GRAVITY WAVES IN THE STABLE PLANETARY BOUNDARY LAYER

Carmen J. Nappo

CJN Research Meteorology Knoxville, Tennessee USA

Workshop on Diurnal cycles and the stable boundary layer 7-10 November 2011, Reading, UK

ORIGINS OF GRAVITY WAVES OBSERVED IN THE STABLE PBL

GLOBAL:

WAVES OUTSIDE THE MODEL DOMAIN

LOCAL:

WAVES THAT CAN BE EITHER RESOLVED BY THE MODEL PHYSICS OR PARAMETERIZED AS FUNCTIONS OF THE MODEL PHYSICS.

EXAMPLES

GLOBAL:

- 1. SOLITARY WAVES
- 2. TROPOSPHERIC UNDULAR BORES
- 3. JET STREAKS
- 4. INERTIA-GRAVITY WAVES
- 5. LARGE-SCALE LEE WAVES

LOCAL:

- 1. DUCTED MODES
- 2. MESOSCALE UNDULAR BORES
- 3. LEE WAVES
- 4. CANOPY WAVES
- 5. VISCOUS WAVES



WE DESCRIBE THREE WAVE MECHANISMS WHICH CAN BE CONSIDERED LOCAL:

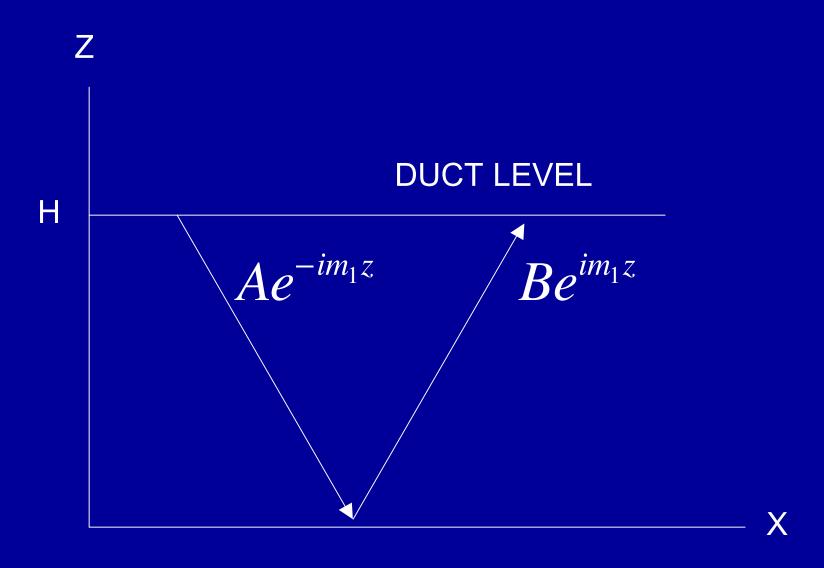
- 1. LEE WAVE DECAY (SMITH *ET AL* 2002, *JAS*)
- 2. DISSIPATIVE WAVES (HOOKE & JONES 1986, *JAS*)
- 3. STABLE PBL INTERNAL WAVES (CHIMONAS, 2002, *BLM*)

LEE WAVE DECAY

OBSERVATIONS AND ANALYSES OF WAVES OVER MONT BLANC BY SMITH et al (2002)¹ SHOW THAT THE AMPLITUDES OF TRAPPED LEE WAVES DECAY DOWNSTREAM DUE TO THE ABSORPTION OF WAVE ENERGY IN THE PBL.

¹Smith etal 2002: *J. of Atmos. Sci. 2073-2092*

THEORY



FOR PERFECT REFLECTION AT THE GROUND SURFACE

$$A+B=0$$

FOR WAVE ABSORPTION

$$A + qB = 0$$
$$0 < q < 1$$

q = REFLECTION FACTOR

RESULTS FROM COAMPS (COUPLED OCEAN/ATMOSPHERIC MESOSCALE PREDICTION SYSTEM)

