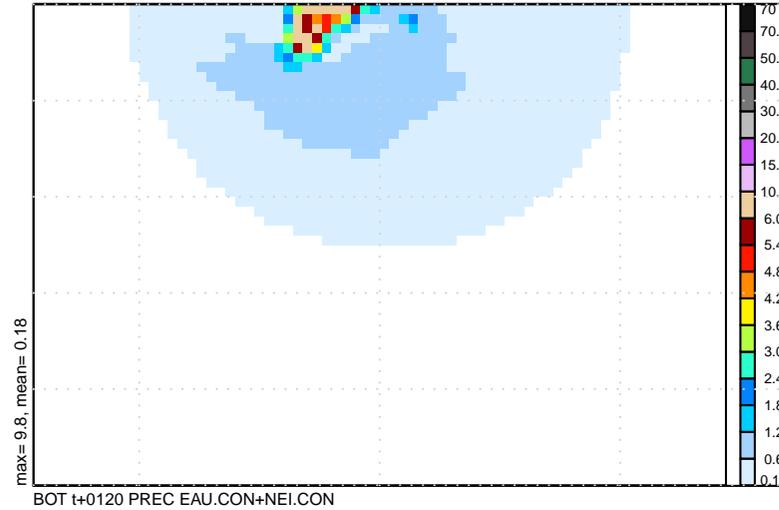


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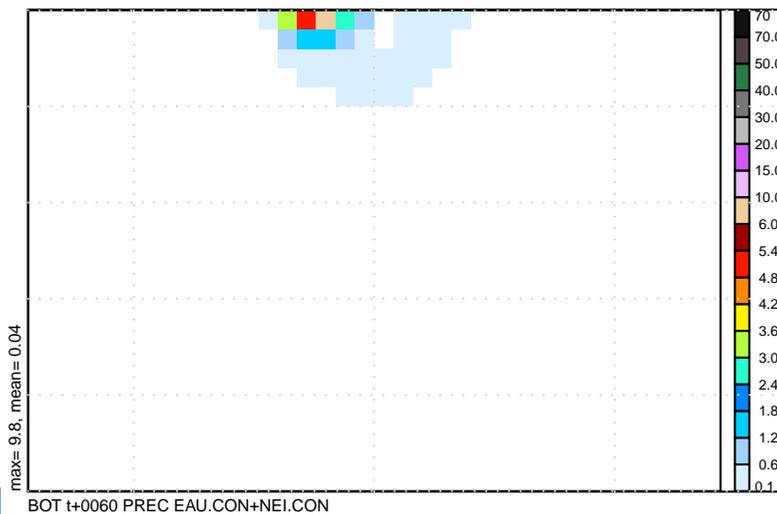
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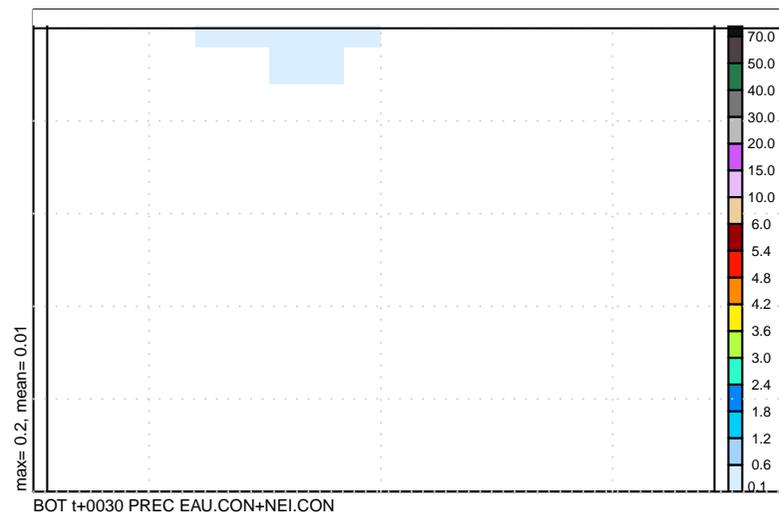
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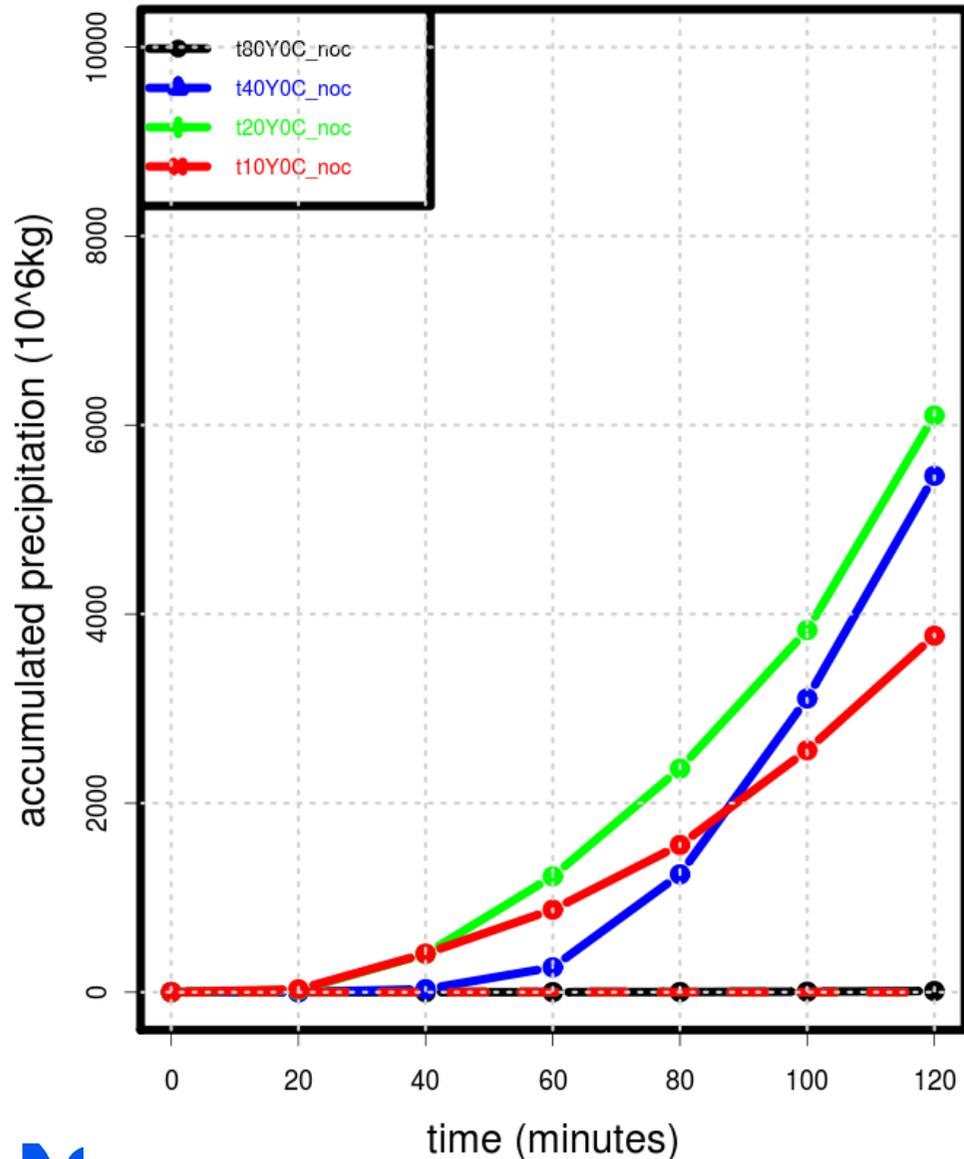
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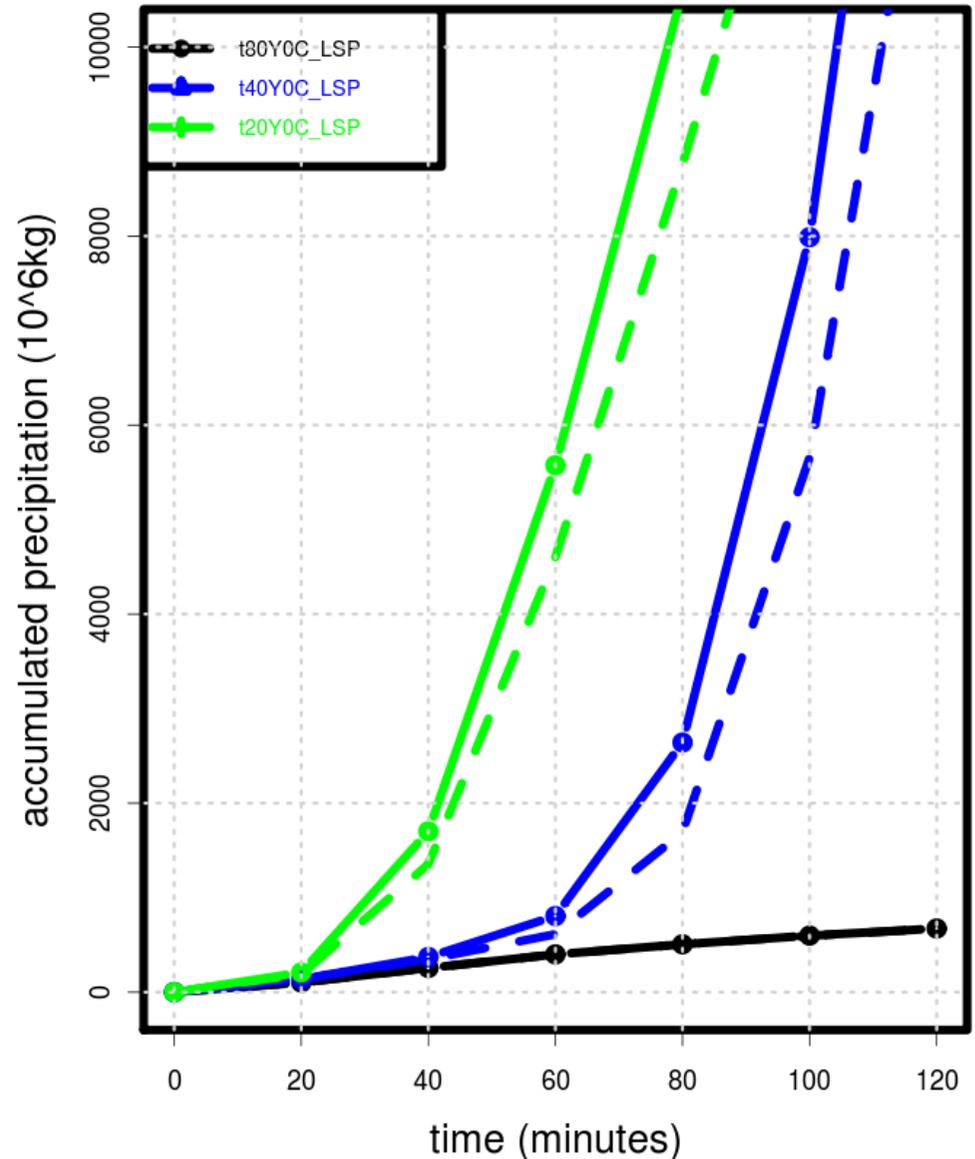
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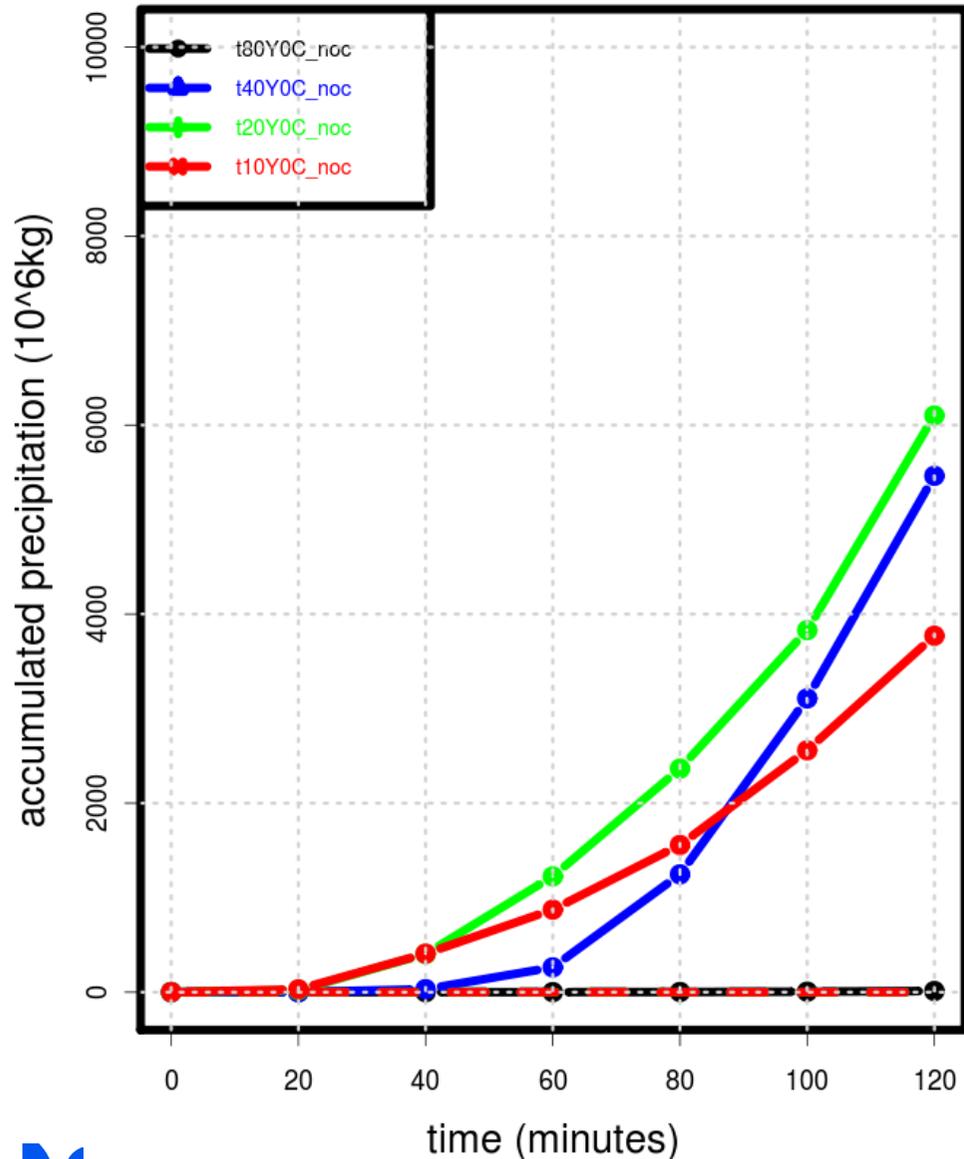
Diagnostic scheme



