

**REPRESENTING CLOUD AND PRECIPITATION
IN
NWP MODELS IN CANADA**

(Peter) M.K. Yau¹ and Jason Milbrandt²

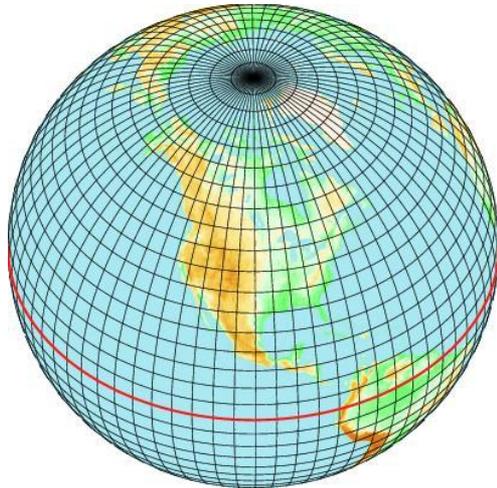
¹McGill University, Montreal, Canada

²Environment Canada [RPN], Dorval, Canada

Environment Canada's forecast model **GEM** (Global Environmental Multiscale)

Grid configurations:

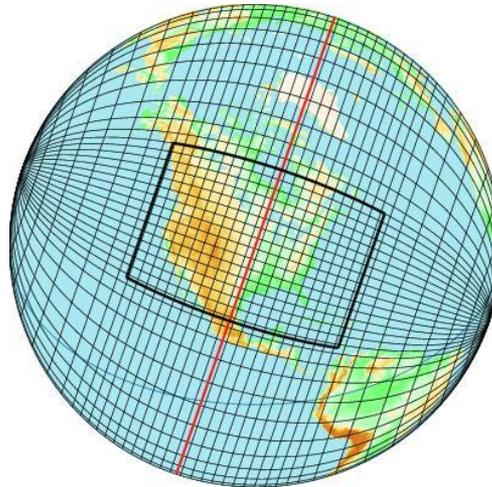
Global Uniform



- medium-range (10-d)
- $\Delta x = 35 \text{ km} \rightarrow 25 \text{ km}$
- $\Delta t = 15 \text{ min}$

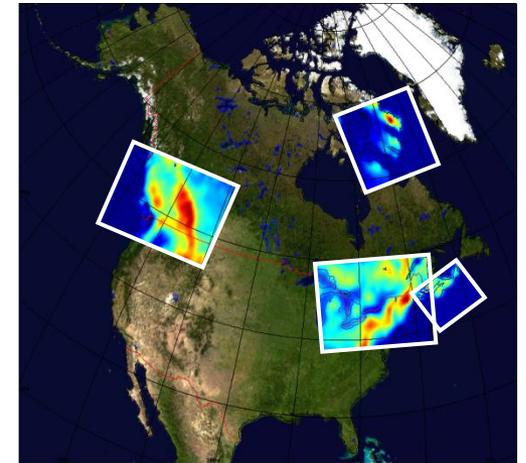
Simple Cloud Scheme

Global Variable



- short-range (48-h)
- $\Delta x = 15 \text{ km} \rightarrow 10 \text{ km}$
- $\Delta t = 7.5 \text{ min}$

Limited Area (LAM)



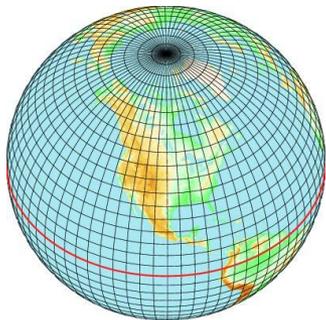
- experimental
- short-range (24-h)
- $\Delta x = 2.5 \text{ km} \rightarrow 1 \text{ km}$
- $\Delta t = 1 \text{ min}$ ($\Delta t = 30 \text{ s}$)

Detailed Microphysics Scheme

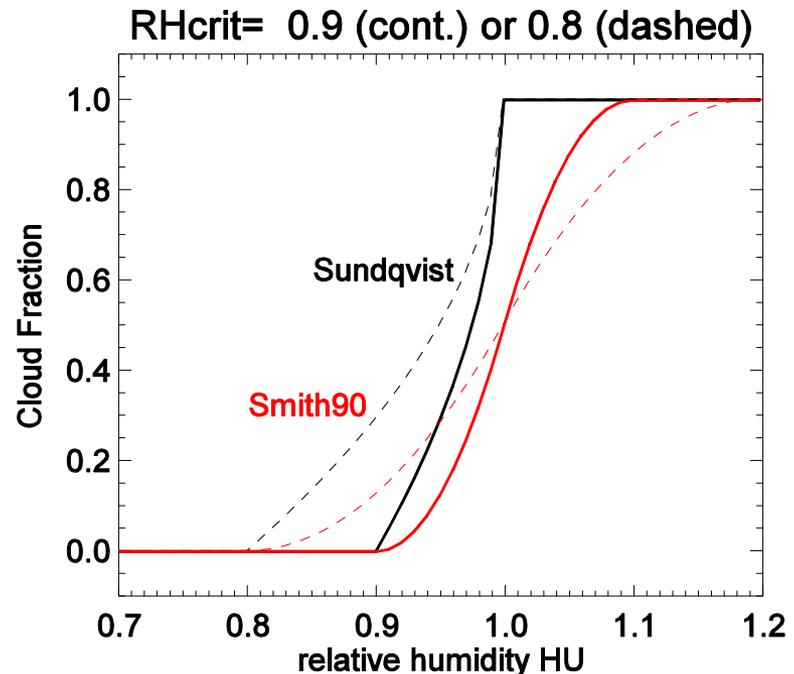
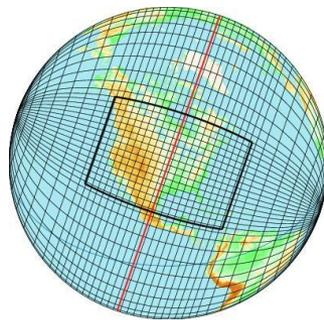
The simple cloud scheme (Sundqvist)

- Cloud-cover fraction is diagnosed (function of RH)
- Condensation occurs when RH exceeds a threshold (80% near surface)
- Total condensate (cloud water/ice) is prognostic (advected)
- Precipitation falls instantly to the ground – there is no advection of precipitation

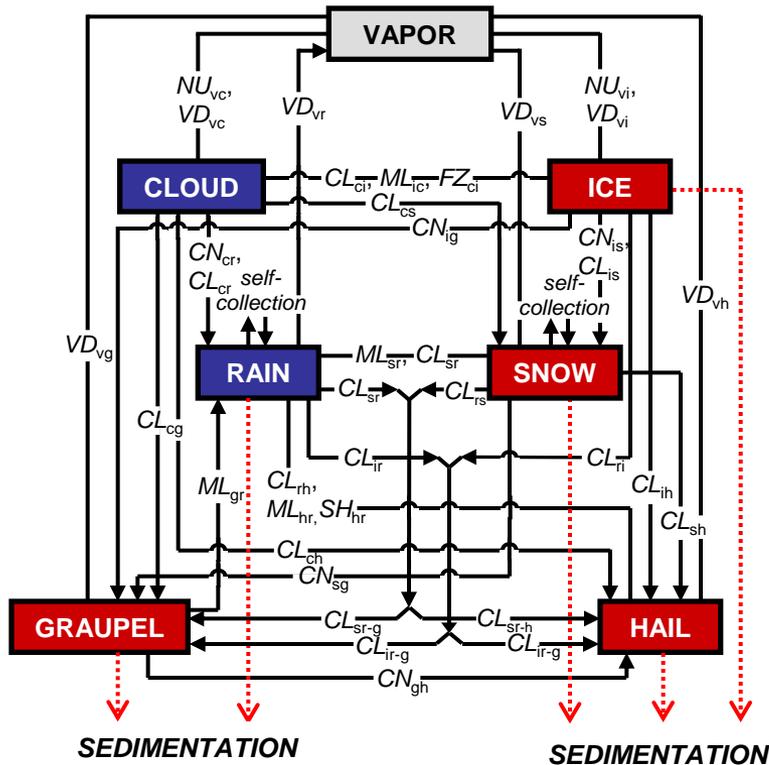
Global Uniform



Global Variable



The detailed microphysics scheme



Six hydrometeor categories:

2 liquid: *cloud, rain*

4 frozen: *ice, snow, graupel, hail*

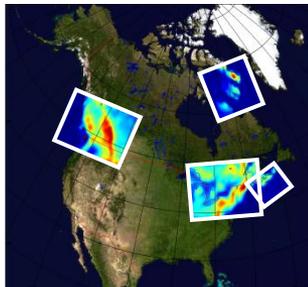
Multi-moment scheme

Milbrandt and Yau (JAS 2005 a,b)
 Milbrandt and Yau (JAS, 2006 a,b)
 Gultepe and Milbrandt
 (Pure App. Geoph., 2007)
 Milbrandt et al. (MWR, 2008)
 Milbrandt et al. (MWR, 2010)
 Dawson et al. (MWR, 2010)

Scheme implemented in

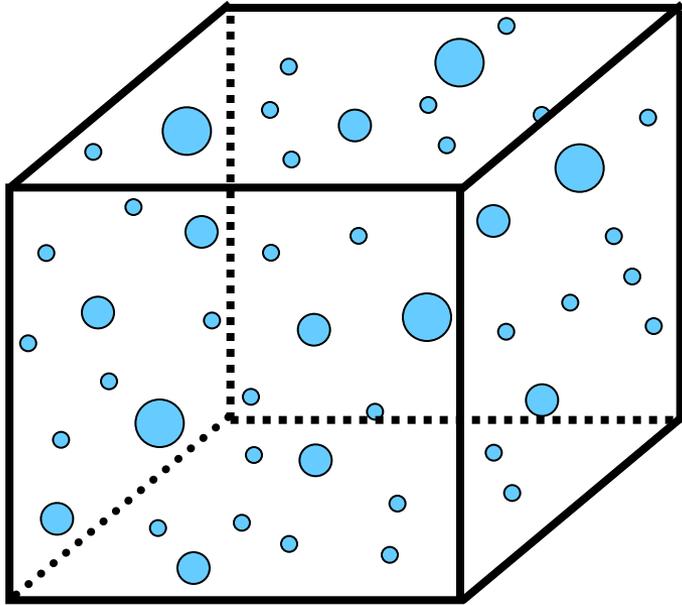
GEM-LAM, Global variable (Canada)
 ARPS (U Oklahoma, US)
 WRF 3.2 (US)