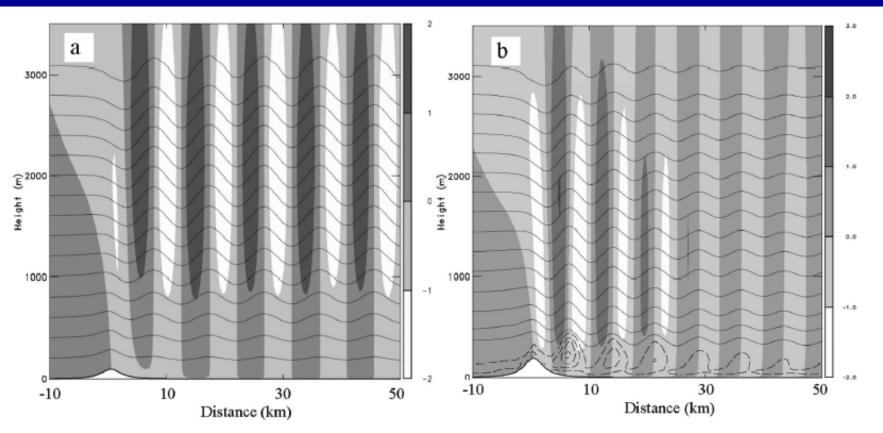


FIG. 2. Cross section of vertical velocity component (in grayscale), isentropes (interval: 4 K), and eddy viscosity coefficient (dashed contours, interval: 5 m² s⁻²) derived from simulations with (a) $h_m = 100$ m and a free-slip condition and (b) $h_m = 200$ m and a no-slip condition with $z_o = 1$ m.

FROM Jiang et al (2006): *J. Atmos. Sci.* 617-633

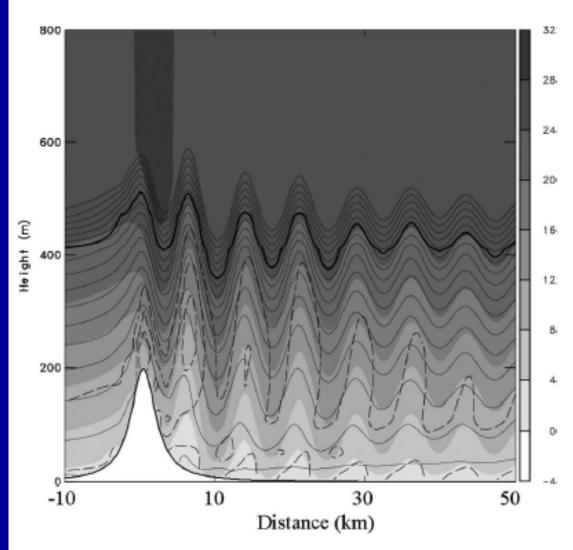
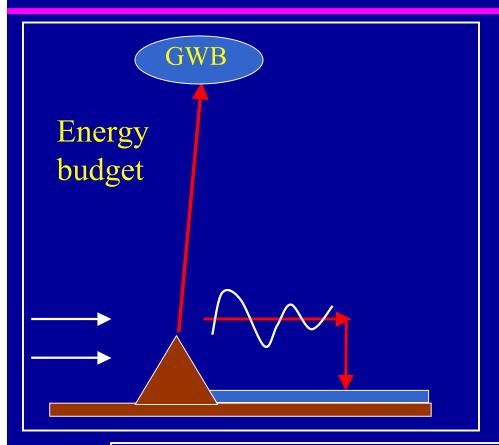
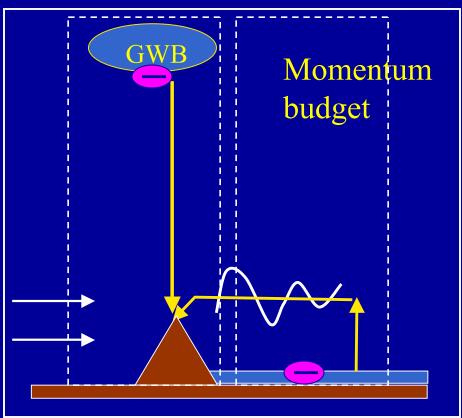


FIG. 3. Cross section of the horizontal wind component (gray-scale), isentropes (solid contours, interval: 1 K), and TKE (dashed contours, interval: 1 $\text{m}^2 \text{ s}^{-2}$) from the same no-slip simulations as in Fig. 2b. The bold contour corresponds to Richardson number Ri = 0.5.