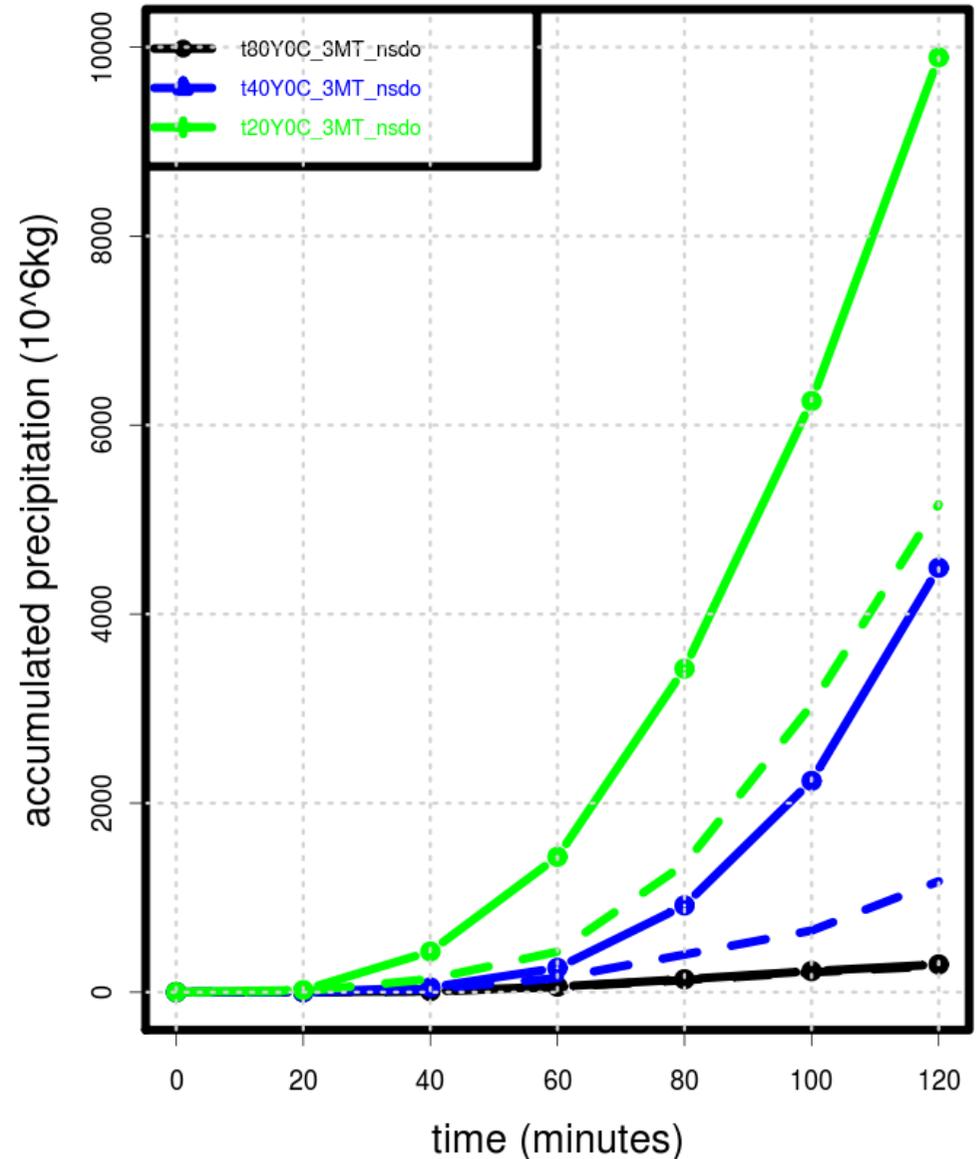
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3MT scheme



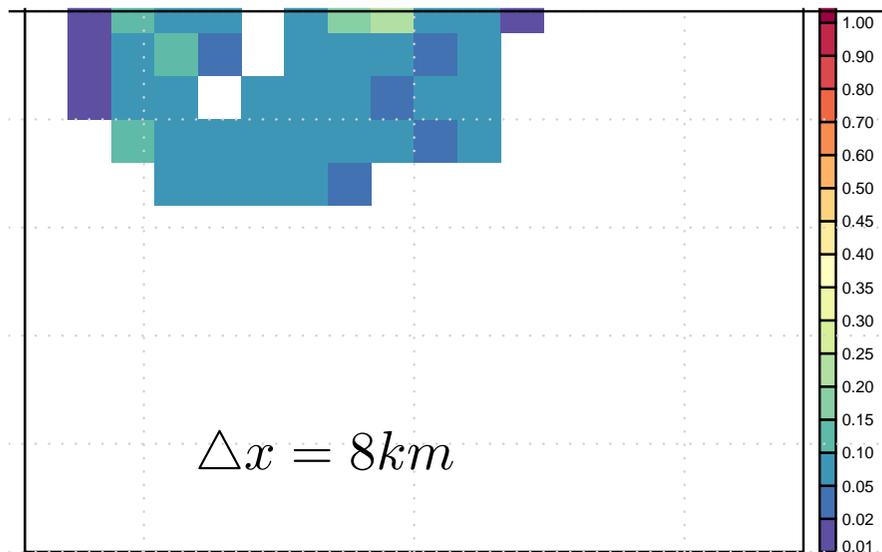
How to smoothly glide towards 100% resolved ?

Statements:

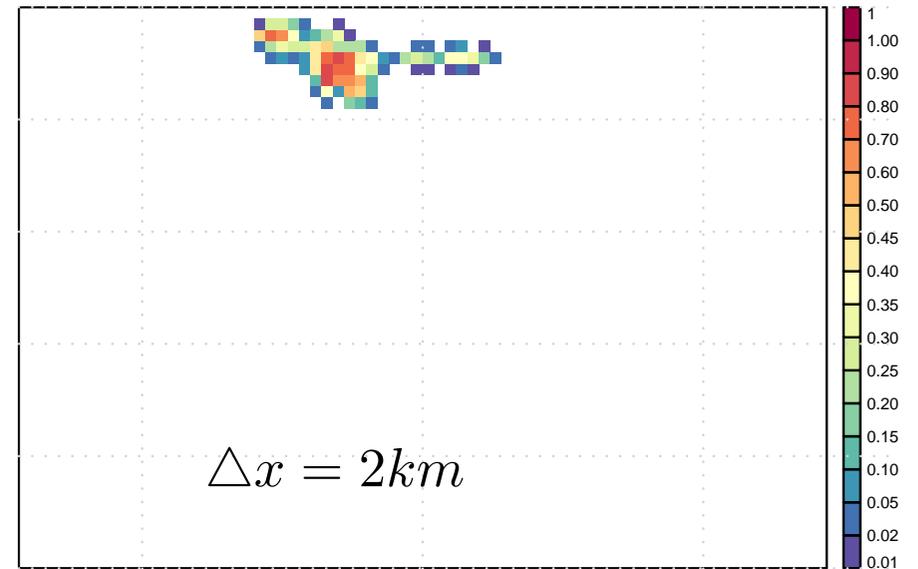
when increasing resolution, e.g. from 10km to 1km, on a region with convective activity:

- The fraction σ_u of the mesh covered by subgrid updraughts increases, depending on the granularity. It should finally tend to 1 in some of the grid boxes, when Δx small enough.

max 15-85 S015UD_MESH_FRAC
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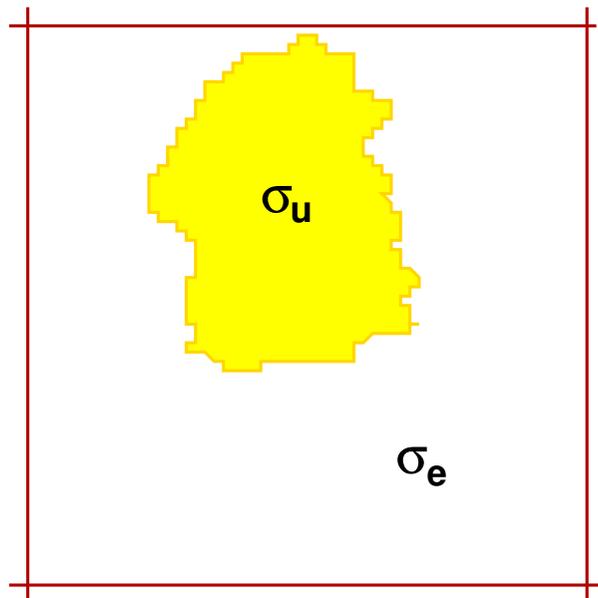
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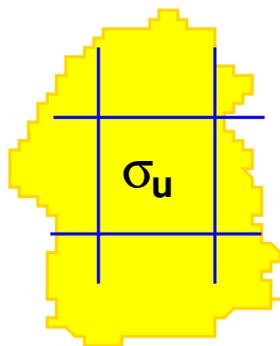


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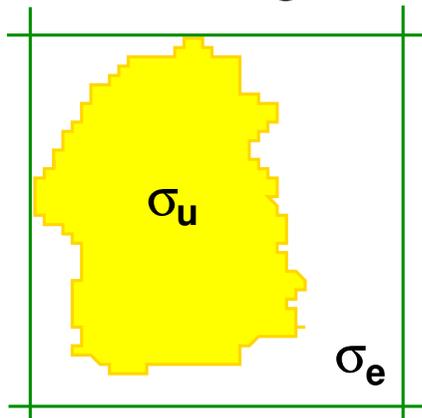


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- Triggering criteria are often badly affected: e.g. buoyancy kick increasing with resolved \overline{w} .

Complementary Subgrid Updraught

Perturbation approach: provide a complementary contribution to the part of the updraught resolved by the mean grid-box: $\psi' = (\psi_u - \bar{\psi})$

- 3MT sequential organization.
- Ascent properties obtained from the anelastic equations. The CSU contribution is confined in the grid column: effects of mesh fraction, of the resolved profile.
- Updraught evolution: prognostic perturbation velocity, evolving mesh fraction, gradually rising top.
- Closure referring to steady state and estimation of 'real world' CAPE.
- CSU triggering: threshold of resolved condensation (together with updraught viability). Threshold $\propto (\Delta x)^{-2}$.

Triggering at 'grey zone' resolutions

- Kain-Fritsch (2004):

$$\Delta T_{v,KF} = \left[\gamma (\bar{w}_{LCL} - w_0 \min(1, \frac{z_{LCL}}{z_0})) \right]^{1/3}, \quad \frac{1}{\gamma} \sim 0.01 \text{m s}^{-1} \text{K}^{-3}, \quad z_0 = 2 \text{km},$$

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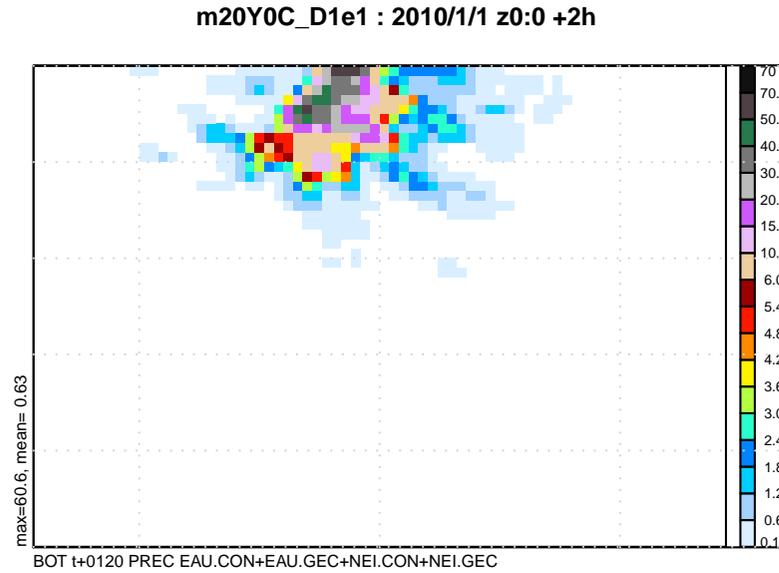
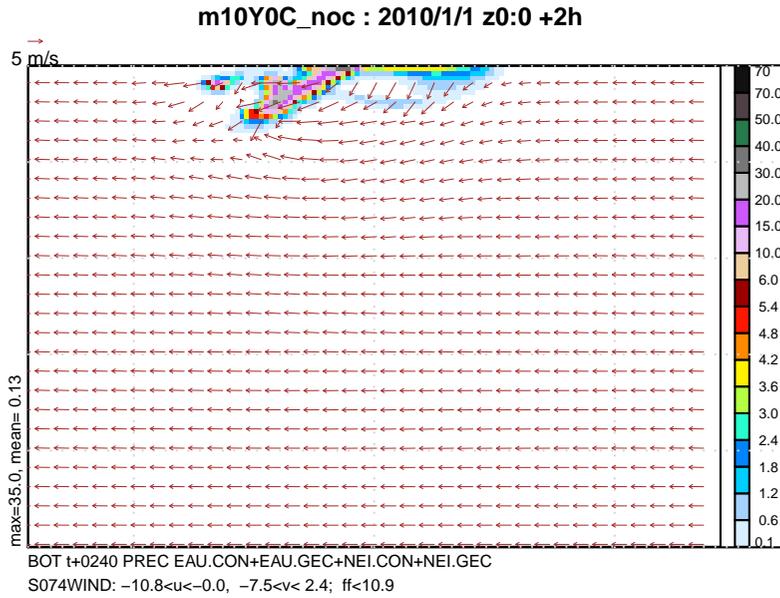
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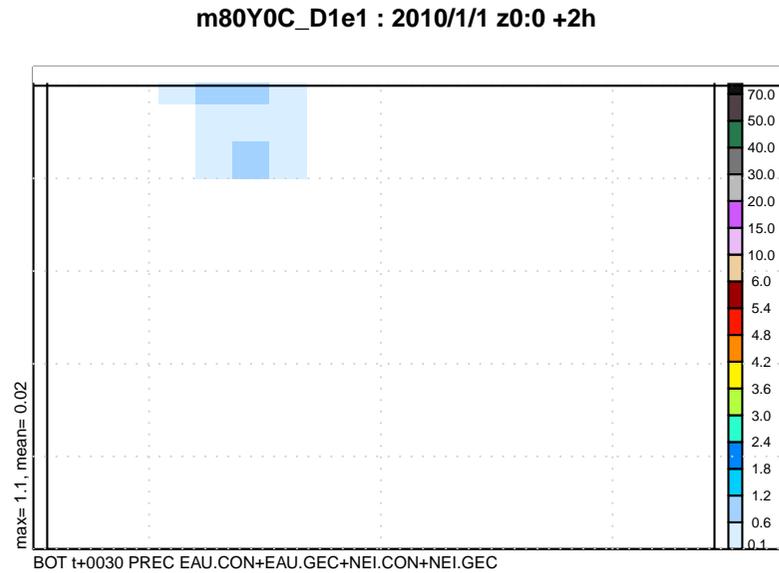
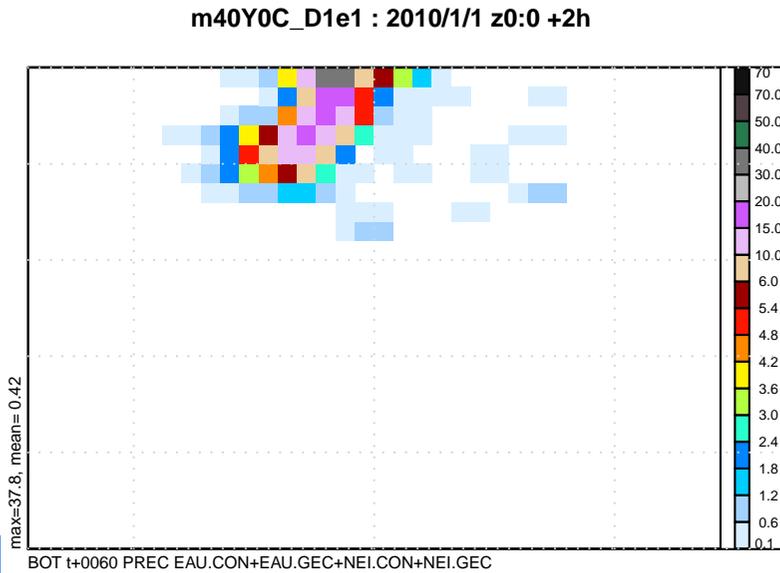
total