Limited Area Model



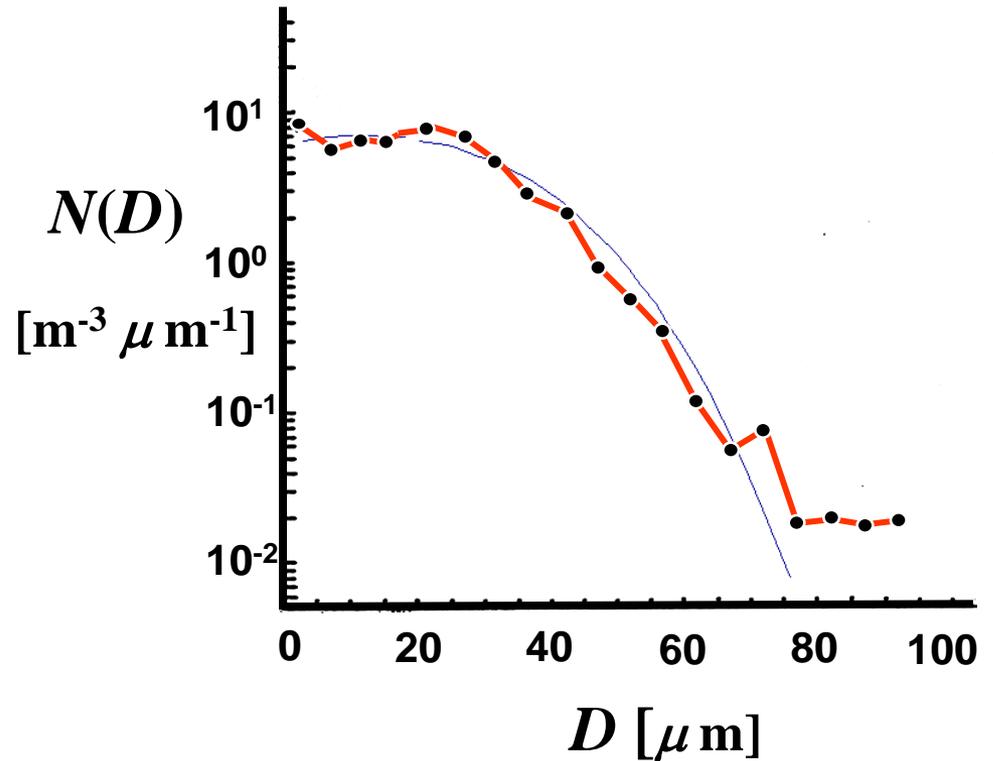
- Overview of the scheme
- Testing and improvement in IMPROVE-2
(GEM-LAM)
- Forecast in winter Olympics 2010
(GEM-LAM)
- Testing over Arctic (GEM-Global Variable)

Representing the size spectrum



1 m³

ANAYLTICAL FUNCTION



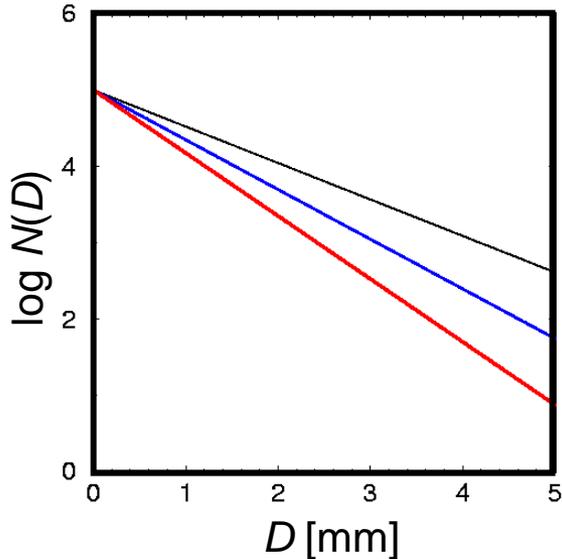
BULK METHOD

Gamma Distribution Function:

$$N(D) = N_0 D^\alpha e^{-\lambda D}$$

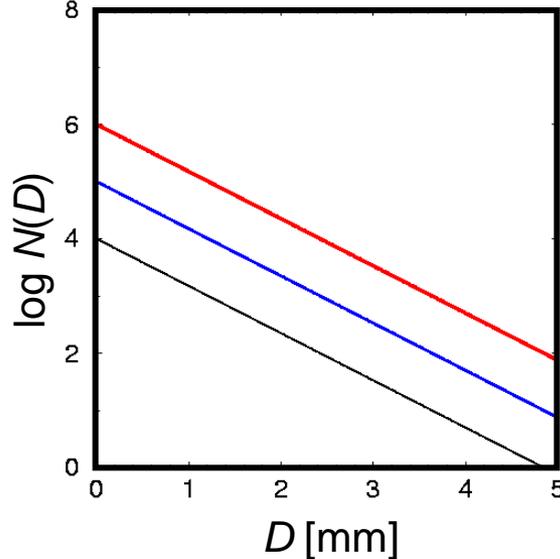
Varying λ (slope)

(N_0 and α constant)



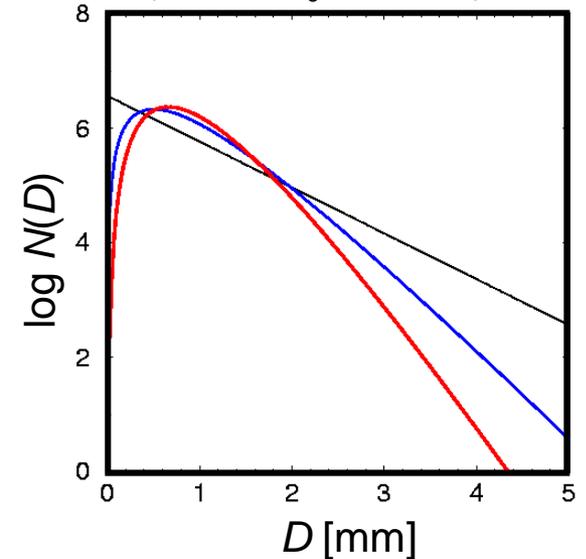
Varying N_0 (intercept)

(λ and α constant)



Varying α (shape)

(Q^* and N_0 constant)



INCREASING
VALUES

(of λ , N_0 and α)

* $Q = \rho q$ (mass content)

BULK METHOD

Predict evolution of specific moment(s)

e.g. q_x, N_{Tx}, \dots



Implies prediction of evolution of parameters

i.e. N_{0x}, λ_x, \dots

Size Distribution Function:

$$N_x(D) = N_{0x} D^{\alpha_x} e^{-\lambda_x D}$$

Total number concentration, N_{Tx}

$$N_{Tx} \equiv \int_0^{\infty} N_x(D) dD = M_x(0)$$

Mass mixing ratio, q_x

$$q_x \equiv \frac{c_x}{\rho} \int_0^{\infty} D^3 N_x(D) dD = \frac{c_x}{\rho} M_x(3),$$

where $m_x(D) = c_x D^3$, $\rho = \text{air density}$

Radar reflectivity factor, Z_x

$$Z_x \equiv \int_0^{\infty} D^6 N_x(D) dD = M_x(6)$$

p^{th} moment:

$$M_x(p) \equiv \int_0^{\infty} D^p N_x(D) dD = N_{0x} \frac{\Gamma(1 + \alpha_x + p)}{\lambda_x^{p+1+\alpha_x}}$$

BULK METHOD

Predict evolution of specific moment(s)

e.g. q_x, N_{Tx}, \dots



Implies prediction of evolution of parameters

i.e. N_{0x}, λ_x, \dots

Size Distribution Function:

$$N_x(D) = N_{0x} D^{\alpha_x} e^{-\lambda_x D}$$

For every predicted moment, there is one prognostic parameter.

The remaining parameters are prescribed or diagnosed.

e.g. **One-moment scheme:**

q_x is predicted;

→ λ_x is prognosed

(N_{0x} and α_x are specified)

Two-moment scheme:

q_x and N_{Tx} are predicted;

→ λ_x and N_{0x} are prognosed;

(α_x is specified)

Three-moment scheme:

q_x, N_{Tx} and Z_x are predicted;

→ λ_x, N_{0x} and α_x is prognosed

p^{th} moment:

$$M_x(p) \equiv \int_0^{\infty} D^p N_x(D) dD = N_{0x} \frac{\Gamma(1 + \alpha_x + p)}{\lambda_x^{p+1+\alpha_x}}$$

CLOSURE OF SYSTEM

Solve for shape parameter α from

$$\frac{c^2 N_T Z}{(\rho q)^2} = G(\alpha) = \frac{(\alpha + 6)(\alpha + 5)(\alpha + 4)}{(\alpha + 3)(\alpha + 2)(\alpha + 1)},$$

where $m(D) = cD^3$, and $\rho = \text{air density}$

Solve for slope parameter λ from

$$\lambda = \left(\frac{c N_T \Gamma(\alpha + 4)}{\rho q \Gamma(\alpha + 1)} \right)^{\frac{1}{3}},$$

Solve for intercept parameter N_0 from

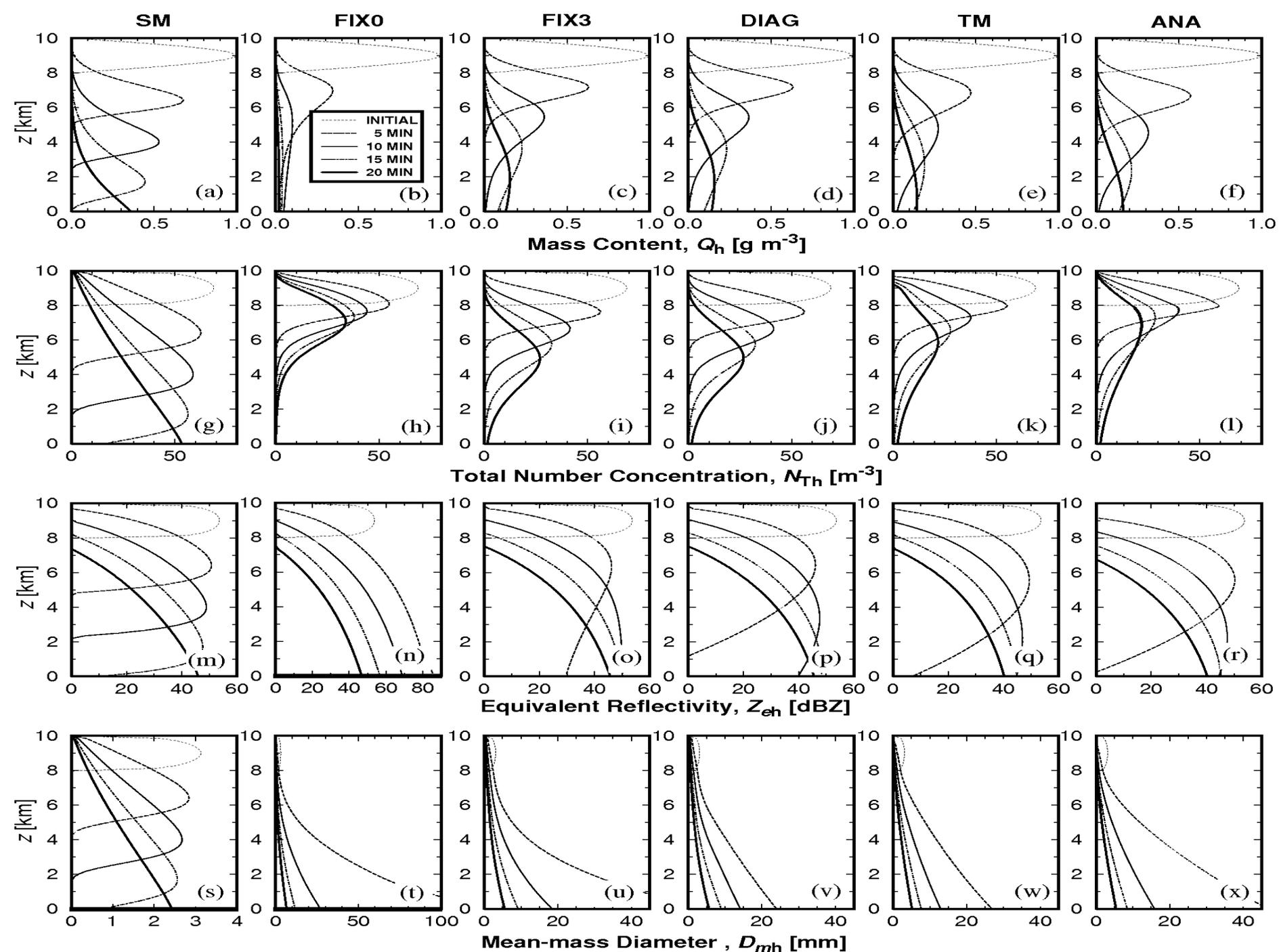
$$N_0 = \frac{N_T \lambda^{\alpha+1}}{\Gamma(\alpha + 1)}$$