MECHANISM (JIANG PC)





Energy flux

Momentum flux

Deceleration

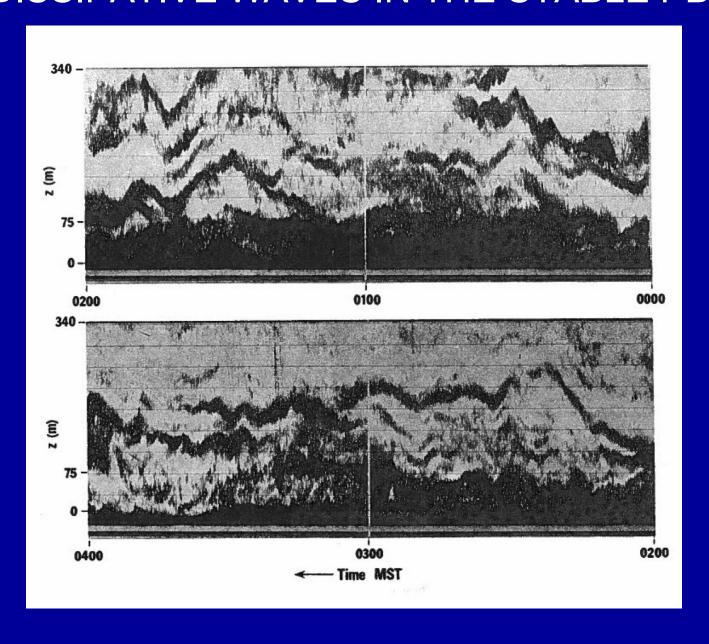
Acceleration



Wavebreaking aloft



DISSIPATIVE WAVES IN THE STABLE PBL



DOWNWARD WAVE FRONTS UPWARD WAVE FRONTS

VISCOSITY

$$\rho_0 \frac{\partial u'}{\partial t} = -\frac{\partial p'}{\partial x} + \mu \frac{\partial^2 u'}{\partial x^2} + \mu \frac{\partial^2 u'}{\partial z^2}$$

$$\rho_0 \frac{\partial w'}{\partial t} = -\frac{\partial p'}{\partial z} - \rho' g + \mu \frac{\partial^2 w'}{\partial x^2} + \mu \frac{\partial^2 w'}{\partial z^2}$$

$$\frac{\partial u'}{\partial x} + \frac{\partial w'}{\partial z} = 0$$

$$\frac{\partial \rho'}{\partial t} + w' \frac{\partial \rho_0}{\partial z} = \frac{\kappa}{\rho_0} \left(\frac{\partial^2 \rho'}{\partial x^2} + \frac{\partial^2 \rho'}{\partial x^2} \right)$$

THERMAL CONDUCTIVITY

WAVE EQUATIONS

INVISCID CASE:

$$w = a_i \exp(\omega t - kx - nz) + a_r \exp(\omega t - kx + nz)$$
$$w(0) = 0 \Rightarrow a_i + a_r = 0$$

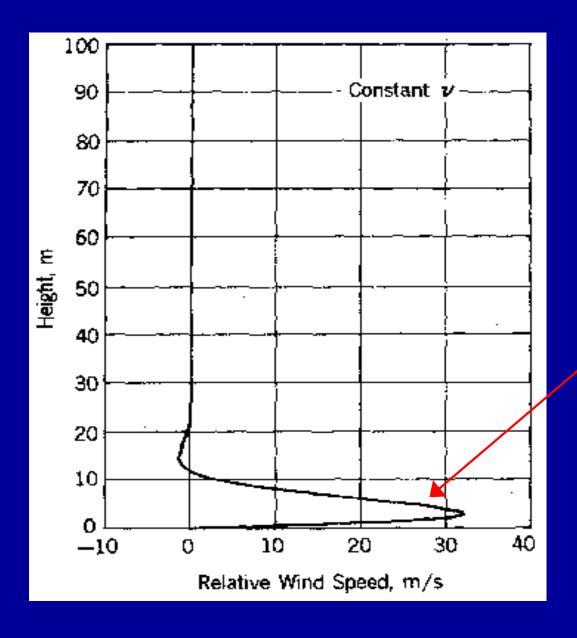
DISSIPATIVE CASE:

$$w = a_i \exp(\omega t - kx - n_g z) + a_r \exp(\omega t - kx + n_g)$$

$$+ a_v \exp(\omega t - kx + n_v z) + a_t \exp(\omega t - kx + n_t z)$$

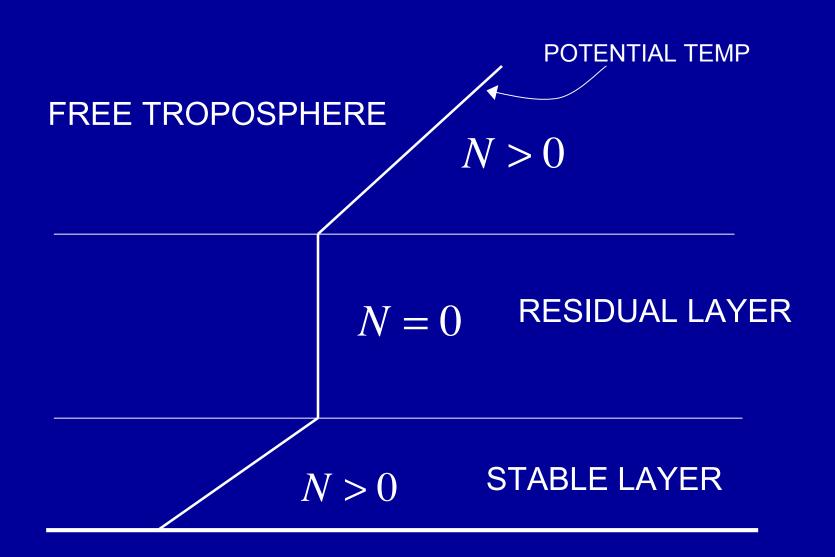
$$w(0) = 0 \Rightarrow a_i + a_r + a_v + a_t = 0$$

THE VISCOUS WAVE

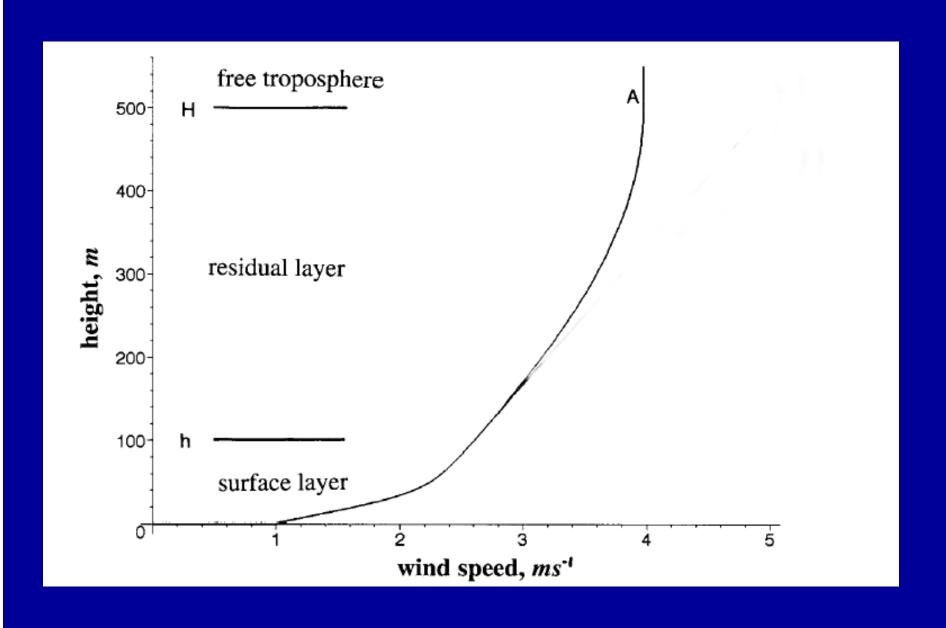


GENERATION OF SHEAR INSTABILITY?