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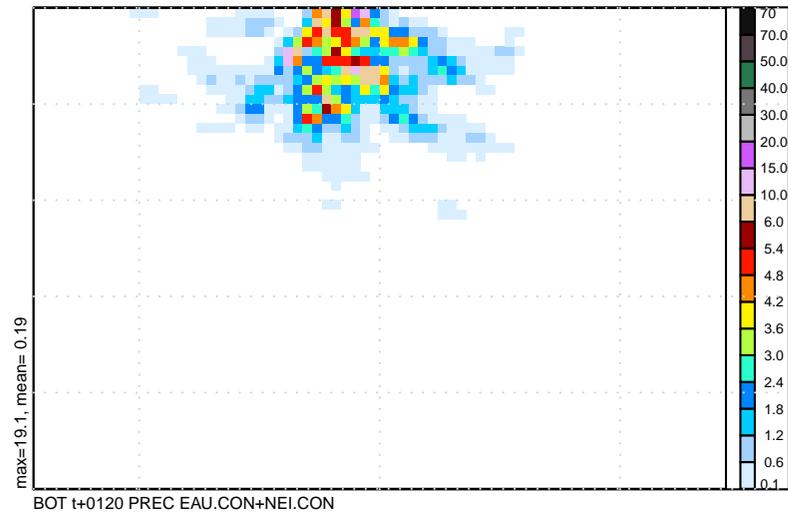


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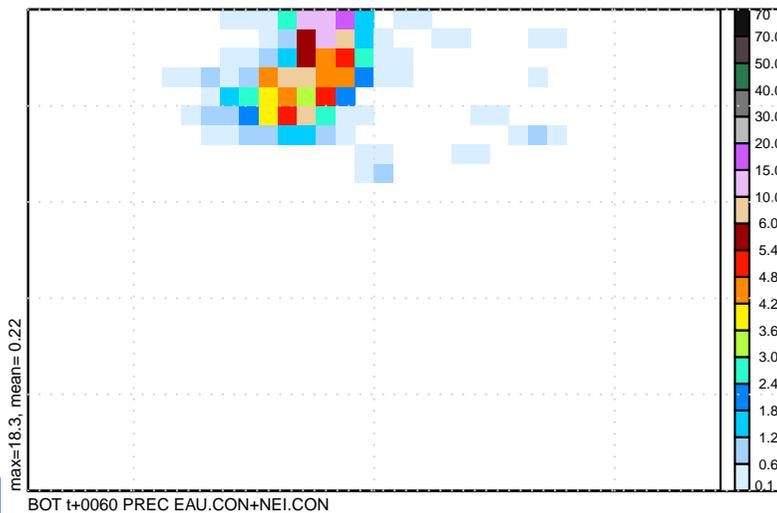
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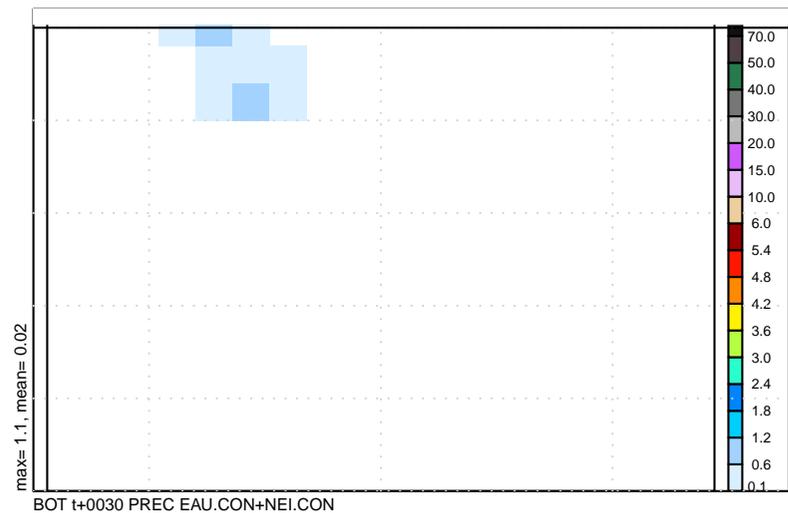
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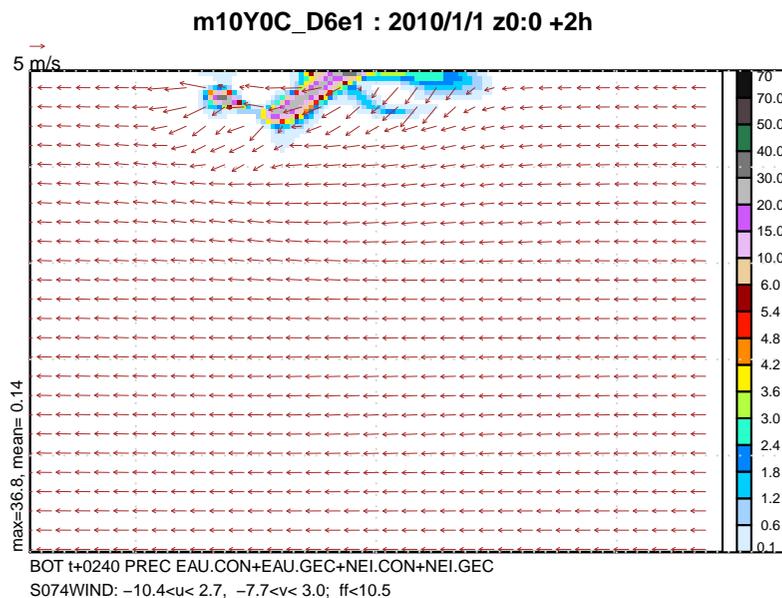
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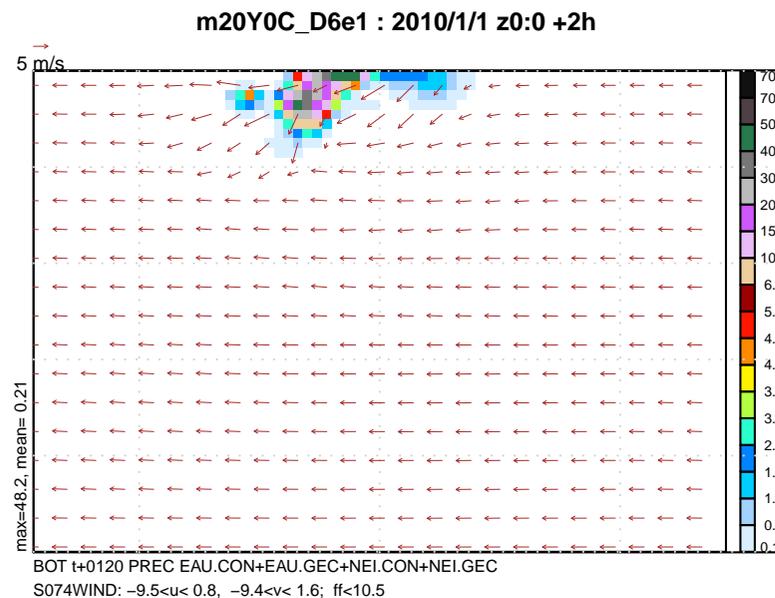
CSU: 2-h accumulated precipitation

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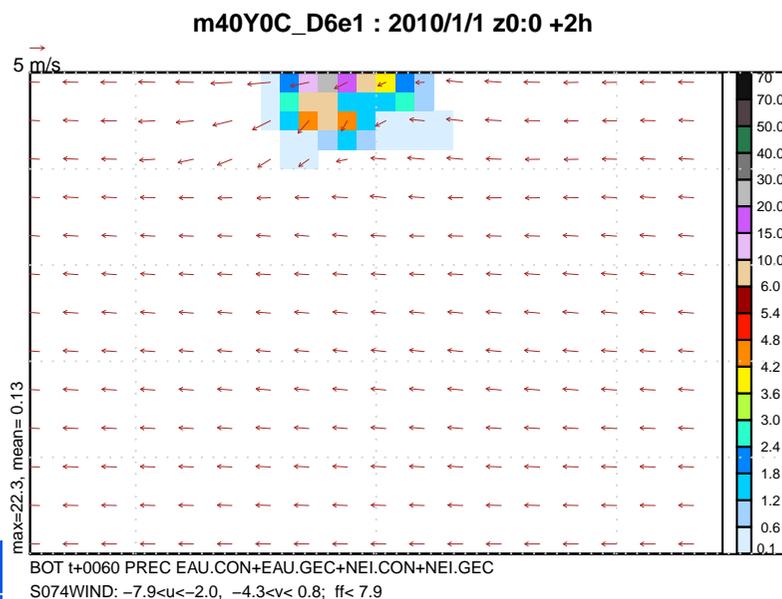
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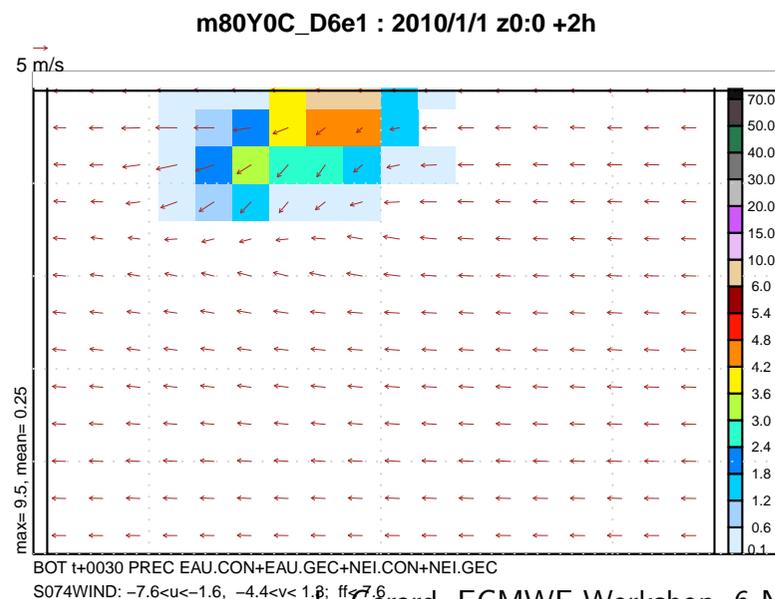
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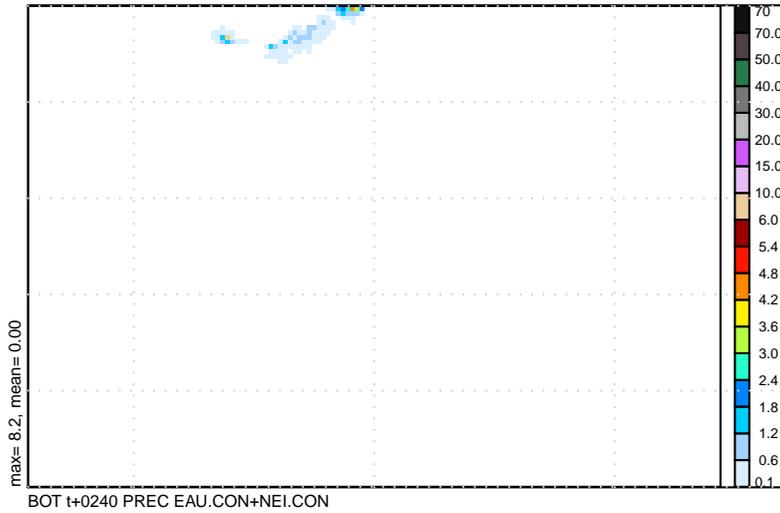


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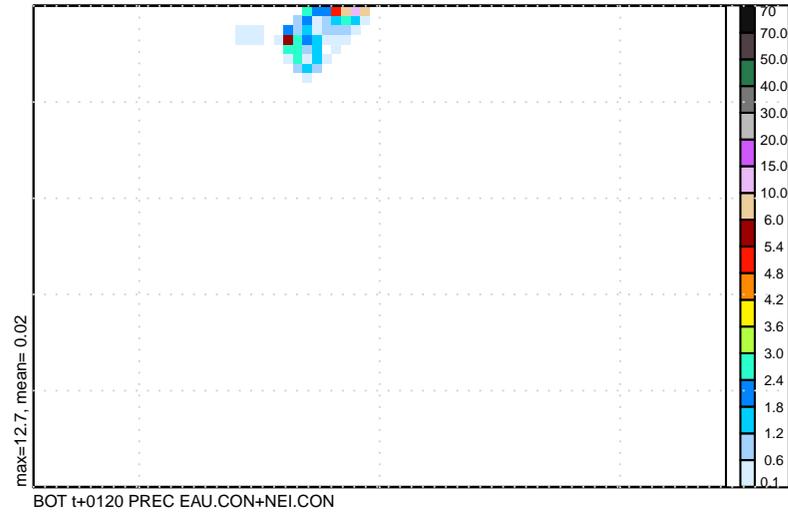
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$\Delta x = 1km$