→ N_T and q vary monotonically in a 1-moment scheme

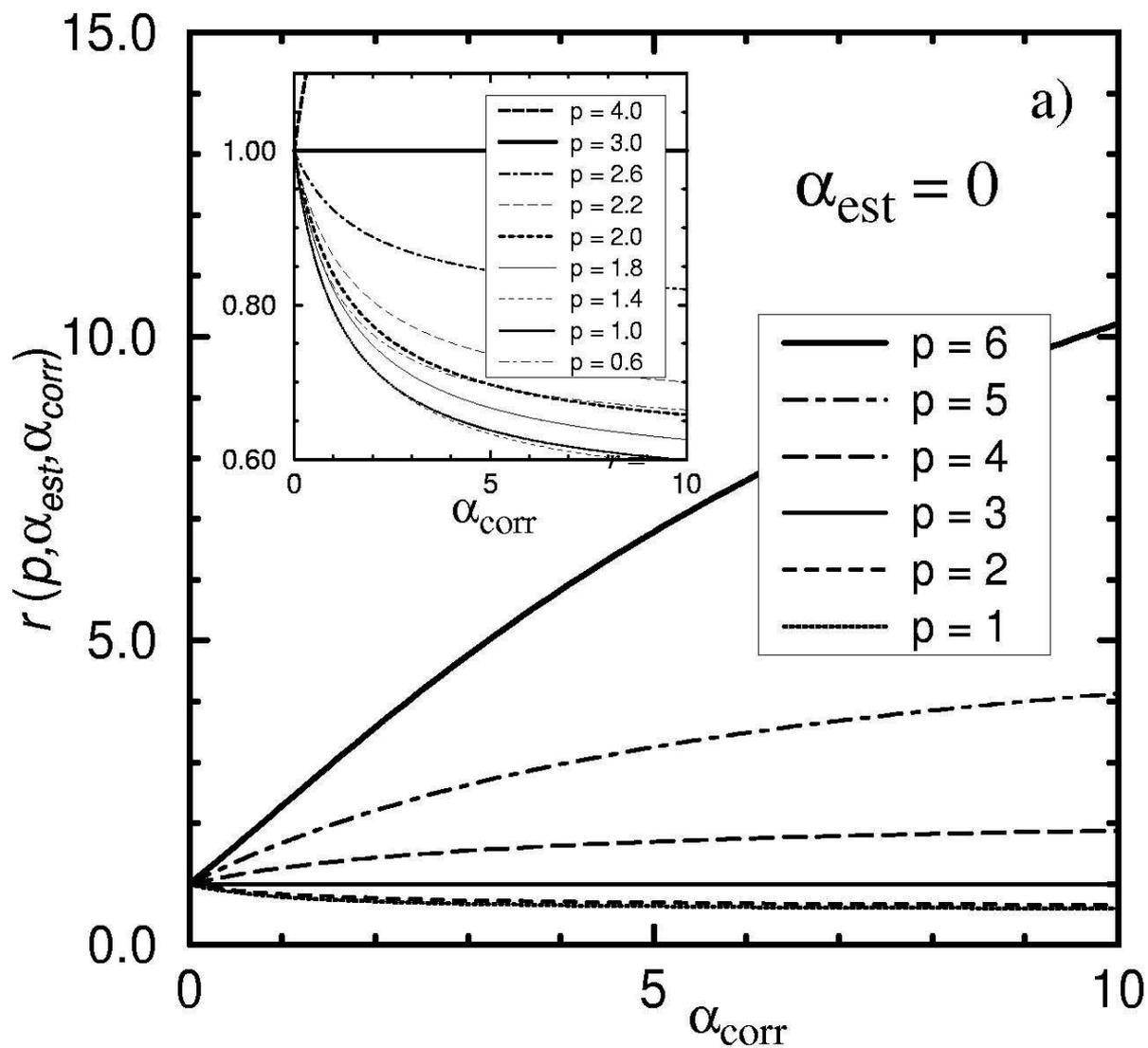
Diagnostic closure for α in 2-moment scheme

$$D_m = \left[\frac{\rho q}{cN_T} \right]^{\frac{1}{3}},$$

$$\alpha = f(D_m)$$



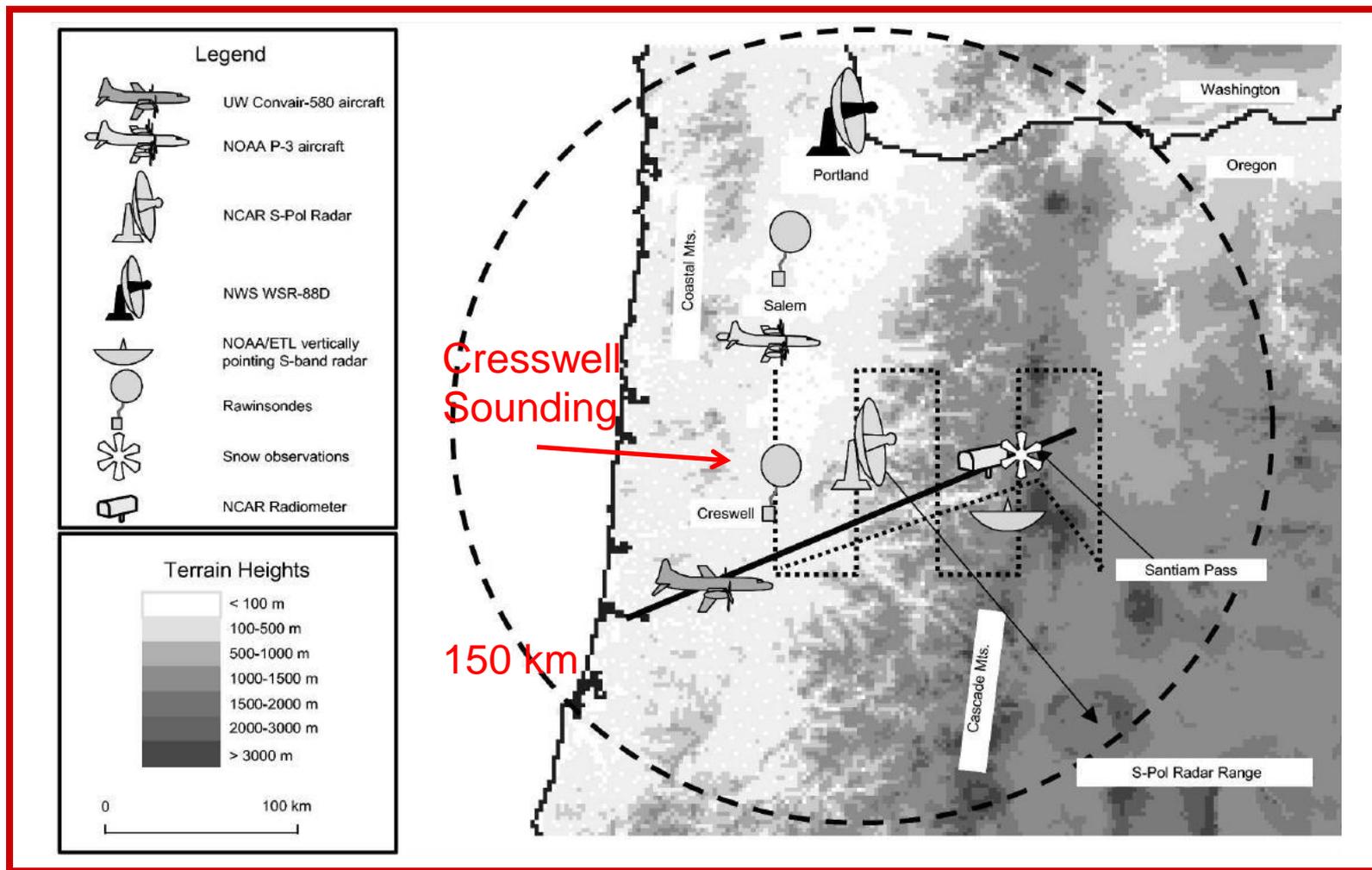
$$r = M(p, \alpha_{\text{est}}) / M(p, \alpha_{\text{corr}})$$



**Verification and
improvement of Multi-
moment scheme in GEM-
LAM (1 km) in IMPROVE-2**

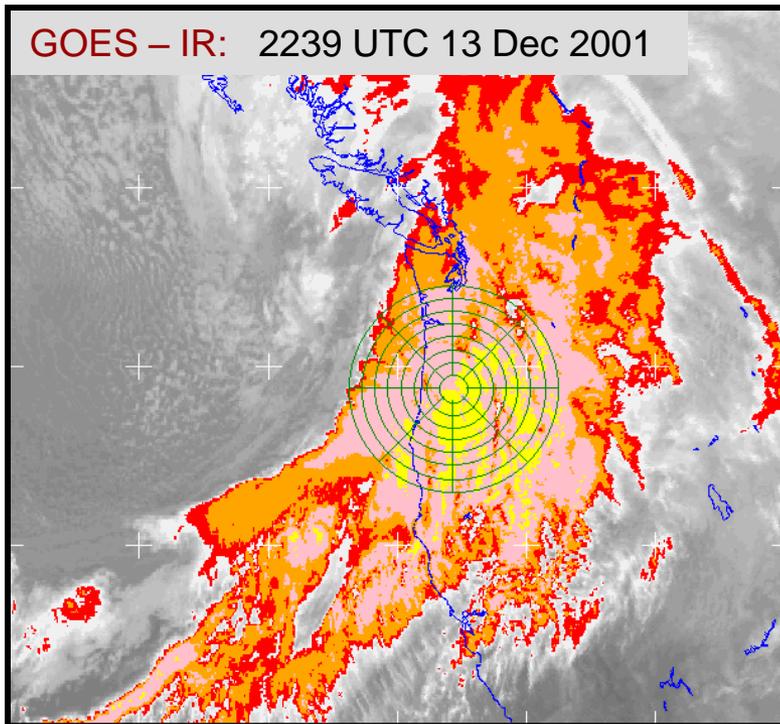
November-December 2001: IMPROVE-2 Observational Campaign

Improvement of Microphysical Parameterization through Observational Verification Experiment



13-14 Dec 2001 case:

- chosen for study at W.M.O. *International Cloud Modeling Workshop*, Hamburg (July 2004)
- special issue of *J. Atmos. Sci.* (October 2005) dedicated to IMPROVE-2



Characteristics:

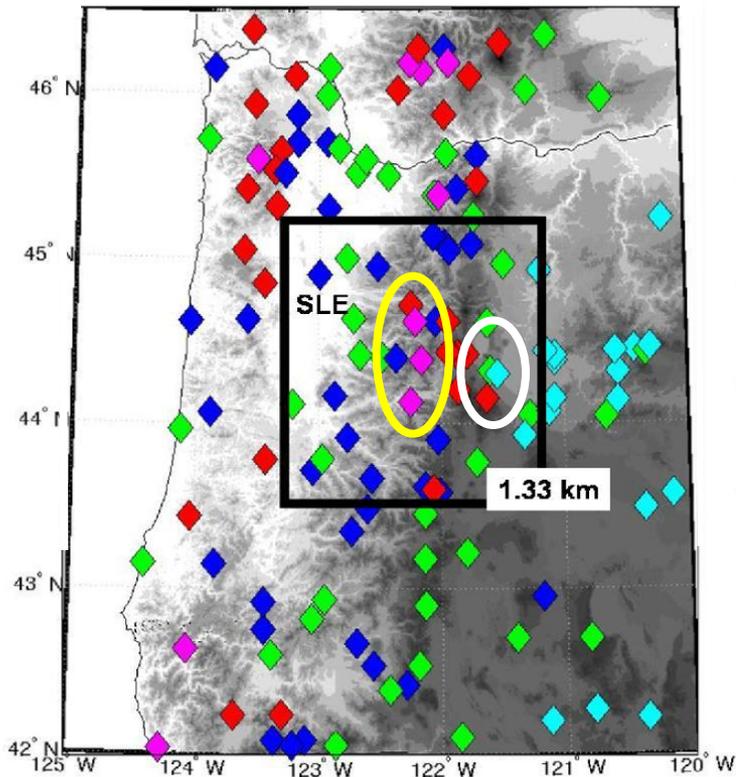
- large-scale baroclinic system
- strong low-level cross-barrier flow