LOCAL WAVE MODES IN THE STABLE PBL



ANALYTICAL WIND PROFILE



GOVERNING EQUATION:

$$\frac{d^{2}w}{dz^{2}} + \left[\frac{N^{2}}{(c-U)^{2}} + \frac{1}{(c-U)}\frac{d^{2}U}{dz^{2}} - \kappa^{2}\right] = 0$$

MODAL BOUNDARY CONDITIONS:

$$w(0) = 0$$
 $w(z \rightarrow \infty) = 0$

REQUIREMENTS FOR A GROWING MODE:

COMPLEX PHASE SPEED: $c = c_r + ic_i$ a Critical Level: $c_r = U(z_c)$

CRITICAL LEVEL RICHARDSON NUMBER: $0 \le Ri(z_c) \le 0.25$

IN THE RESIDUAL LAYER:

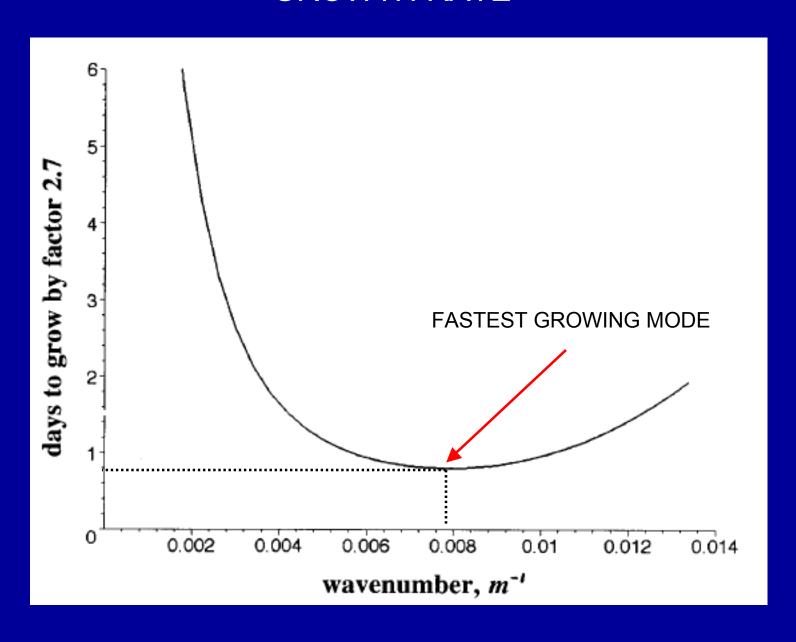
$$\frac{d^2w}{dz^2} + \left[\frac{1}{(c-U)}\frac{d^2U}{dz^2} - \kappa^2\right] = 0$$

A CRITICAL LEVEL IS VERY LIKELY TO EXIST.

THEN WAVE AMPLITUDE GROWS AS:

$$A(t) = A(0)e^{c_i \kappa t}$$

GROWTH RATE



ABOVE THE RESIDUAL LAYER:

$$\frac{d^2w}{dz^2} + \left[\frac{N_{trop}^2}{\left(c - U_{trop}\right)^2} - \kappa^2\right] w = 0$$

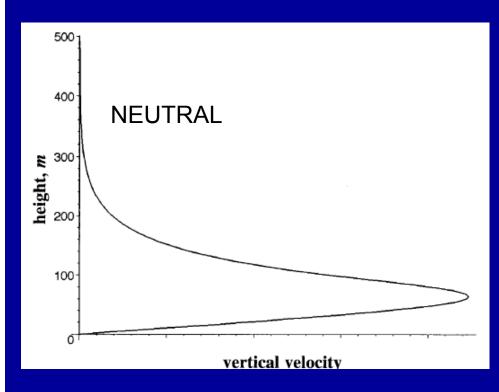
AT SOME HEIGHT THE VERTICAL WAVENUMBER BECOMES COMPLEX

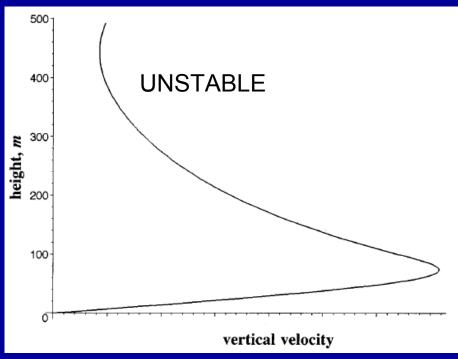
$$m = m_r + i m_i$$

NOW SOLUTION IS:

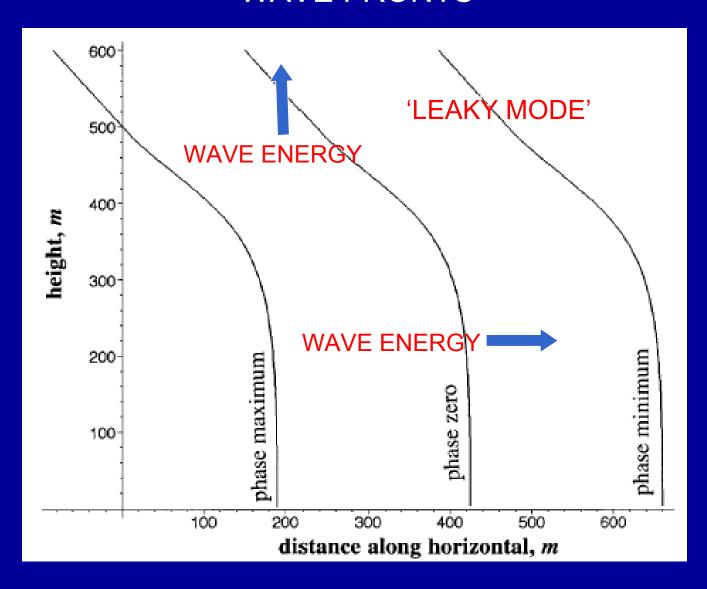
$$w(z > H) = A(0) \exp[i\kappa(c_r t - x)] \exp[c_i t - m_i z]$$
PERIODIC GROWTH DECAY

FUNDAMENTAL MODE

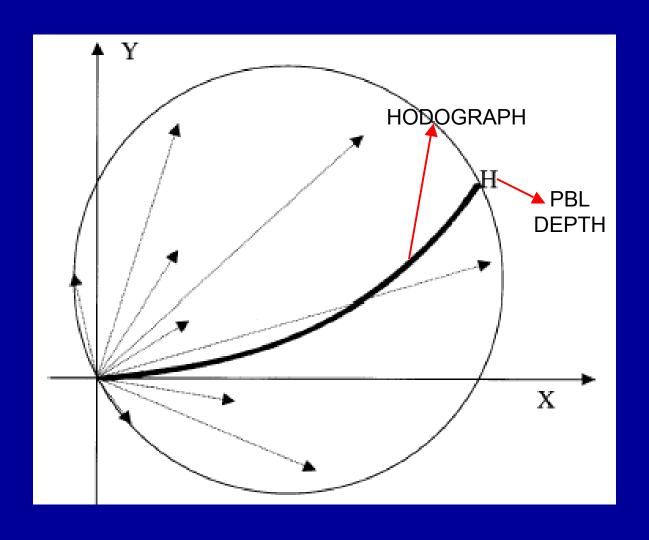




WAVE FRONTS



DOMAIN OF POSSIBLE PHASE-VELOCITIES OF BOUNDARY LAYER MODES IF THE PROJECTION OF THE WIND PROFILE INTO THE PLANE OF THE MODE HAS A CRITICAL LEVEL.



CONCLUSIONS

WE HAVE DESCRIBED THREE DIFFERENT MECHANISMS FOR GENERATING LOCAL GRAVITY WAVES IN THE STABLE PBL.

THESE MECHANISMS LEND THEMSELVES
TO PARAMETERIZATIONS WITHIN A
MESOSCALE MODEL.

AT ANY TIME ALL THREE MECHANISMS CAN BE ACTIVE.



IT'S PURE WEAPON'S GRADE BALONIUM