m10Y0C_D6e1 : 2010/1/1 z0:0 +2h



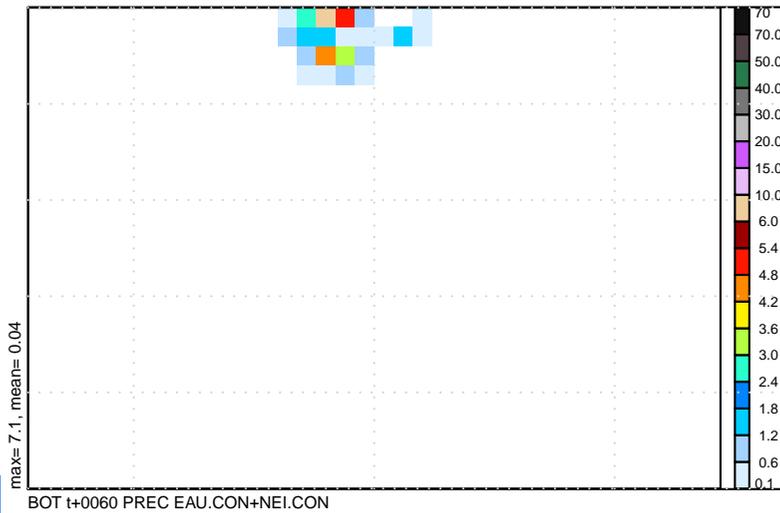
m20Y0C_D6e1 : 2010/1/1 z0:0 +2h



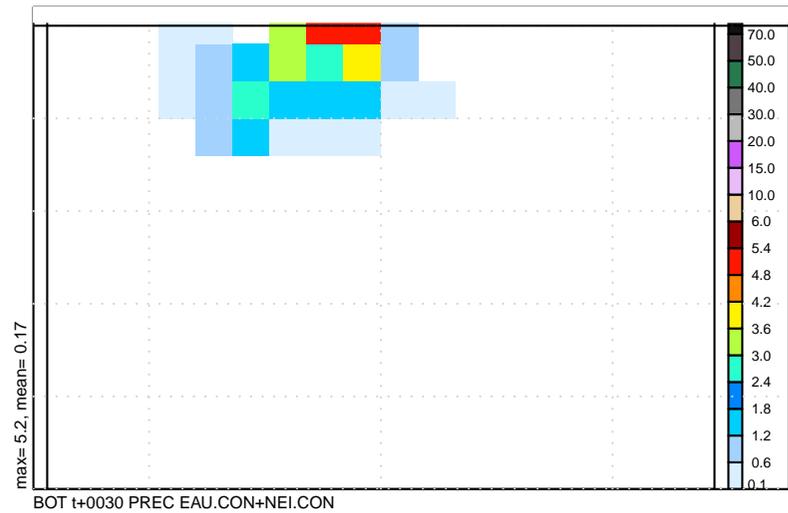
$\Delta x = 2km$

$\Delta x = 4km$

m40Y0C_D6e1 : 2010/1/1 z0:0 +2h



m80Y0C_D6e1 : 2010/1/1 z0:0 +2h

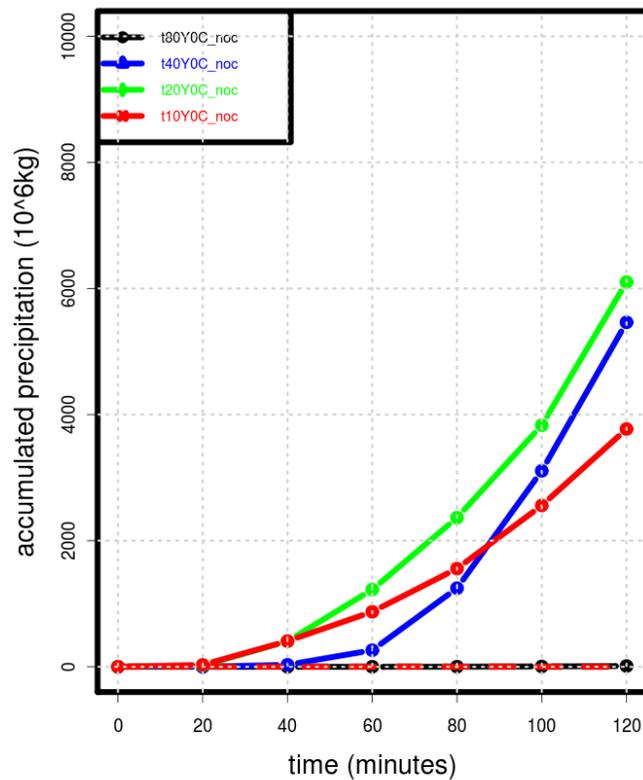


$\Delta x = 8km$



Domain-total accumulated precipitation evolution

accumulated water [$10^6 kg$] over domain 100x200km

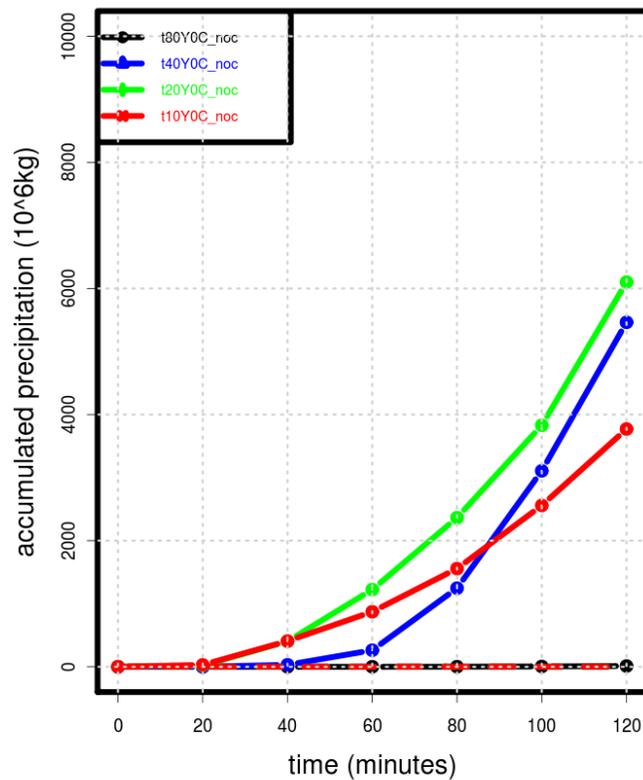


no conv param.

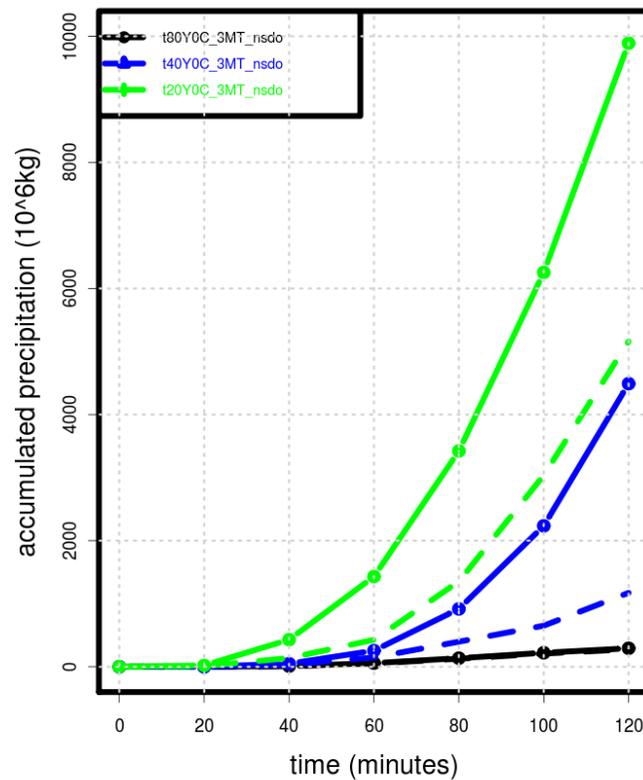
1km, 2km, 4km, 8km.

Domain-total accumulated precipitation evolution

accumulated water [$10^6 kg$] over domain 100x200km



no conv param.



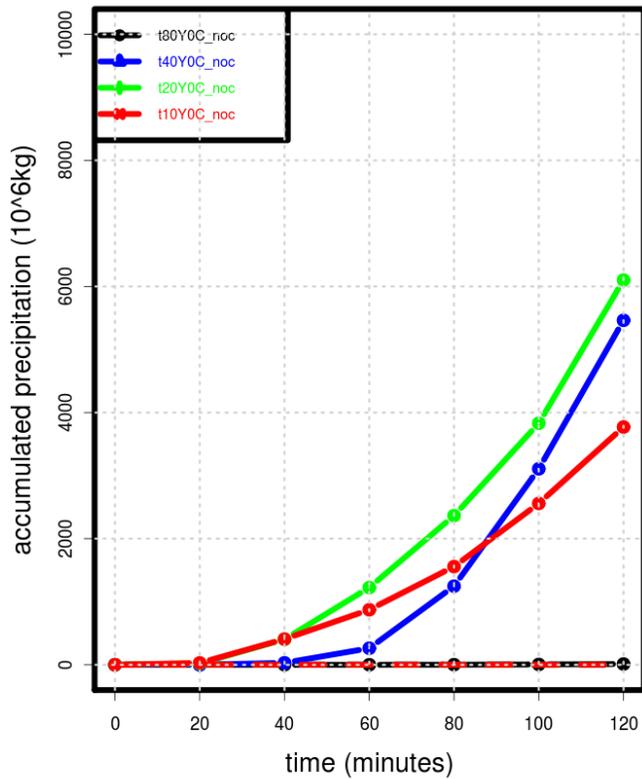
3MT

1km, 2km, 4km, 8km.

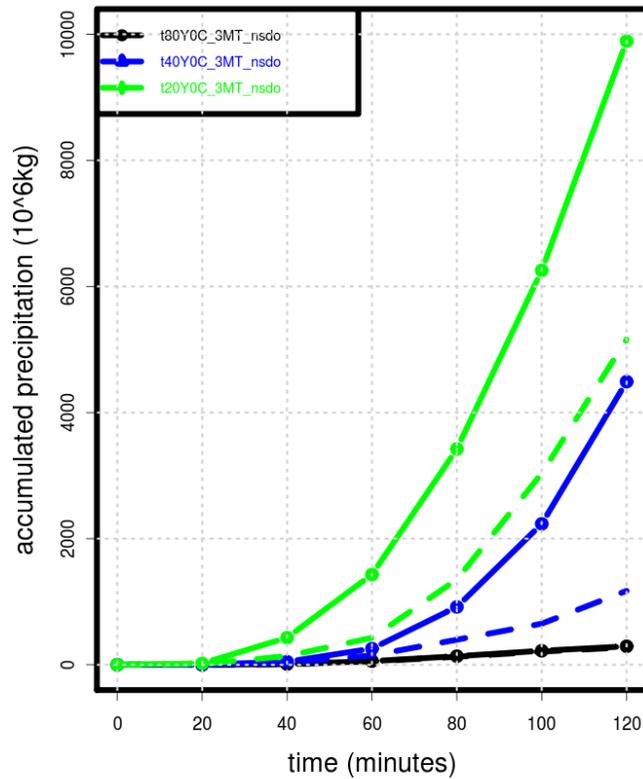


Domain-total accumulated precipitation evolution

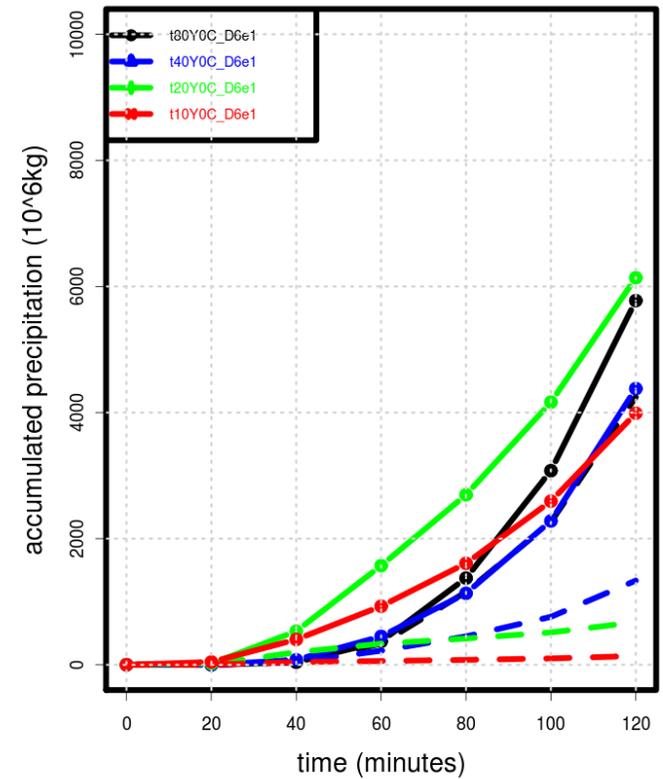
accumulated water [$10^6 kg$] over domain 100x200km



no conv param.



3MT



CSU

1km, 2km, 4km, 8km.



Real case

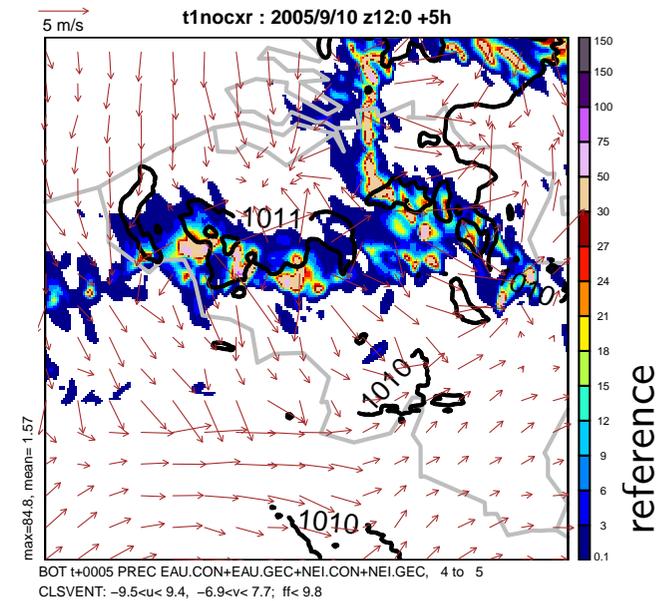
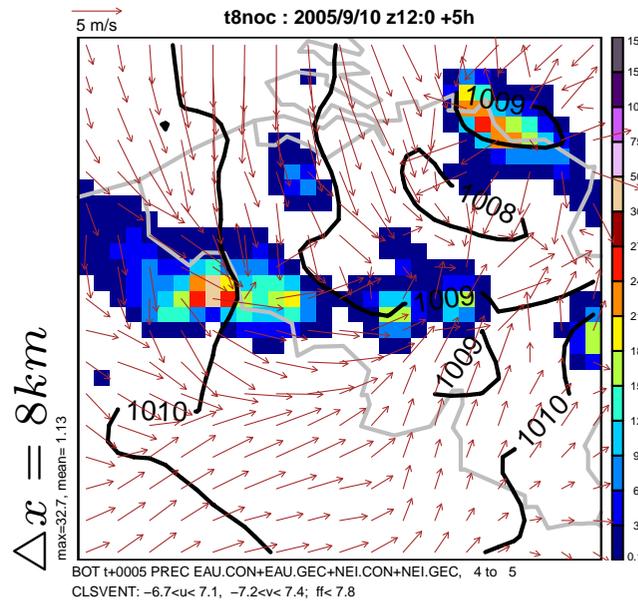
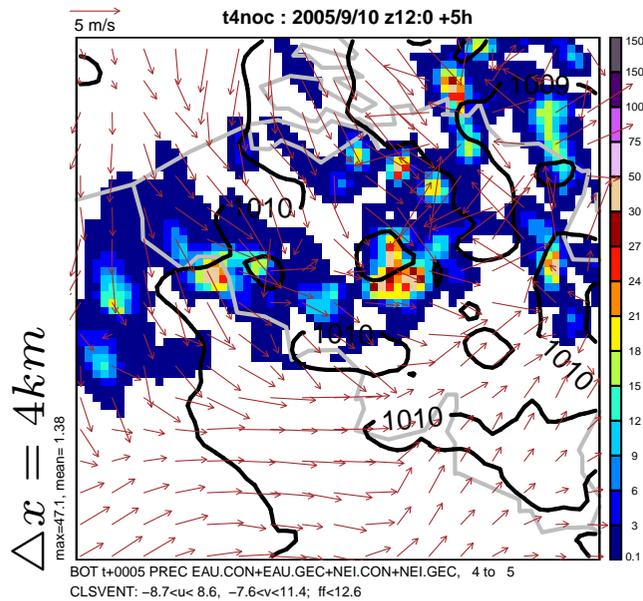
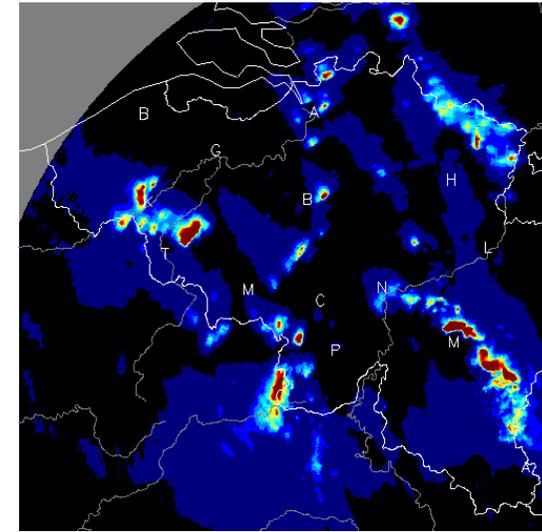
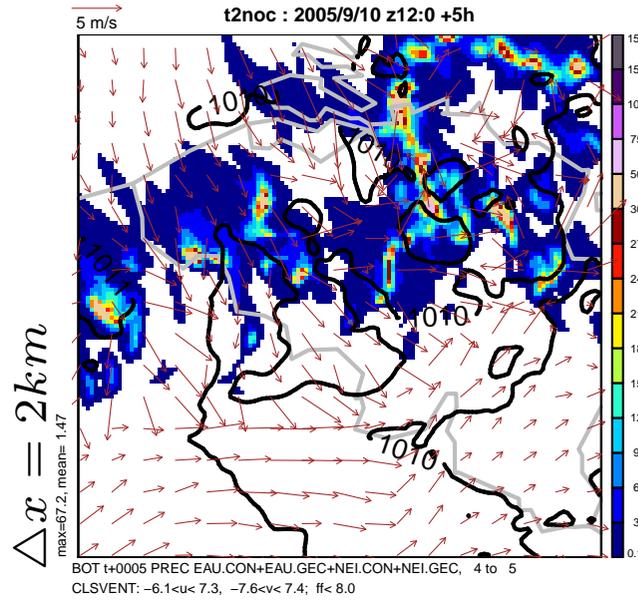
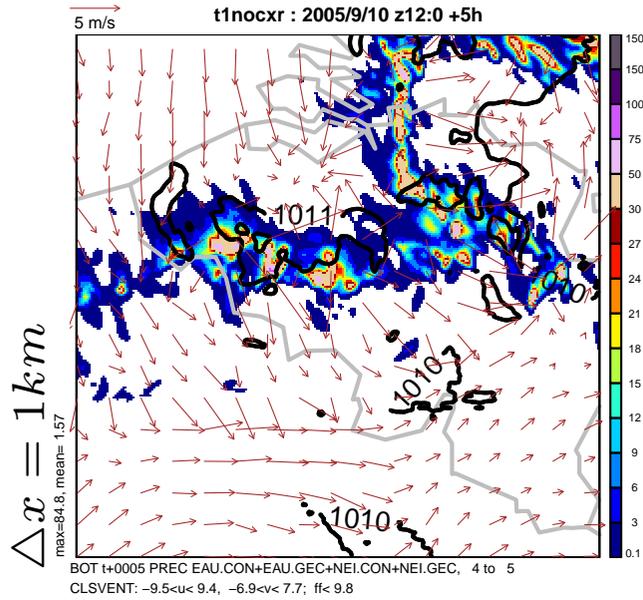
thunderstorms over Belgium on 10 September 2005