Precipitation in IOP region:

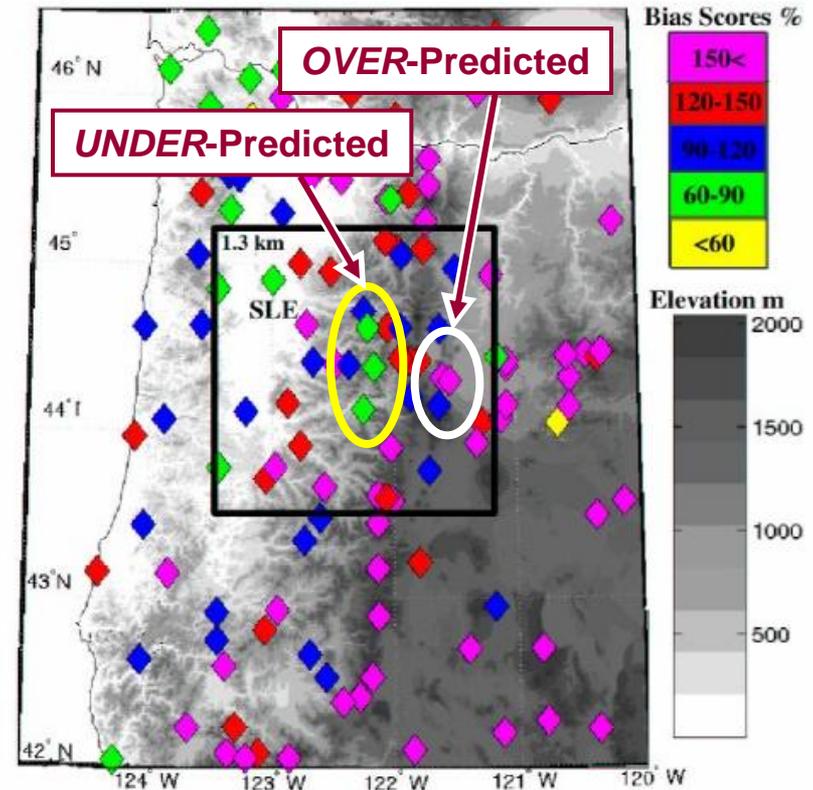
- prefrontal showers;
- moderate to heavy stratiform rain (associated with mid-level baroclinic zone);
- surface frontal rain-band;
- transition to sporadic showers

13-14 Dec 2001 case:

- MM5 runs at 4-km and 1.3 km exhibited errors in surface precipitation attributed to problems associated with the microphysics (SM Reisner-2)

**OBSERVED PRECIPITATION**

1600 UTC 13 Dec – 0800 UTC (18 h)

**BIAS SCORES**

4-km MM5 Simulation

NCAR S-Pol Radar

Portland Radar

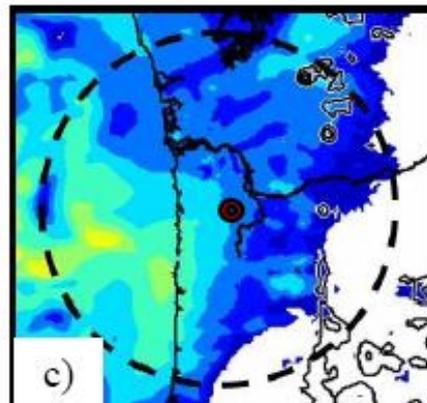
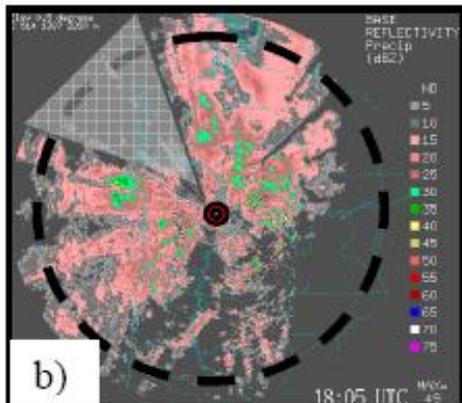
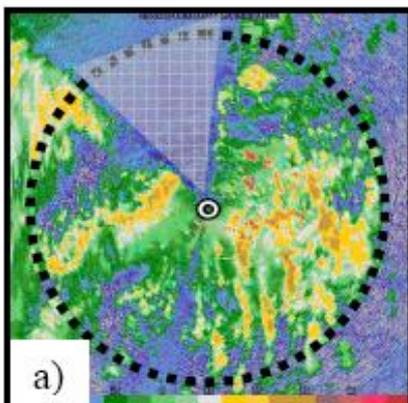
4-km Simulation

S-Pol
1.5 PPI
R 150 km

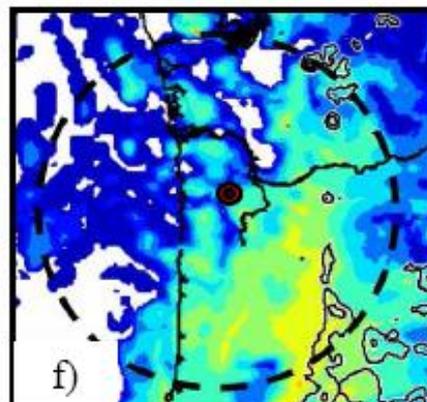
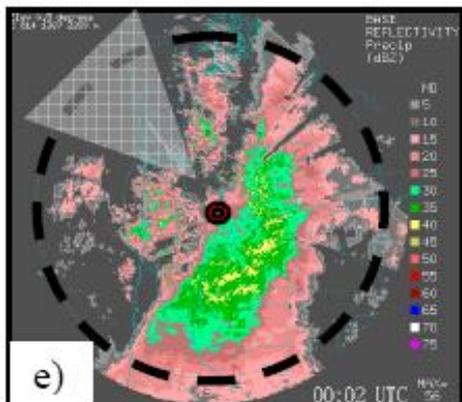
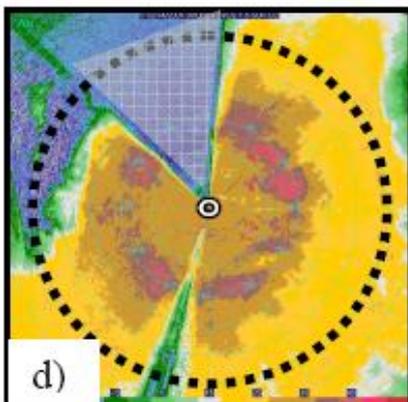
Portland
0.5 PPI
R 200 km

4km-GEM
700 hPa

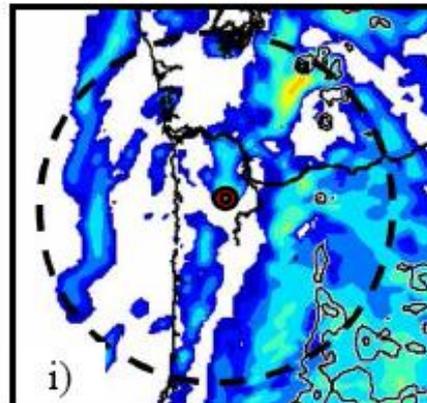
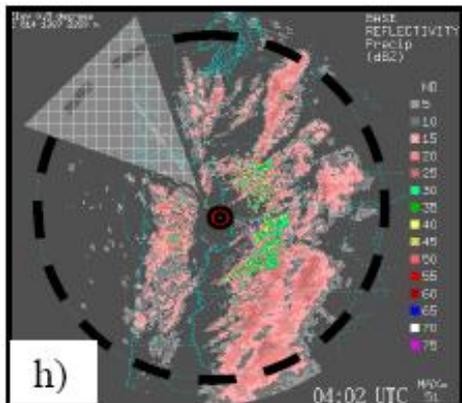
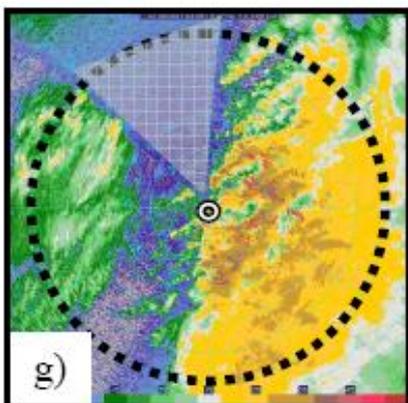
1800 UTC



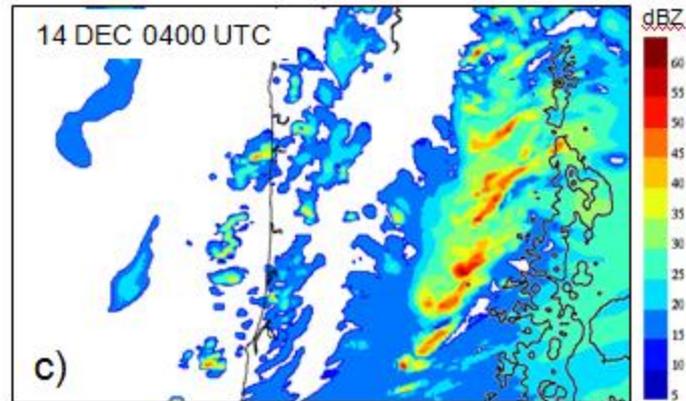
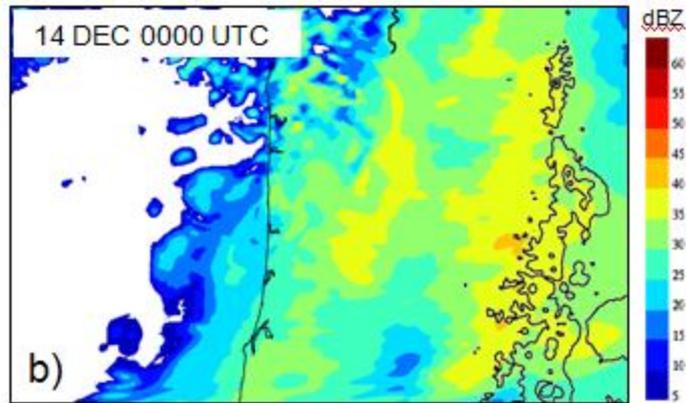
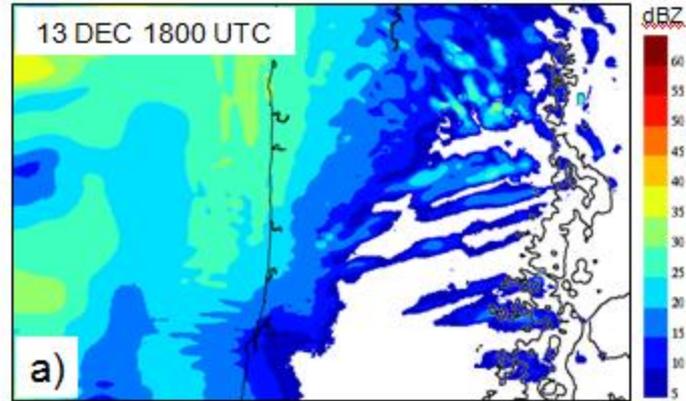
0000 UTC



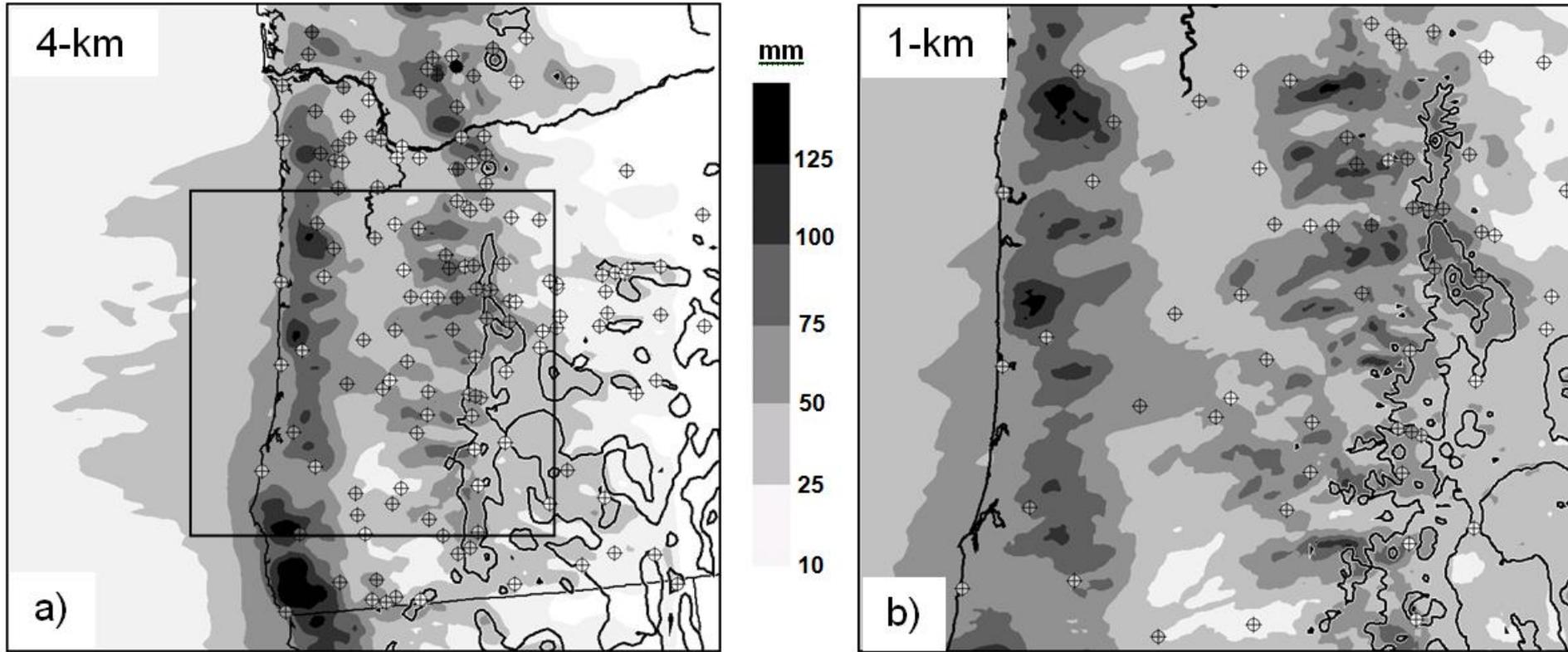
0400 UTC



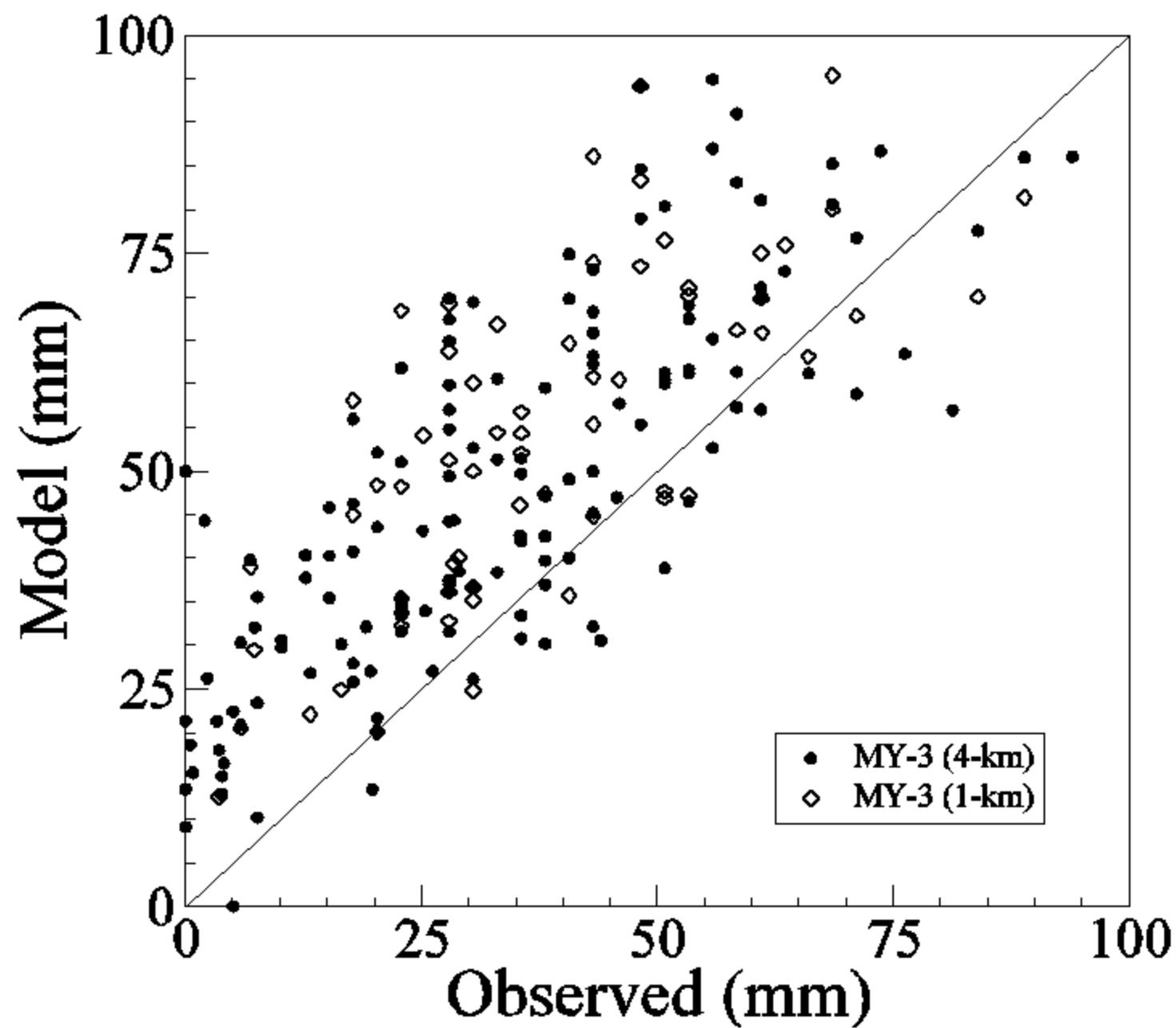
1-km GEM
E. Reflectivity



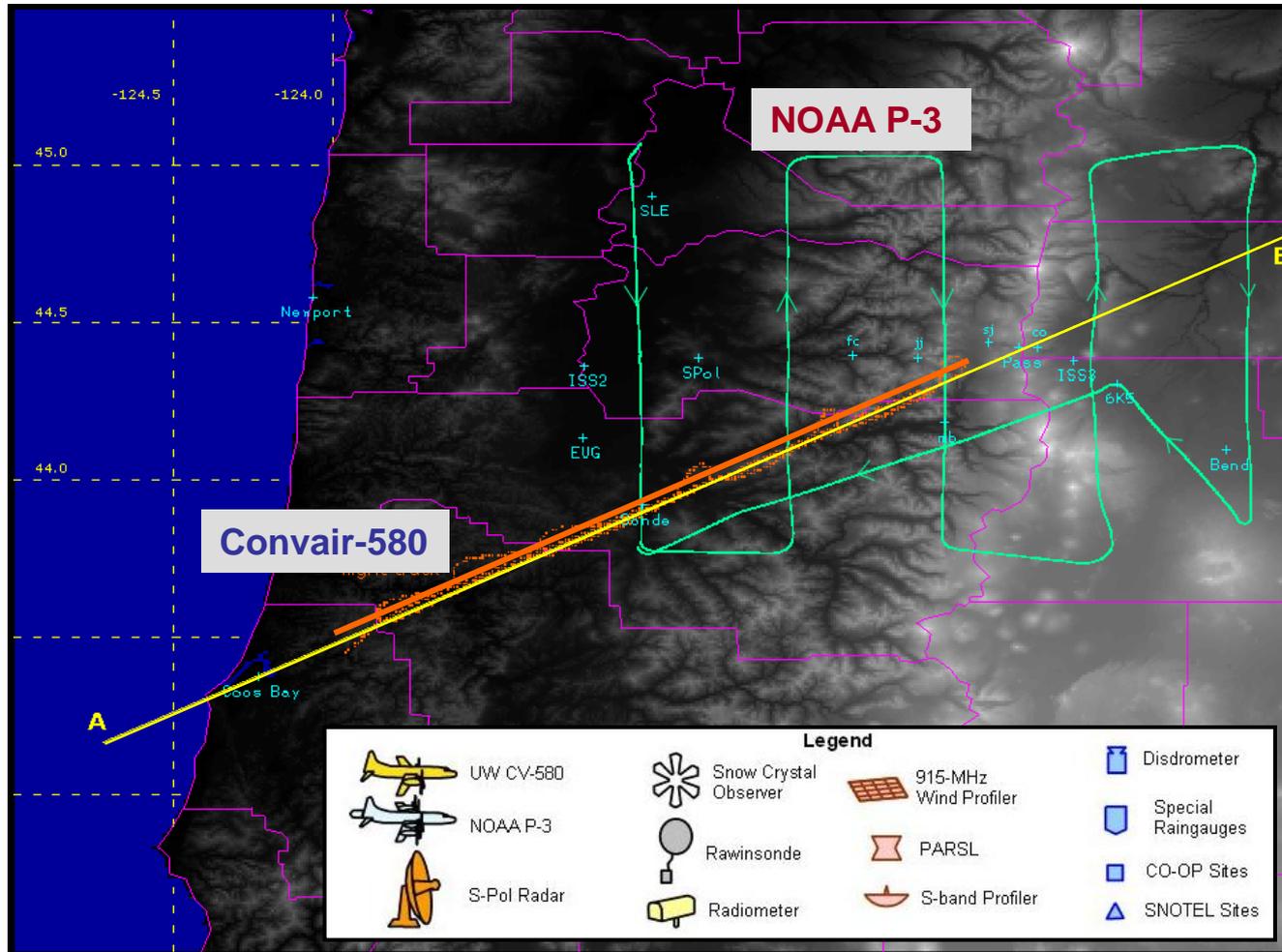
18-h Accumulated Precipitation Observed vs. Simulated