Hydrostatic model at 8km and 4km, non-hydrostatic at 2 and 1km

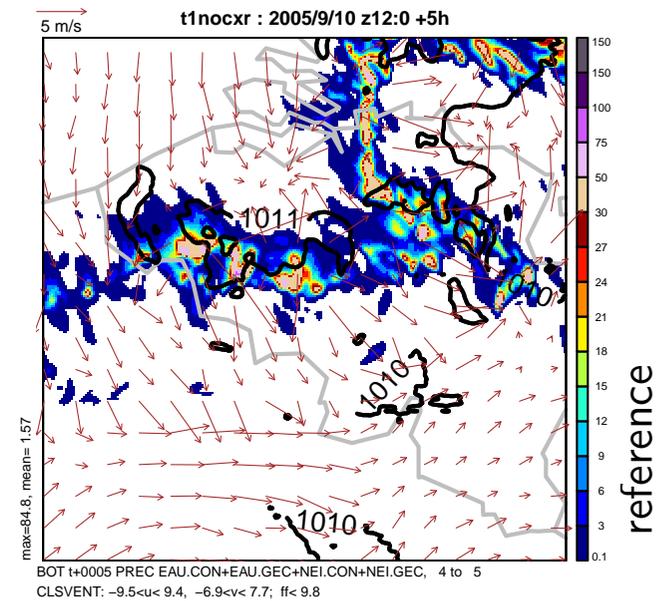
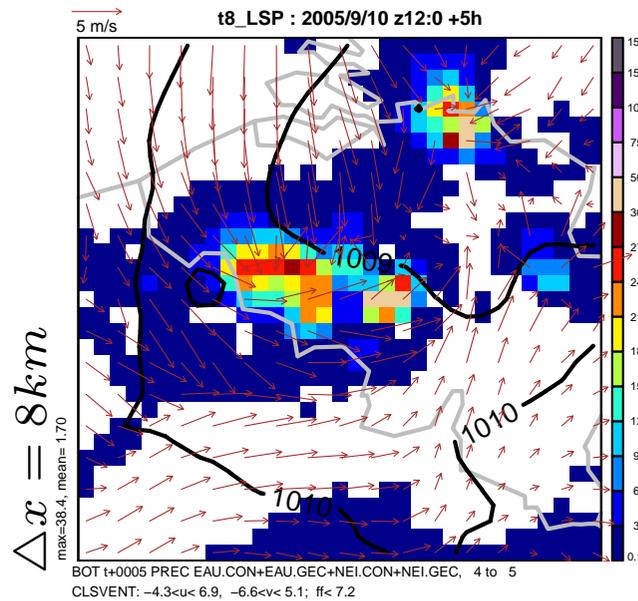
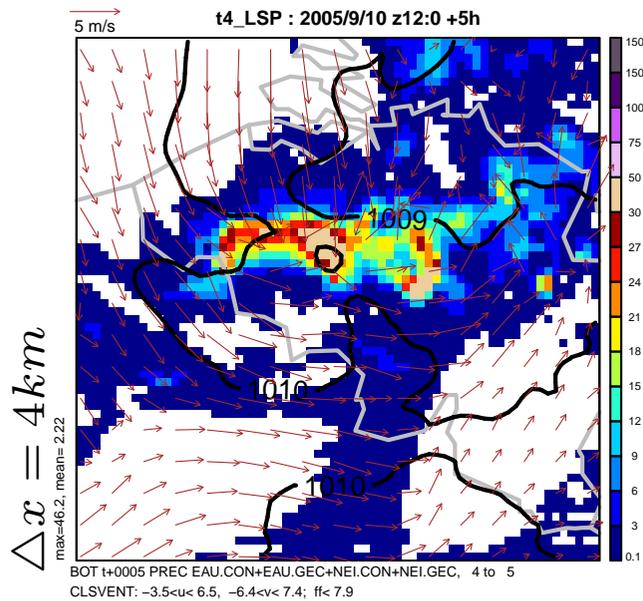
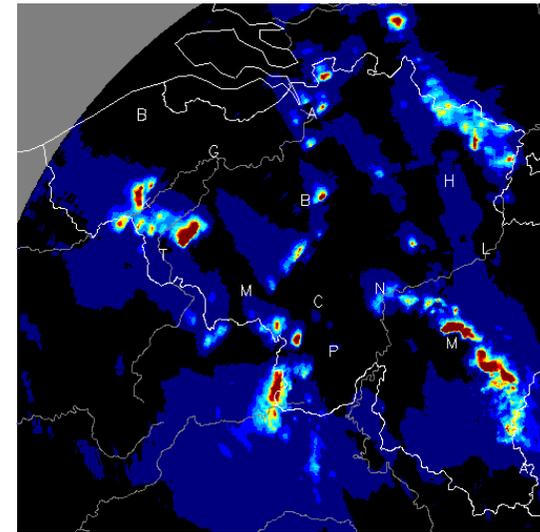
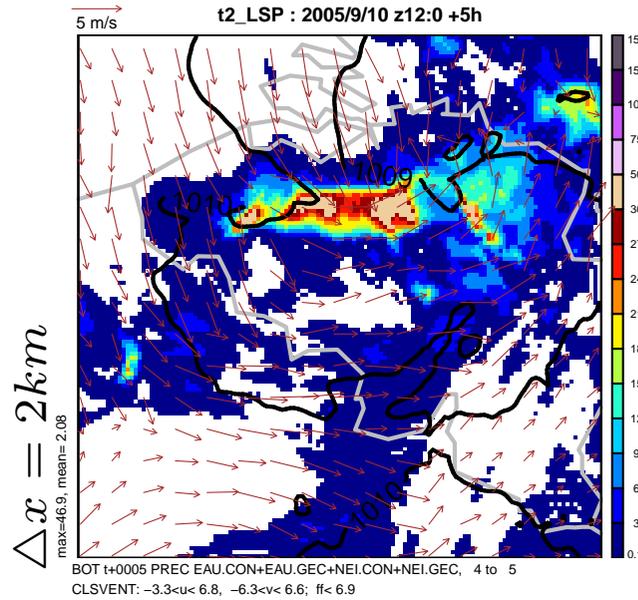
41 vertical hybrid levels

Remark: Belgian territory is quite flat.

No cp: 1h-accumulated precipitation

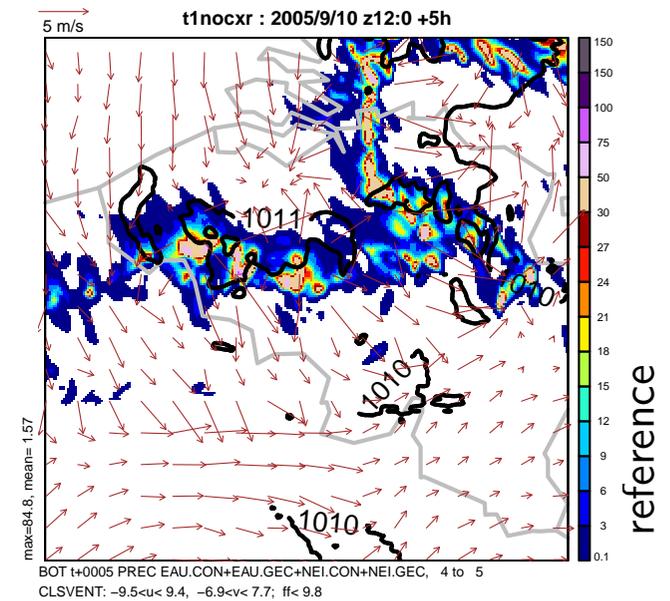
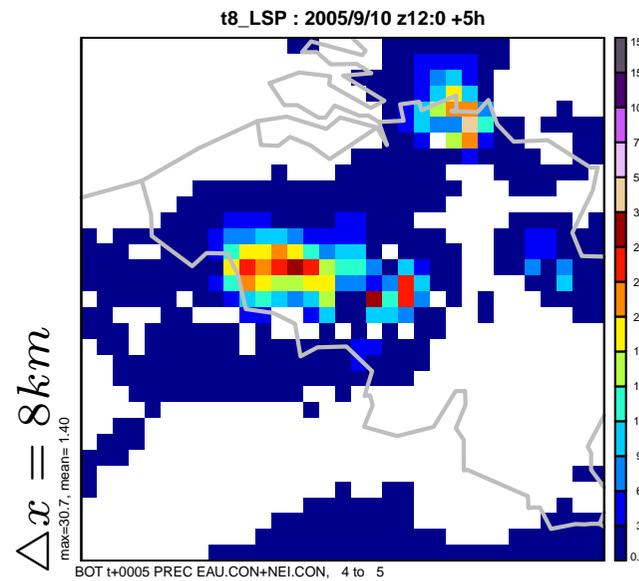
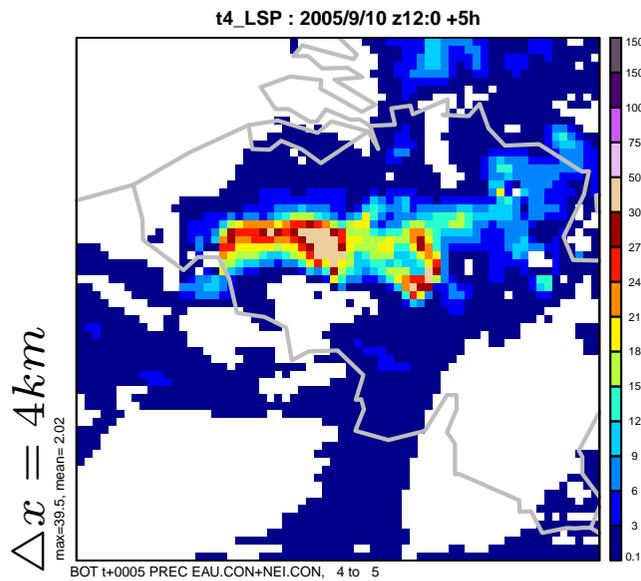
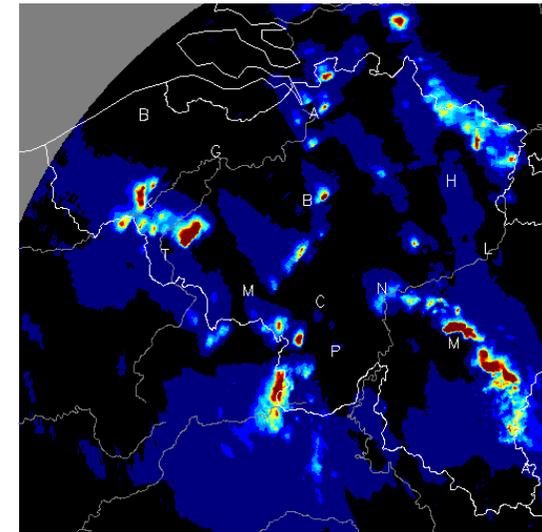
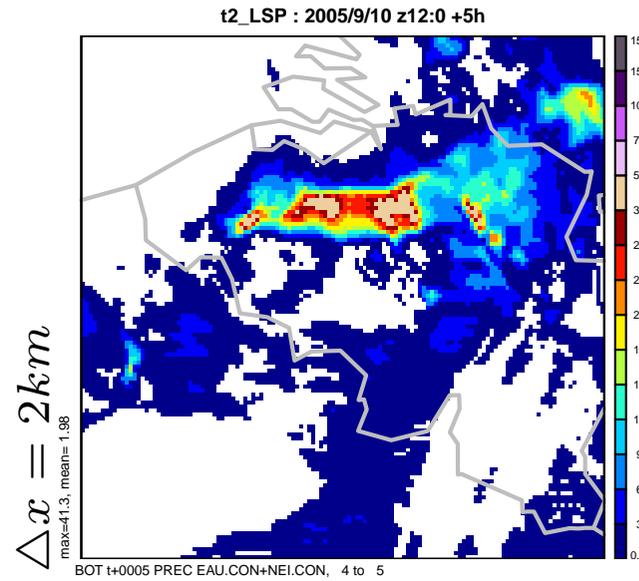