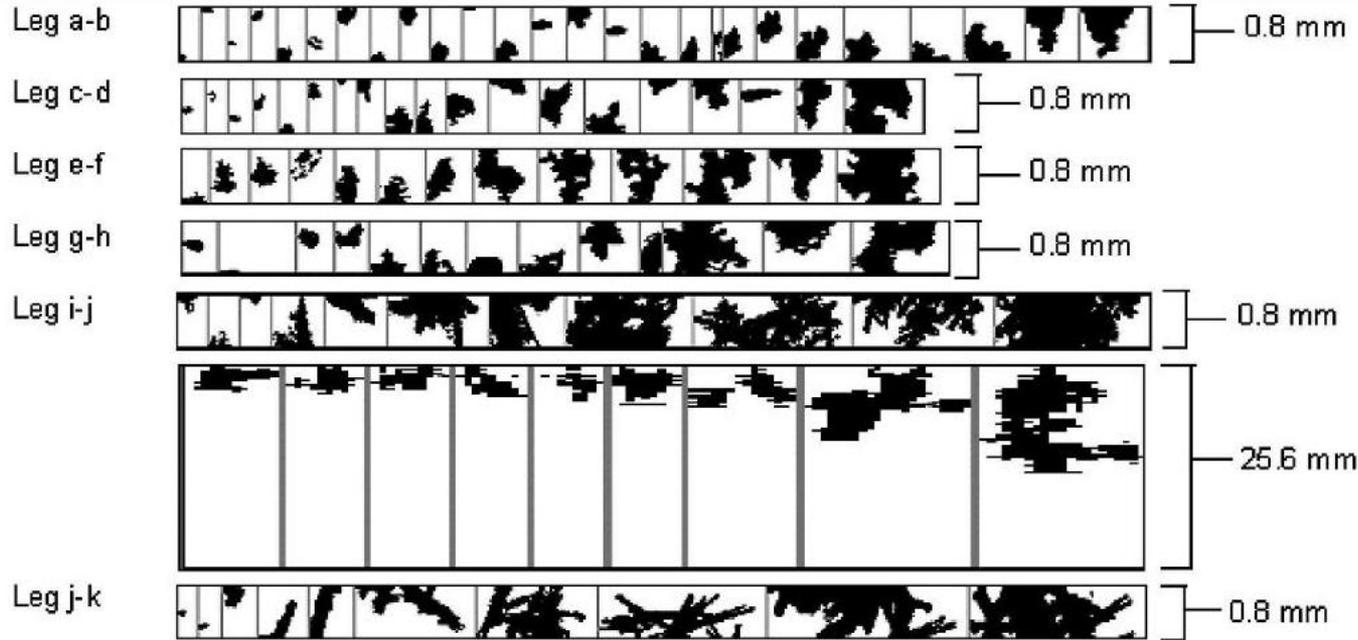
No pronounced over prediction along lee side of Cascade



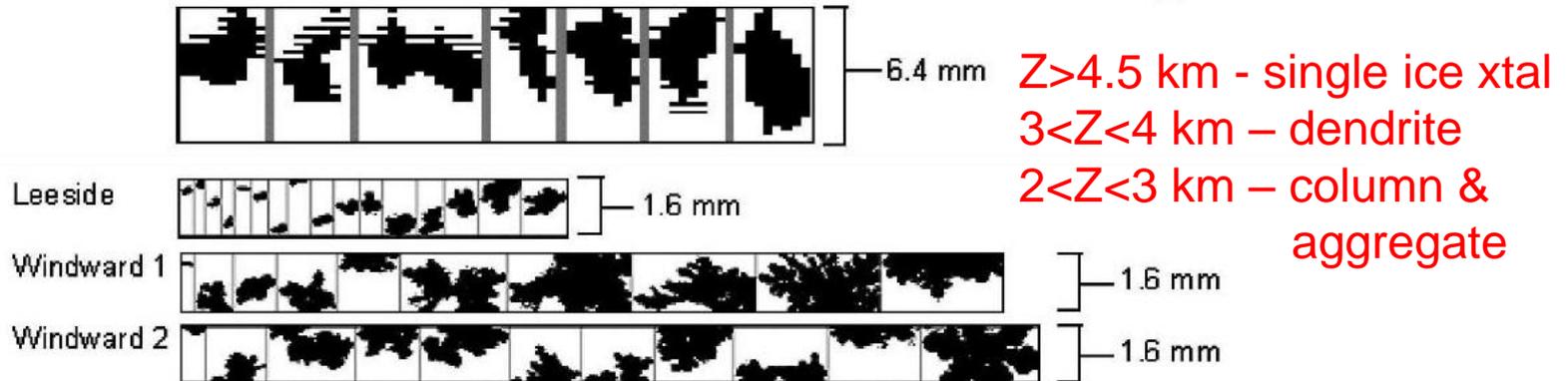
Aircraft flight tracks (2200 – 0200 UTC)



Convair-580

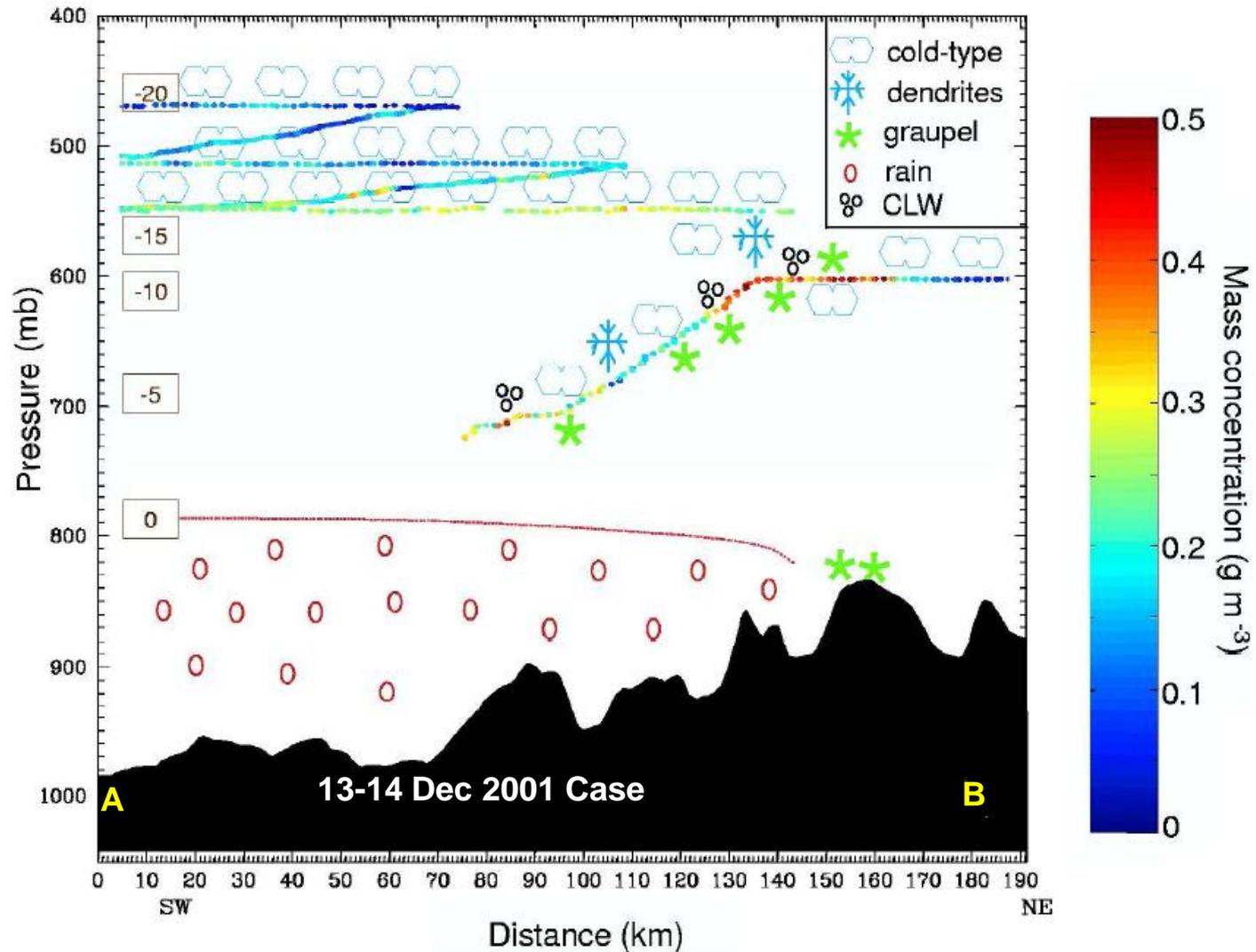


NOAA P-3



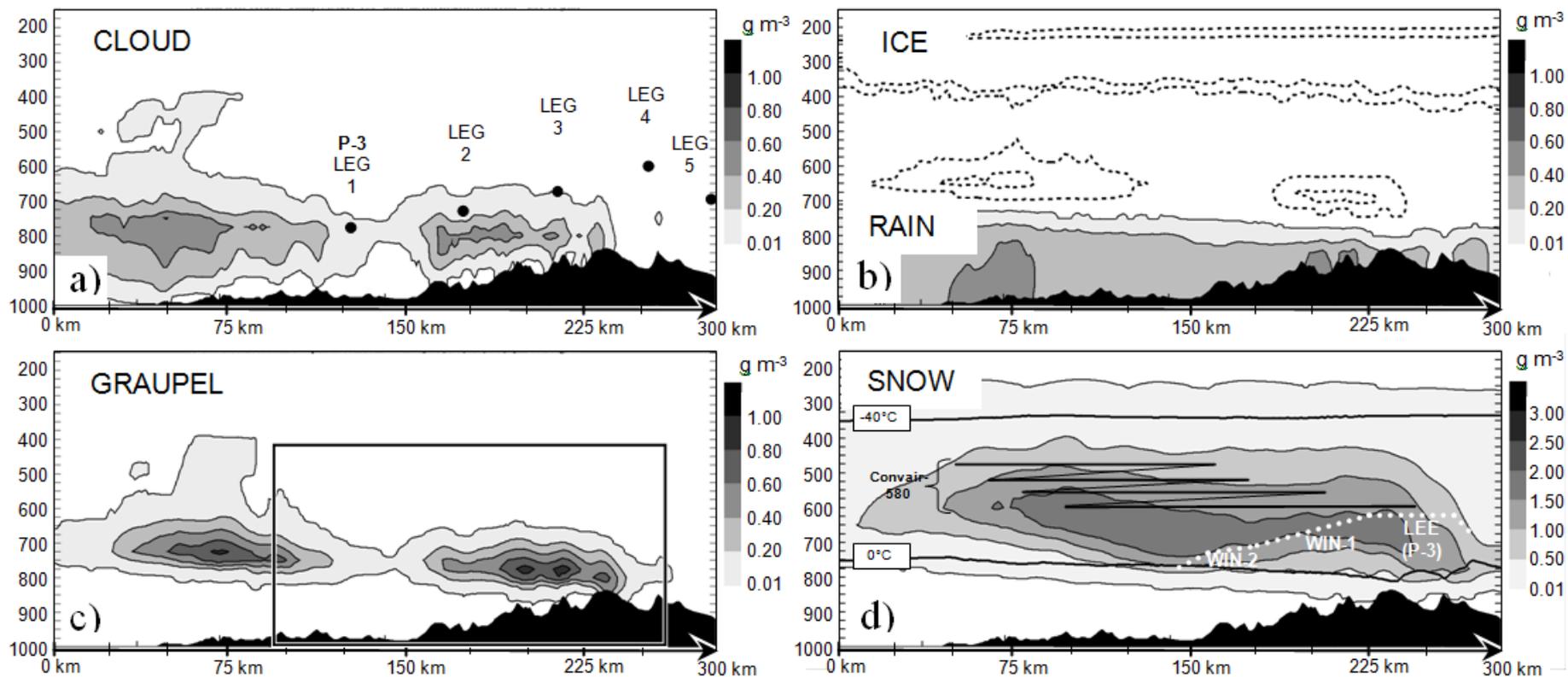
Mean size inc. with dec. height

Combined Observations for 2200–0200 UTC



Q_x [g m⁻³] for 1-km (MY-3) Simulation

Time-Averaged, 2300-0100 UTC



Cloud liquid water along P-3 flight legs

Flight Leg	<i>Valley</i>	<i>Windward</i>		<i>Lee</i>	
	Leg 1	Leg 2	Leg 3	Leg 4	Leg 5
Elevation [m] (Pressure level [hPa])	2000 (775)	2500 (725)	3450 (650)	4000 (600)	3100 (675)
Observation (g m ⁻³) Ave. [Peak]	0.14 [0.40]	0.26 [0.50]	0.20 [0.25]	0.12 [0.15]	0.04 [0.10]
Model (1-km) (g m ⁻³) Ave. [Peak]	0.22 [0.27]	0.08 [0.34]	0.00 [0.09]	0.00 [0.00]	0.01 [0.02]

Under prediction of vertical extent of cloud water

Ice/snow content along Corvair flight legs

Flight Leg	Leg a-b	Leg c-d	Leg e-f	Leg g-h
Elevation [m] (Pressure level [hPa])	6000 (450)	5300 (500)	4900 (525)	4300 (625)
Observed Ave. (g m^{-3})	0.12	0.17	0.25	0.27
Model (1-km) (g m^{-3}) Ave. [Peak]	0.85 [1.34]	0.93 [1.33]	1.15 [1.67]	1.67 [1.94]

Over prediction of concentration of snow mass

→ too large deposition and/or riming

IMPROVEMENTS OF SNOW CATEGORY

- **Diffusional growth**
- **Growth by riming**

Electrostatic Analogy for Diffusional Growth of Ice Crystals

$$\frac{dm}{dt} = \frac{4\pi C(S_i - 1)}{AB_i}$$

“The electrostatic analogy of the capacitance theory of ice crystal growth is highly flawed and does not produce the observed growth rates of ice crystals.

It severely overpredicts the growth rates in almost all cases [by a factor of 3 to 8+ for plates and 2 to 4 for columns] involving even simple hexagonal shapes.”

Bailey and Hallet (2006)

Add **CORRECTION FACTOR** to DIFFUSIONAL GROWTH EQUATION

$$\frac{dm}{dt} = \frac{4\pi C(S_i - 1)}{AB_i} \longrightarrow \frac{dm}{dt} = \frac{4\pi C \cdot f_{corr} \cdot (S_i - 1)}{AB_i}$$

where f_{corr} must be < 1 , with value justified by results

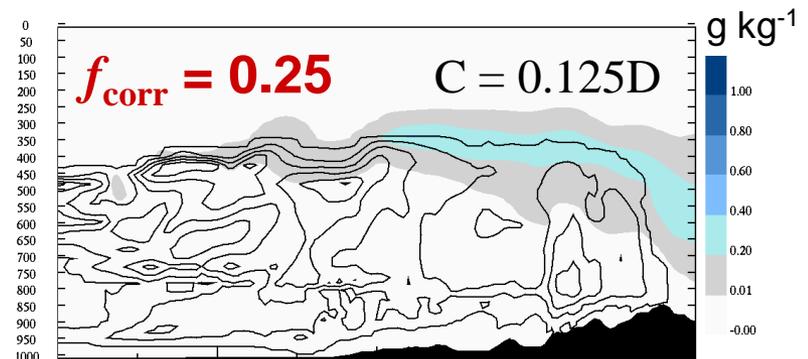
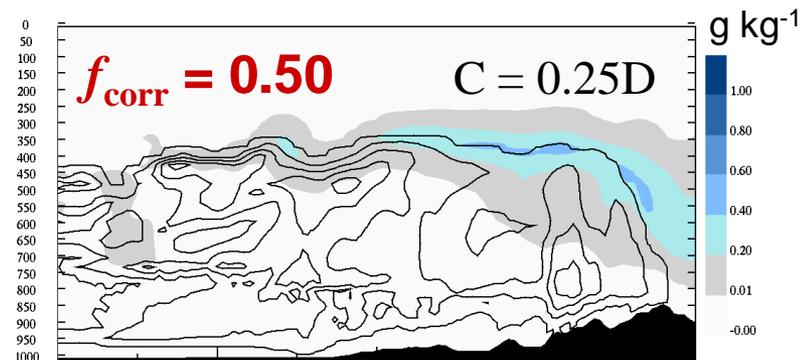
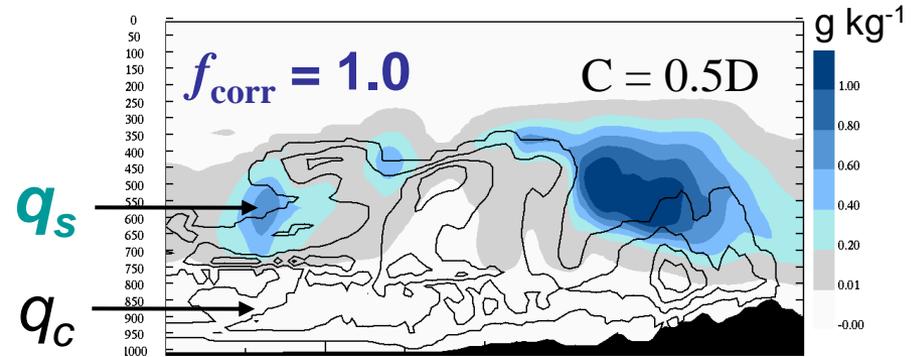
Sensitivity Tests for IMPROVE-2:

With decreasing f_{corr} ,
SNOW content (q_s) is reduced
and
CLOUD LWC (q_c) is increased

Other evidence:

Field et al. (2008)

Westbrook et al. (2008)



RIMING of SNOW

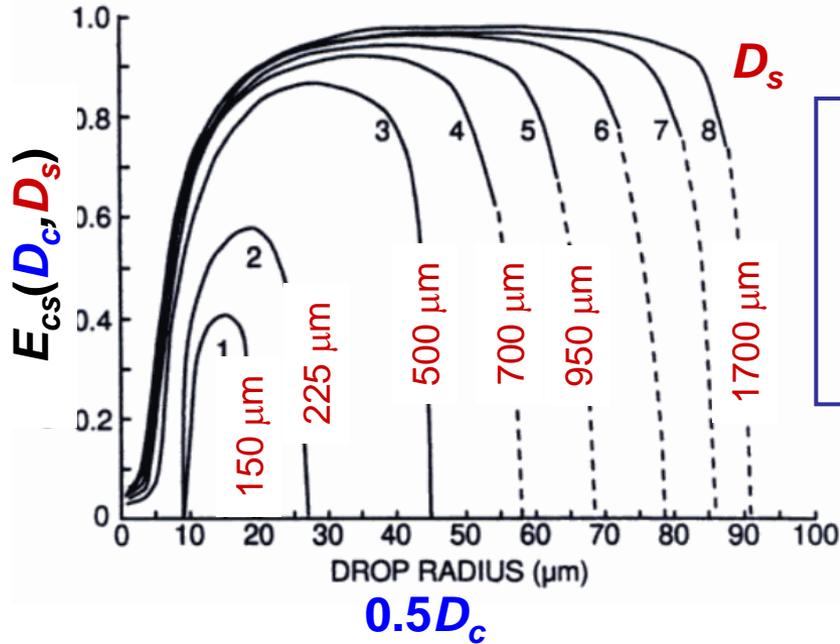
Stochastic collection equation: (for category x collecting category y)

$$CL_{yx} = \frac{1}{\rho} \frac{\pi}{4} \int_0^{\infty} \int_0^{\infty} |V_x(D_x) - V_y(D_y)| (D_x + D_y)^2 m_y(D_y) \underbrace{E_{xy}(D_x, D_y)}_{\text{COLLECTION EFFICIENCY}} N_y(D_y) N_x(D_x) dD_y dD_x$$

**COLLECTION
EFFICIENCY**

- For the collection efficiency, $E_{cs} = 1$ is often assumed (for collection of *cloud* by *snow*)
- If $E_{cs} < 1$, the snow riming rate will be overestimated

RIMING of SNOW



Approximation:

$$E_{cs}(D_c, D_s) = \frac{\min(D_c, 30 \mu\text{m})}{30 \mu\text{m}} \cdot \left[\frac{\min(D_s, 1000 \mu\text{m})}{1000 \mu\text{m}} \right]^{0.5}$$

- Works for $D_c \sim 15\text{-}30 \mu\text{m}$, and $D_s \sim 150\text{-}1500 \mu\text{m}$
- Reduces riming rate 10-80% (vs. $E_{cs} = 1$)

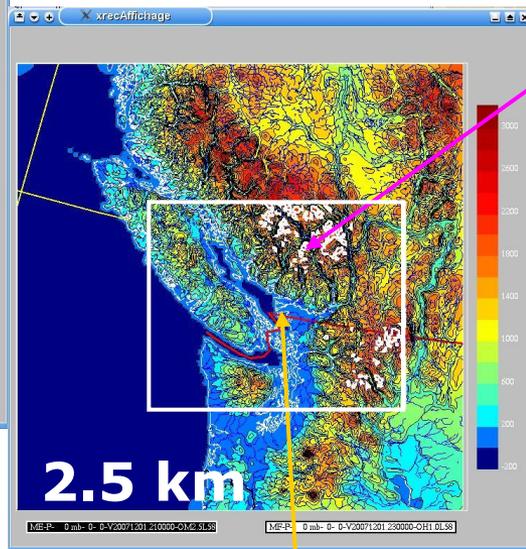
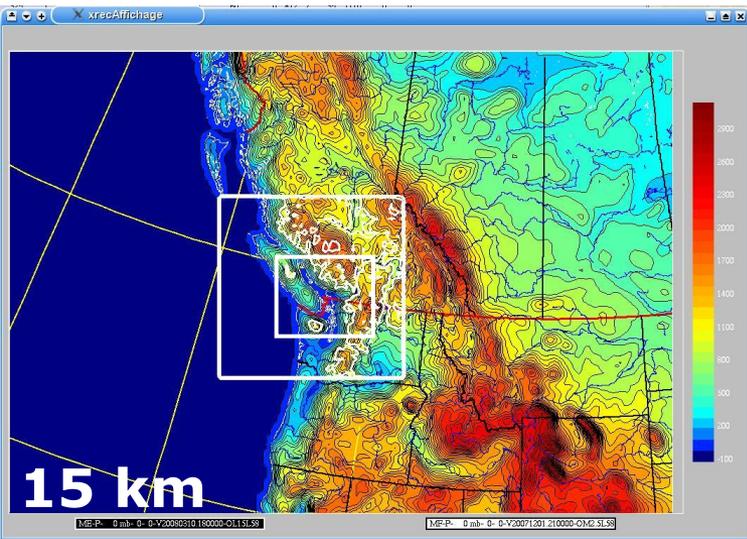
*Wang and Ji, 1992

Test of 2-moment microphysics in
Vancouver Olympics 2010 in 1 km
GEM-LAM

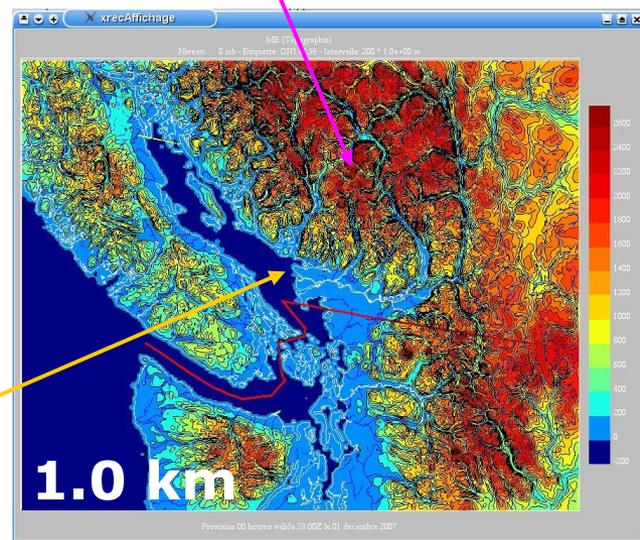
Nesting strategy for LAM-V10 system

- **3 nested LAM integrations twice daily from 0000 and 1200 UTC GEM-Regional forecasts:**