Diag param: 1h-accumulated precipitation

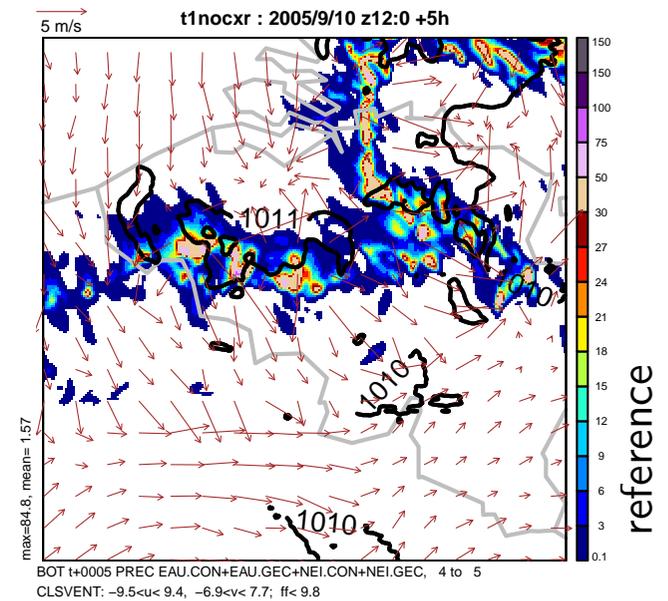
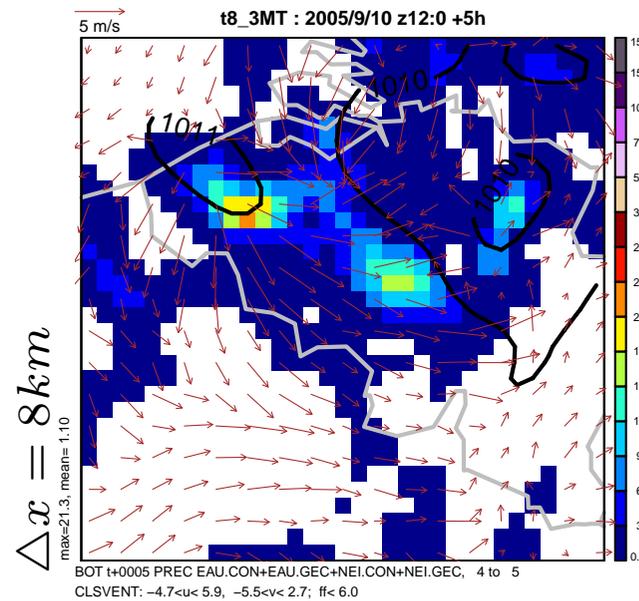
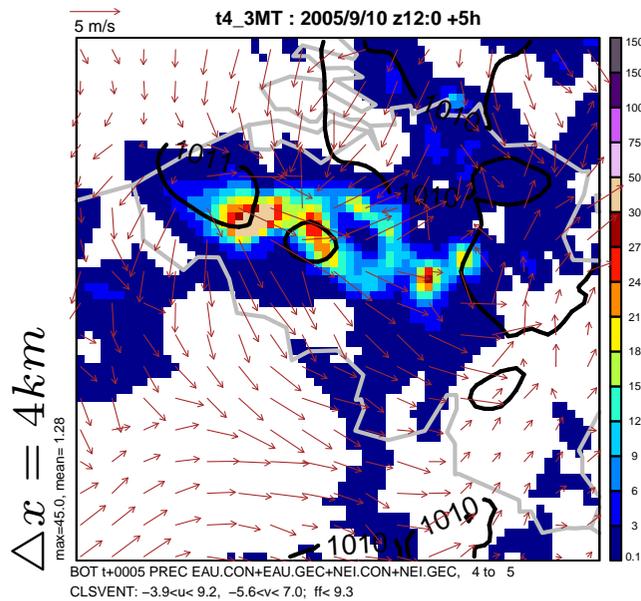
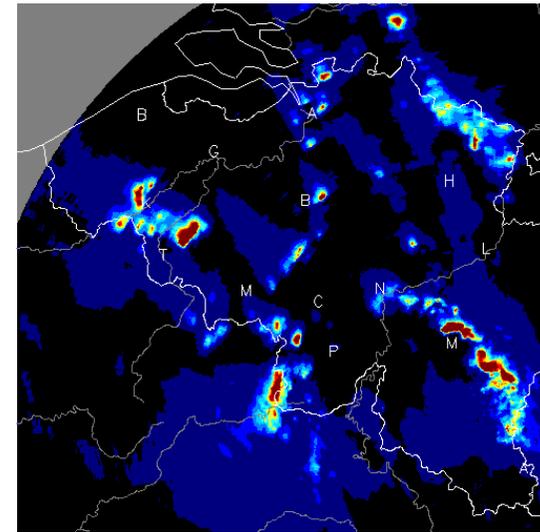
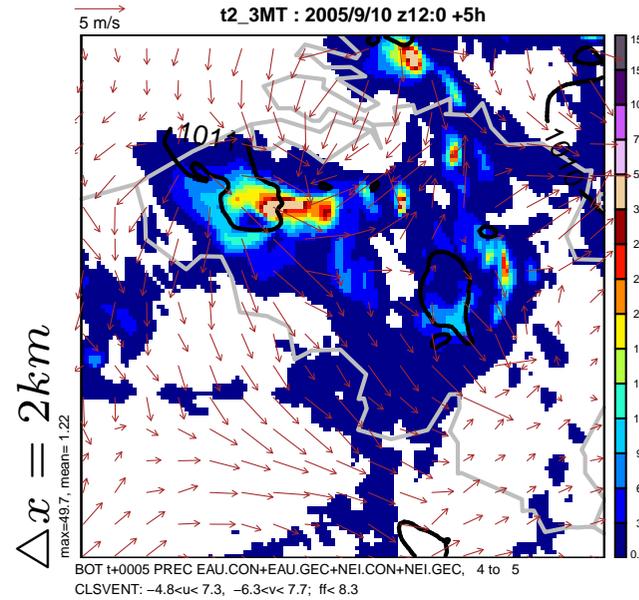


total

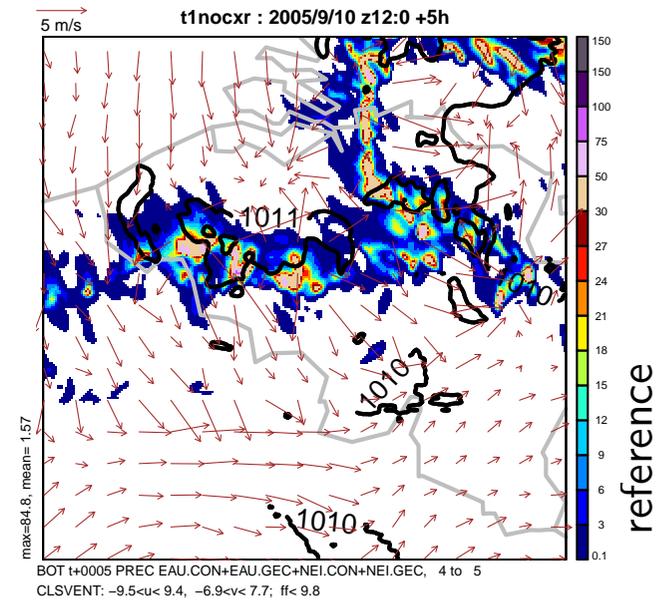
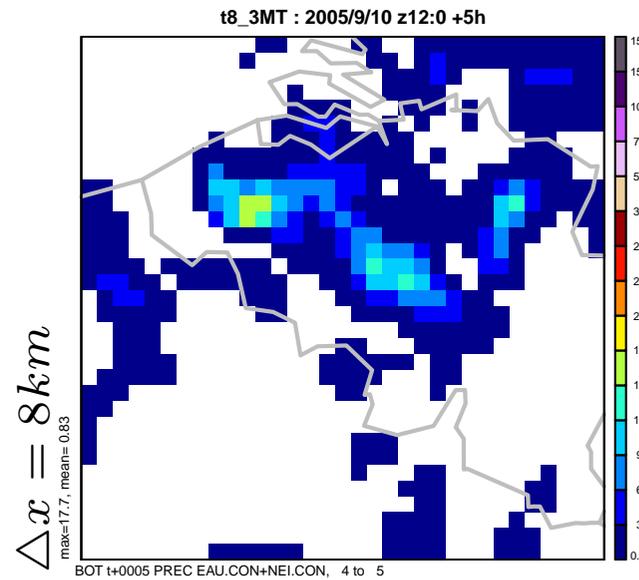
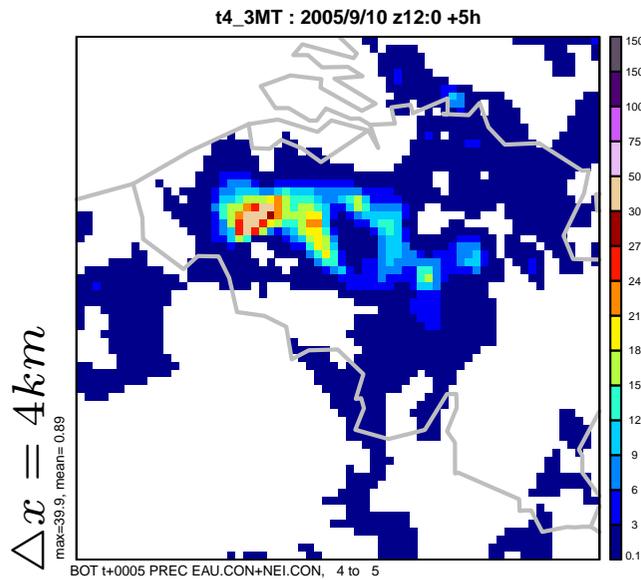
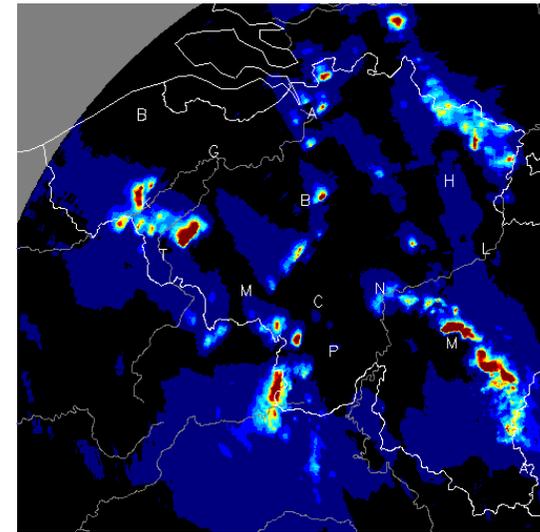
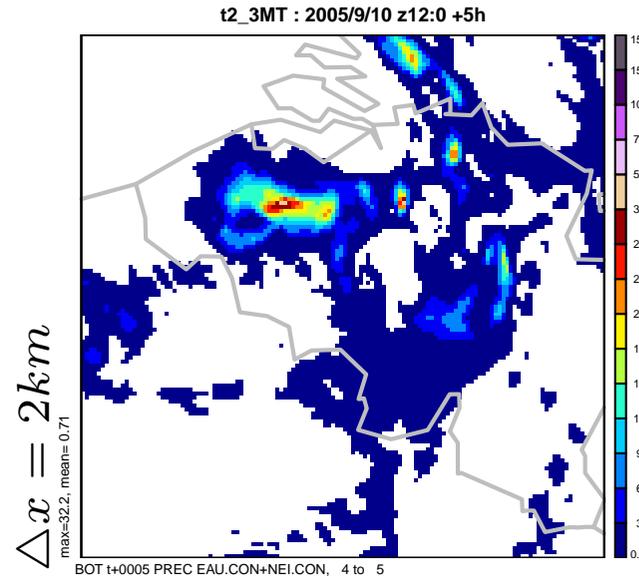
Diag param: 1h-accumulated precipitation



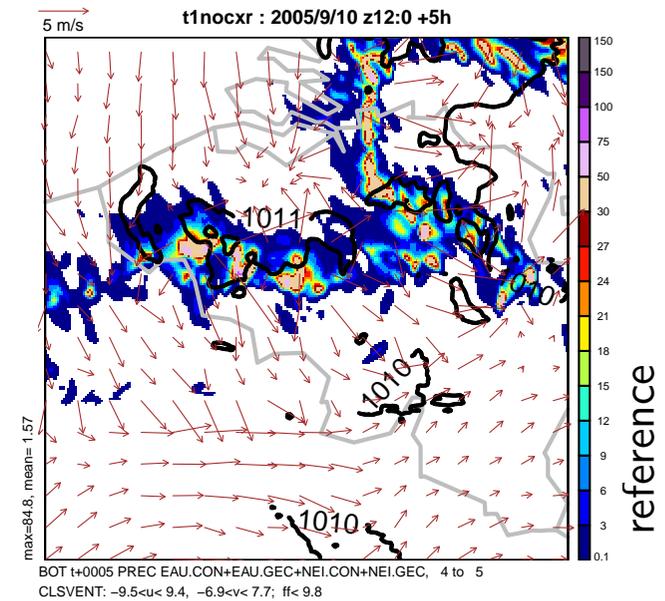
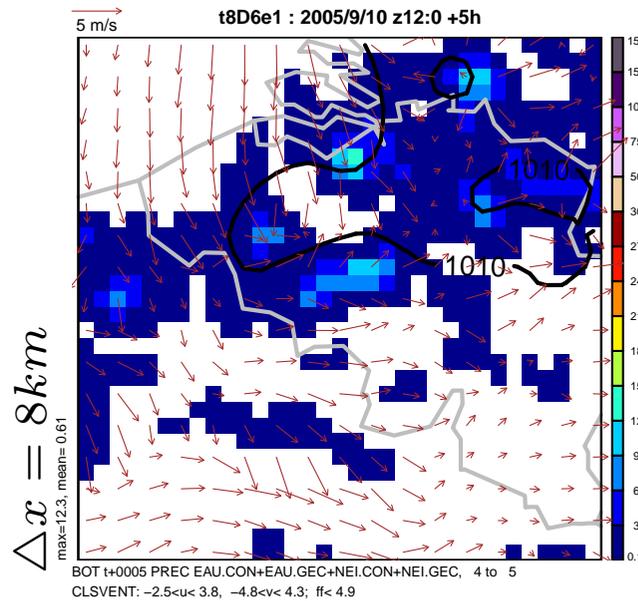
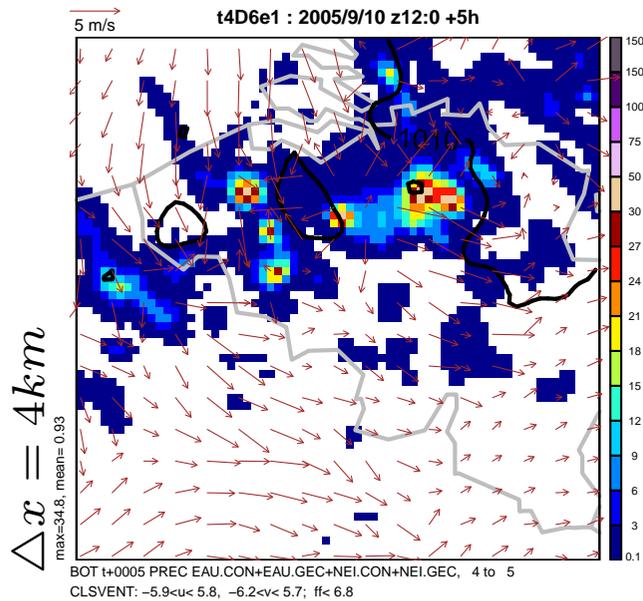
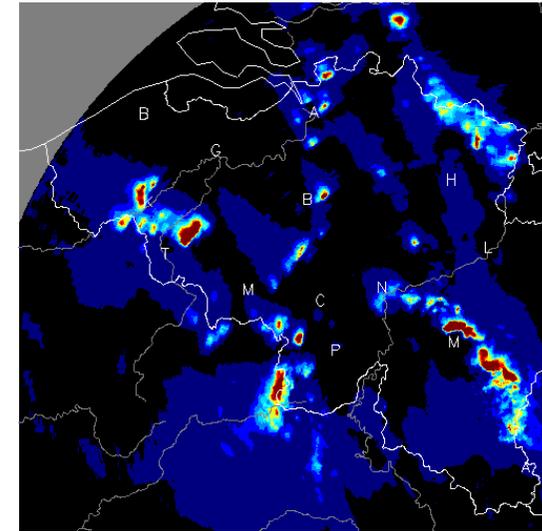
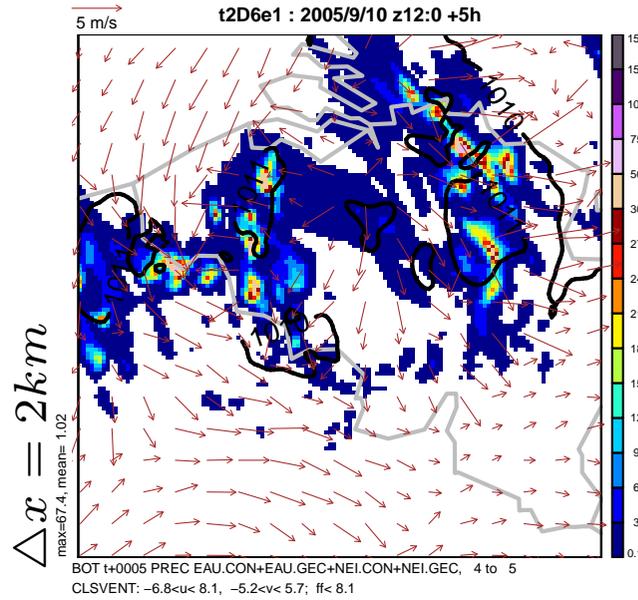
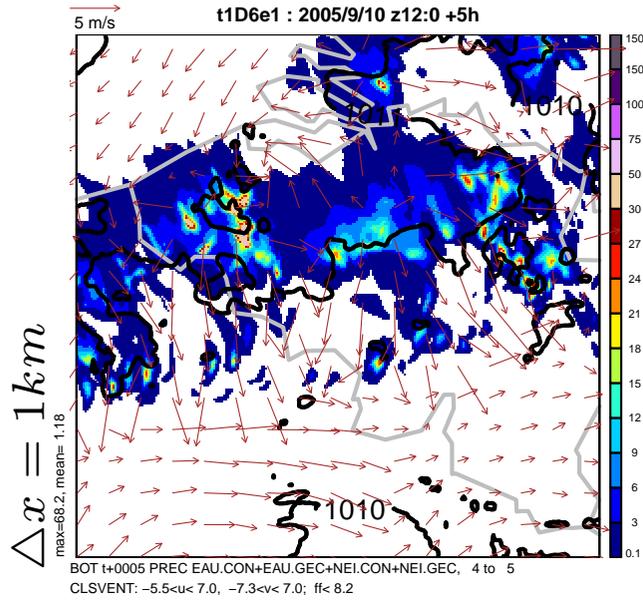
3MT: 1h-accumulated precipitation



3MT: 1h-accumulated precipitation

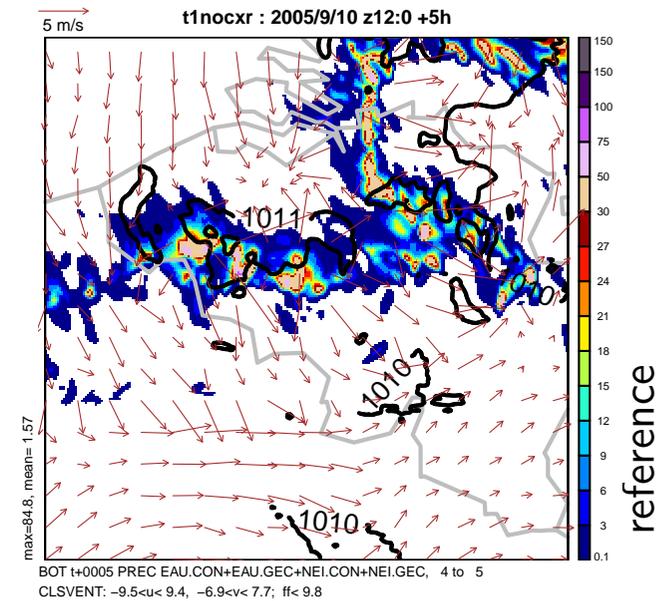
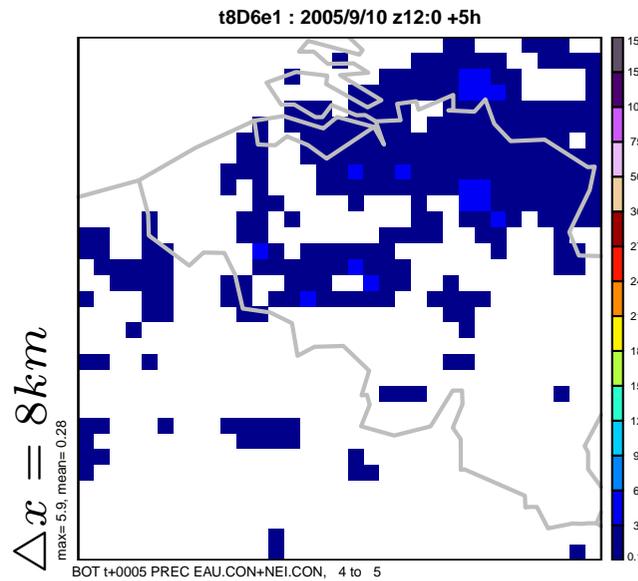
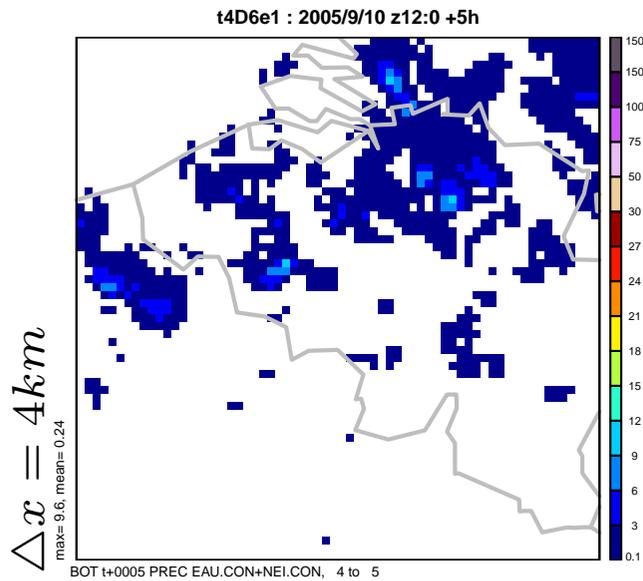
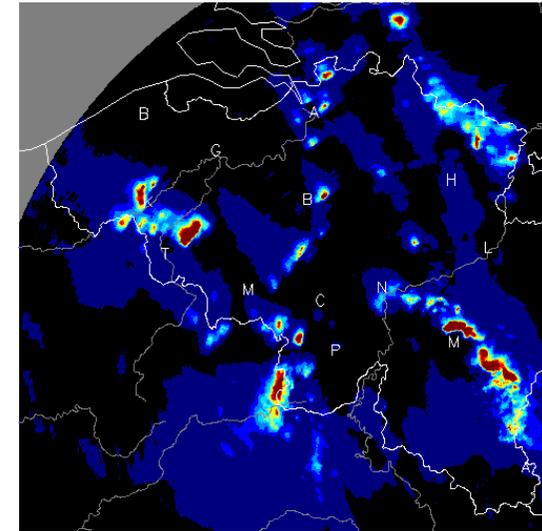
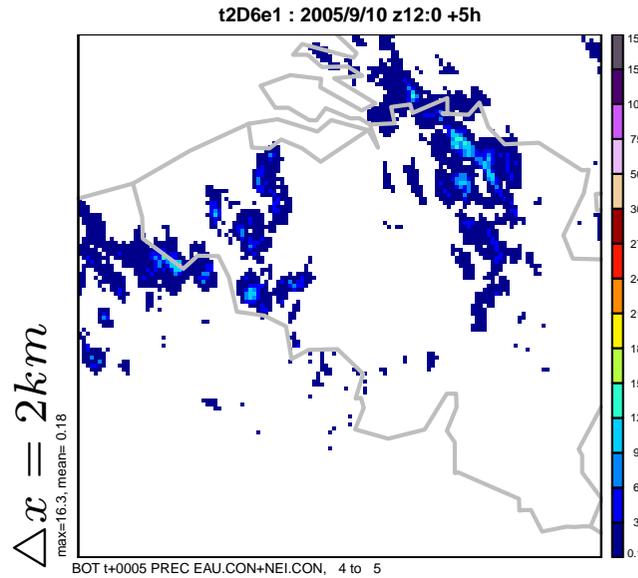
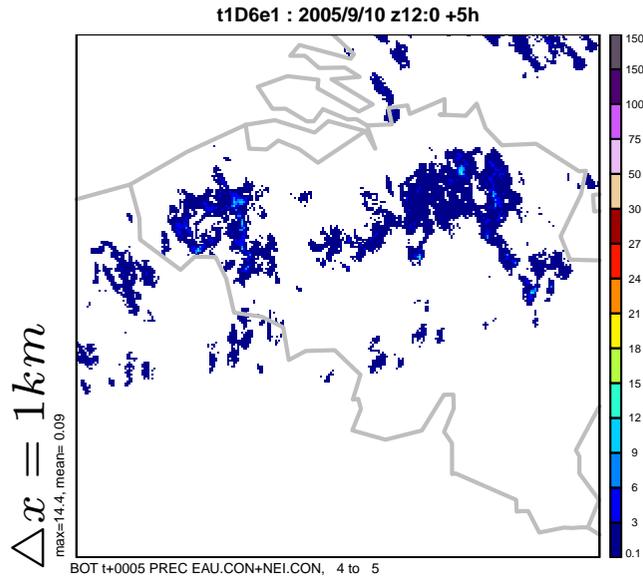


CSU: 1h-accumulated precipitation



total

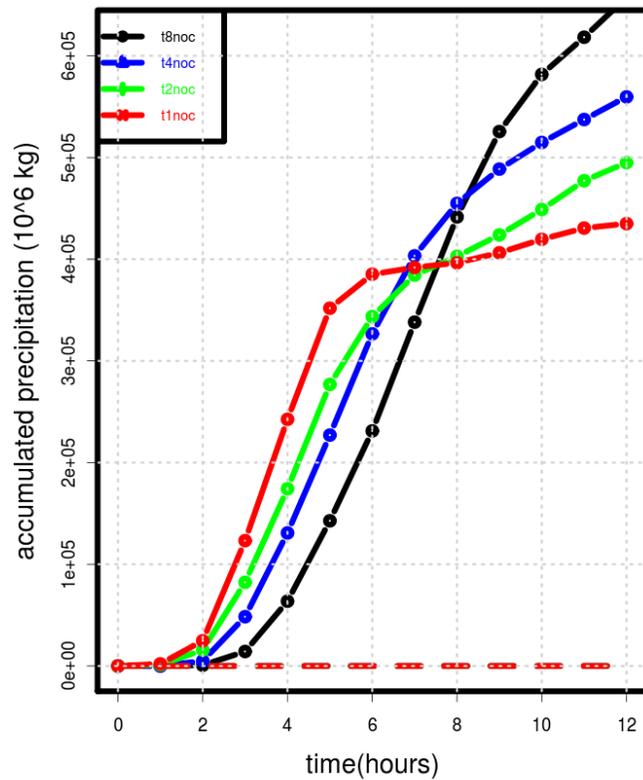
CSU: 1h-accumulated precipitation



subgrid

Domain-total accumulated precipitation evolution

accumulated water [$10^6 kg$] over domain 264x264 km

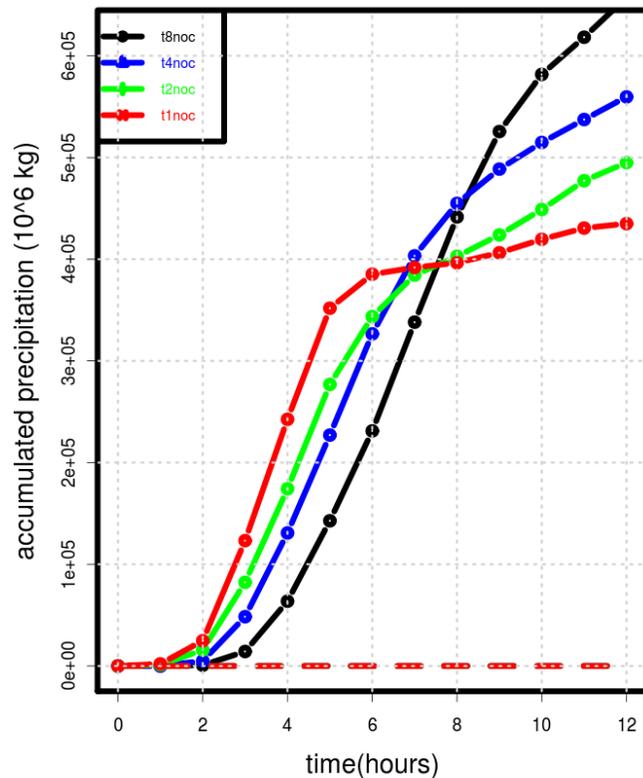


no conv param.

1km, 2km, 4km, 8km.

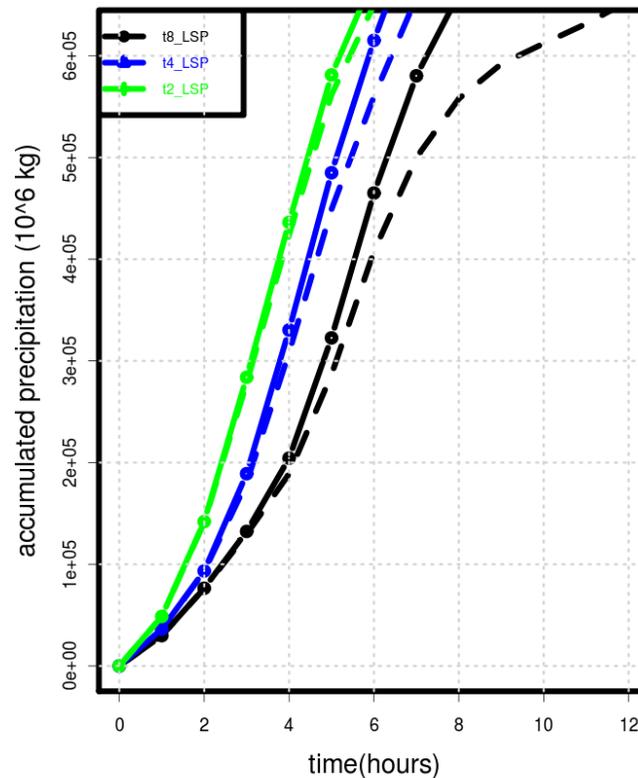
Domain-total accumulated precipitation evolution

accumulated water [$10^6 kg$] over domain 264x264 km



no conv param.

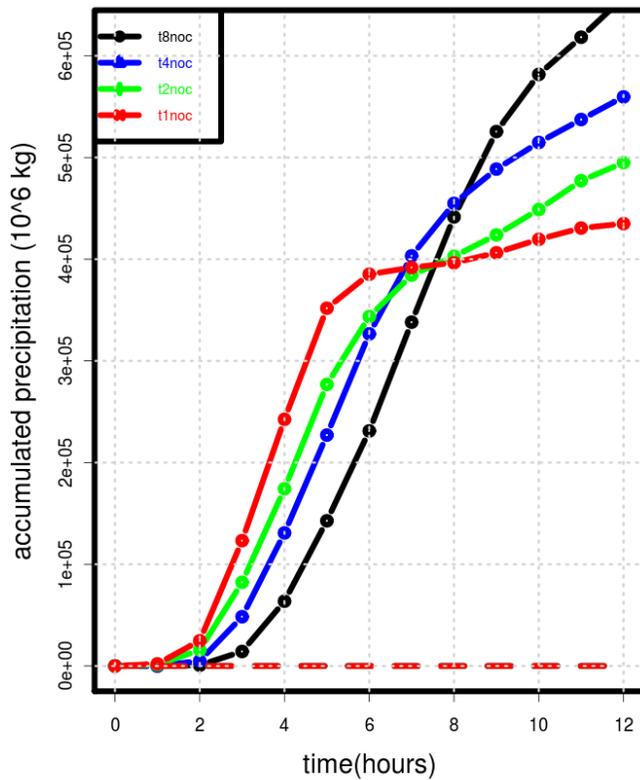
1km, 2km, 4km, 8km.



Diagn. param

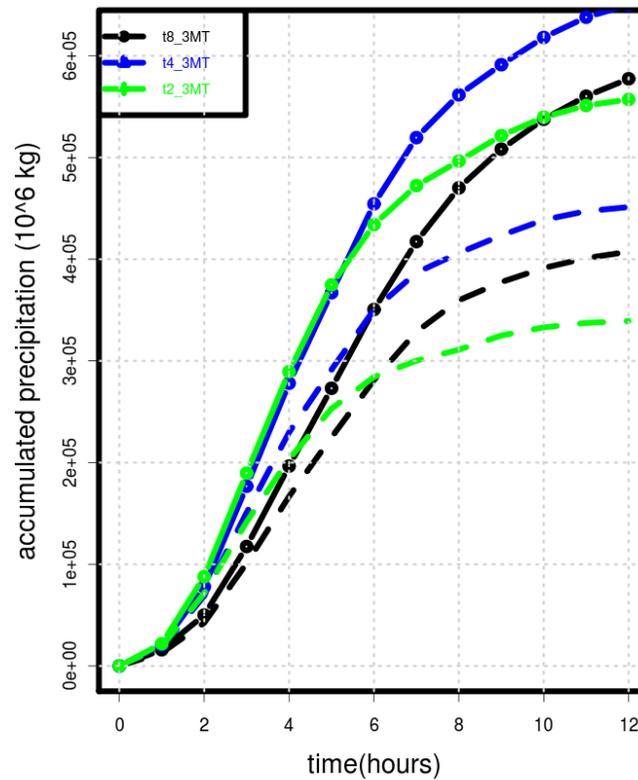
Domain-total accumulated precipitation evolution

accumulated water [$10^6 kg$] over domain 264x264 km



no conv param.

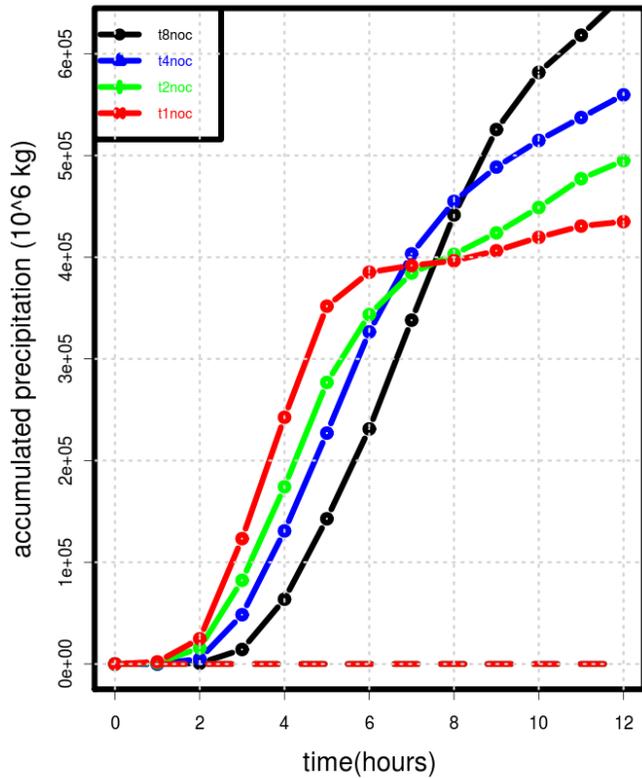
1km, 2km, 4km, 8km.



3MT

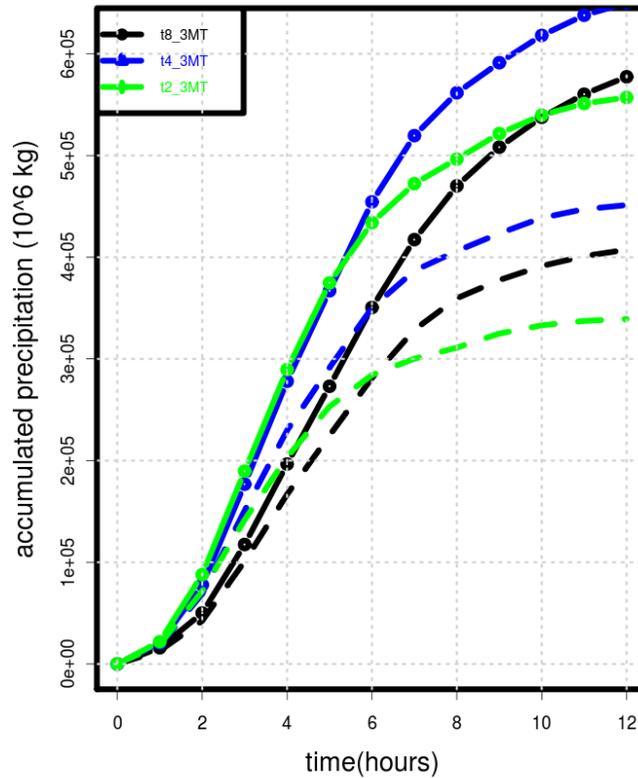
Domain-total accumulated precipitation evolution

accumulated water [$10^6 kg$] over domain 264x264 km

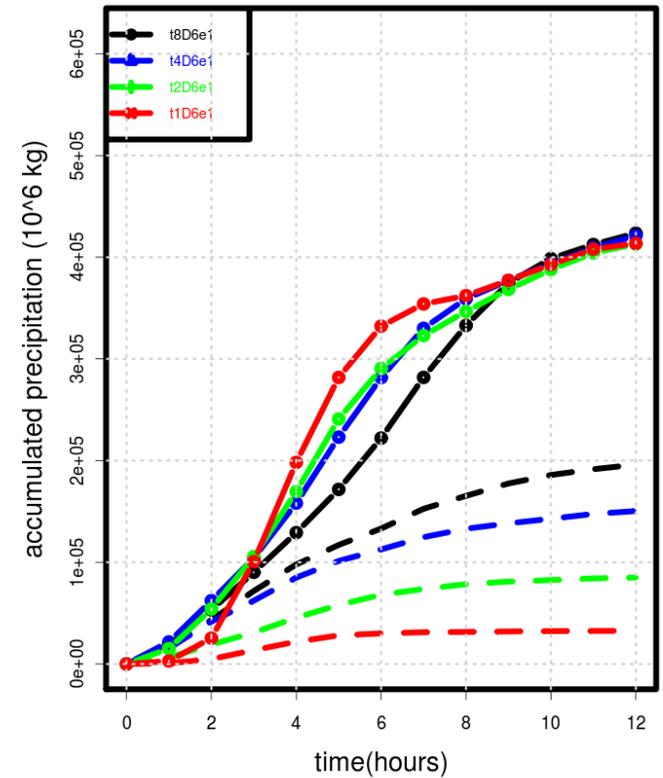


no conv param.

1km, 2km, 4km, 8km.



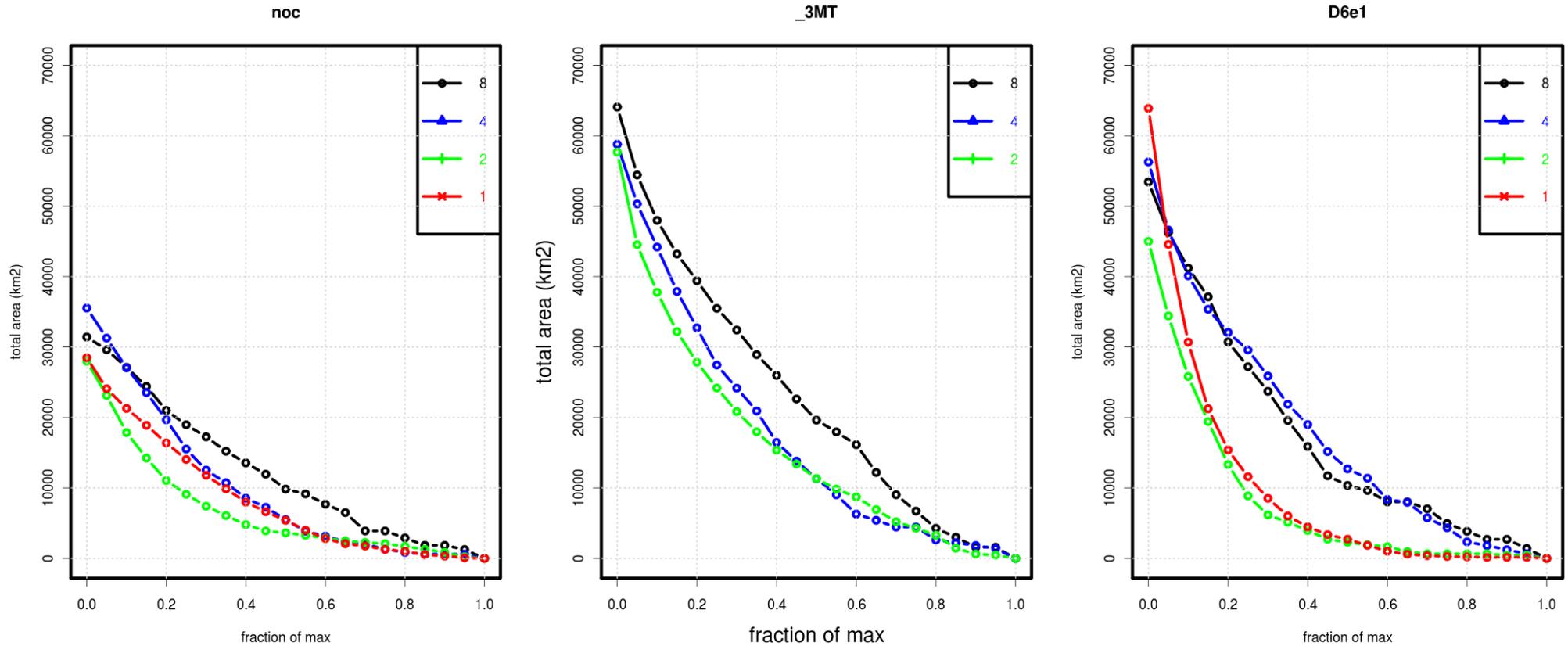
3MT



CSU

Precipitation area distribution

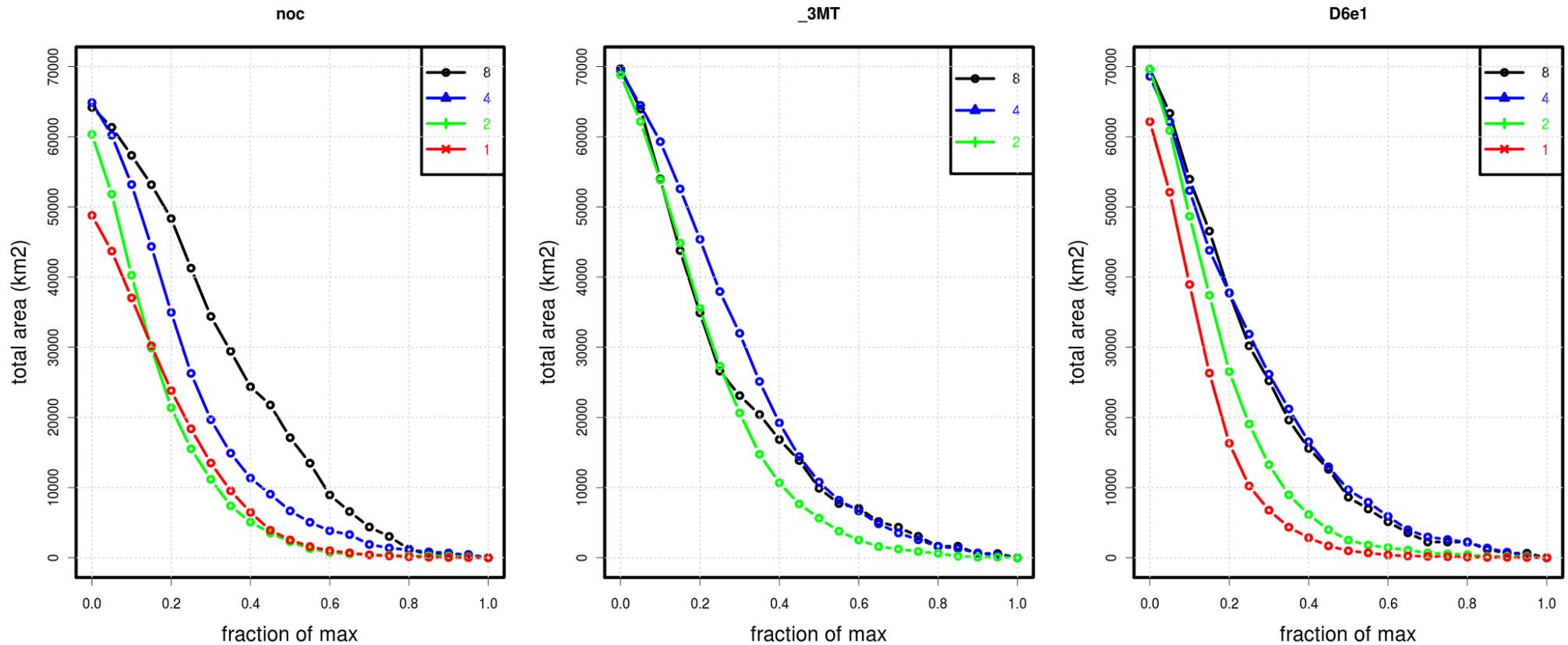
surface covered by precipitation $>$ fraction of maximum



accumulation from +5h to +6h

Precipitation area distribution

surface covered by precipitation $>$ fraction of maximum



accumulation from +0h to +12h

Final highlights

- The CSU approach allows to maintain the total accumulation at all resolutions: benefits for hydrology.
- Consistency of precipitation is a necessary condition to ensure consistency of evolution.
- Finer scale features and evolution are improved at higher resolutions, even down to 1km.
- Multiscale behaviour important for coupling or variable resolution.