LAM-15 km → 2.5 km → 1 km



Whistler

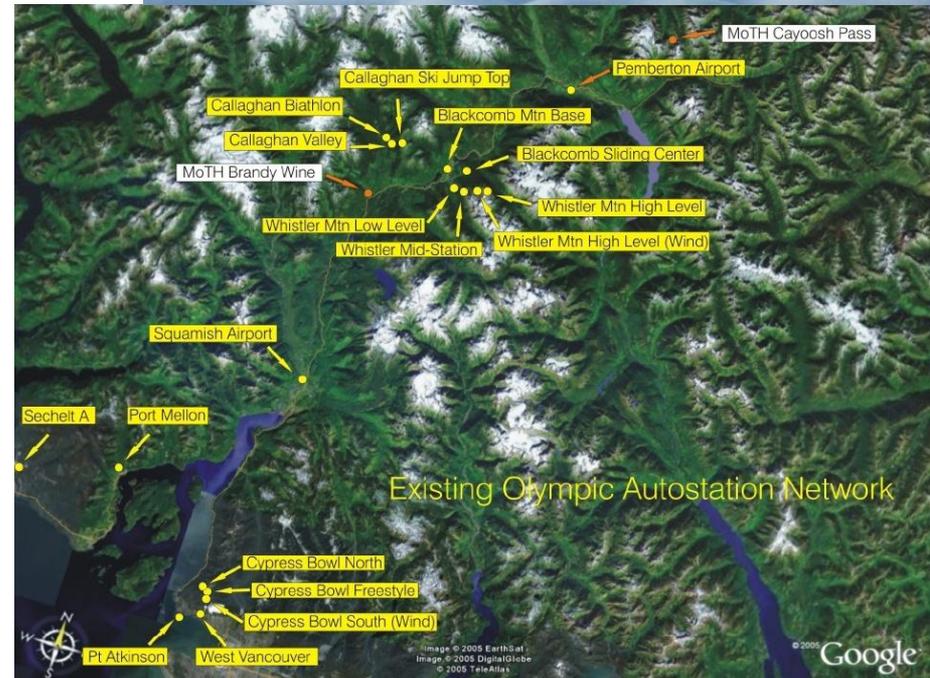


Vancouver

Verification for LAM-V10

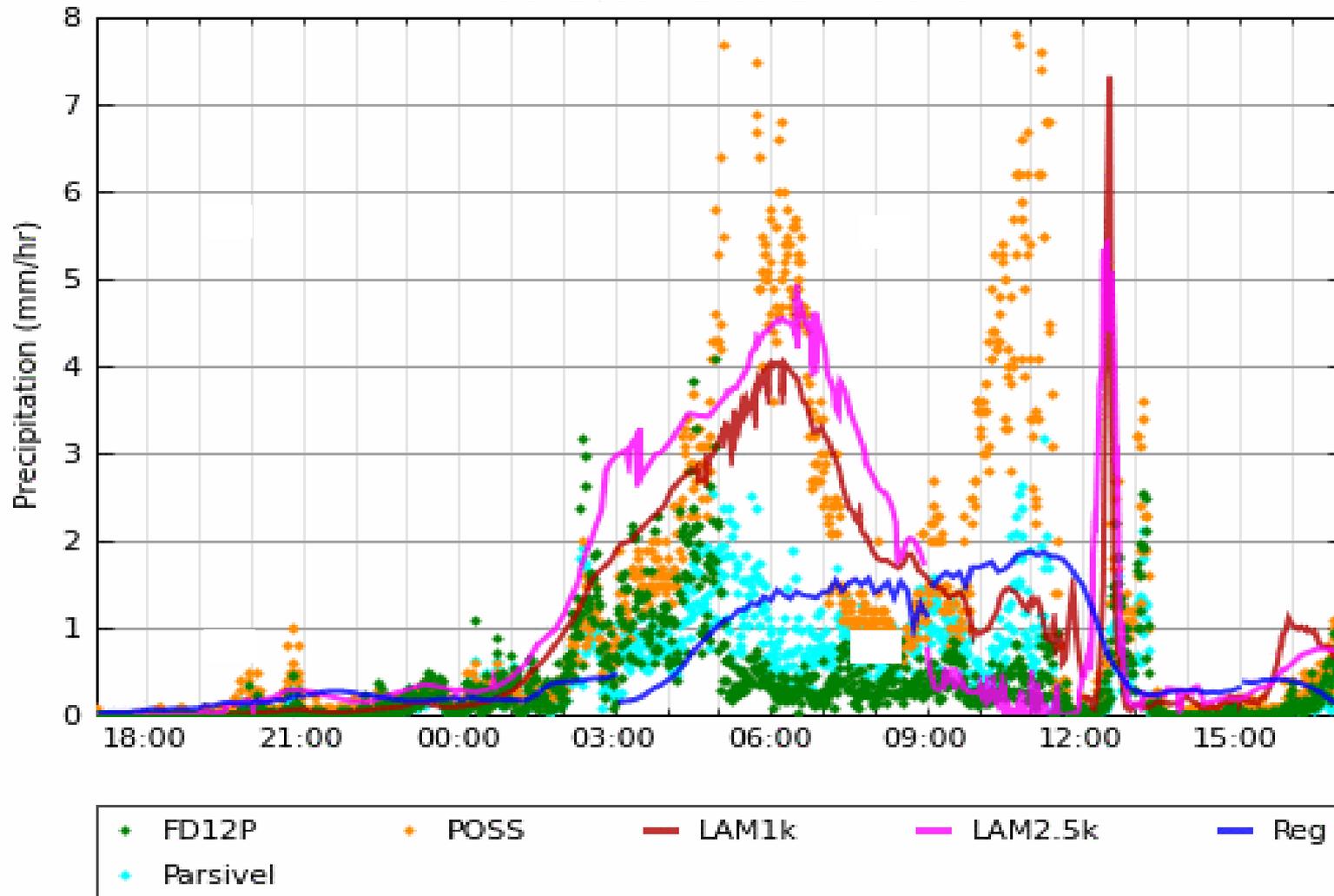
Olympic Autostation Network (OAN):

- approx. 40 standard and special surface observing sites (hourly or synop available on GTS)
- large number (relatively) of surface stations
- concentrated in small region



Verification Examples

VOC: 2010-05-03 17:00 UTC



Observations courtesy of George Isaac

Experimental field: *Solid-to-Liquid ratio*

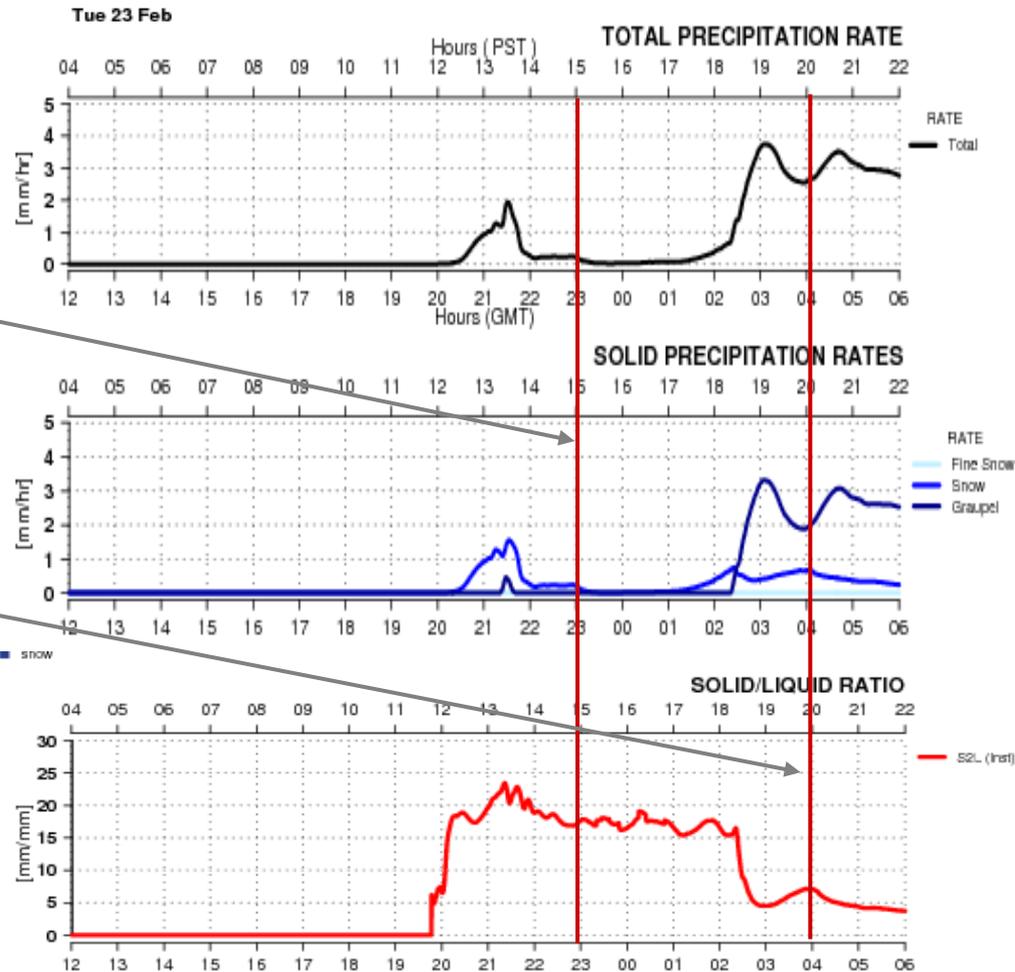
1.0km LAM Model 18 hour Solid Precipitation Meteogram issued 23 February 2010, 12 UTC (04:00 AM local)
Cypress Bowl – South (wind)
 TC ID: VOG LAT: 49.38 N LON: -123.19 W ELEV: 960 m

Observed:*

FLUFFY
SNOWFLAKES

SNOW
PELLETS

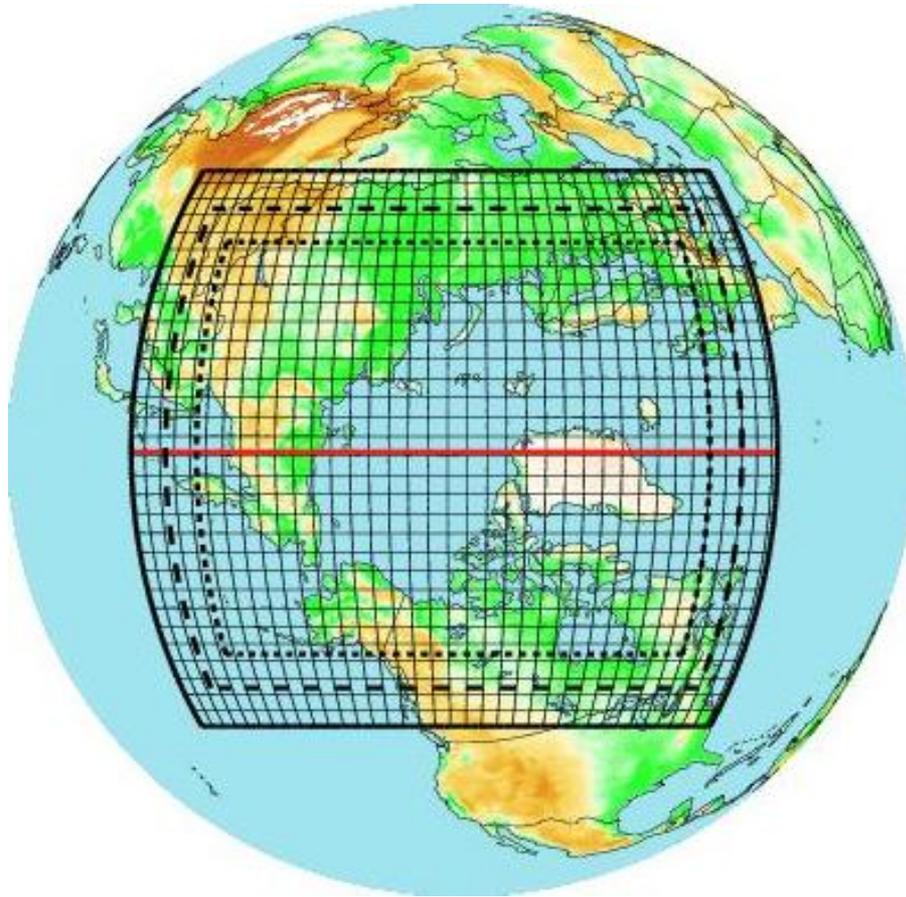
*Forecaster:
Michael Gélinas



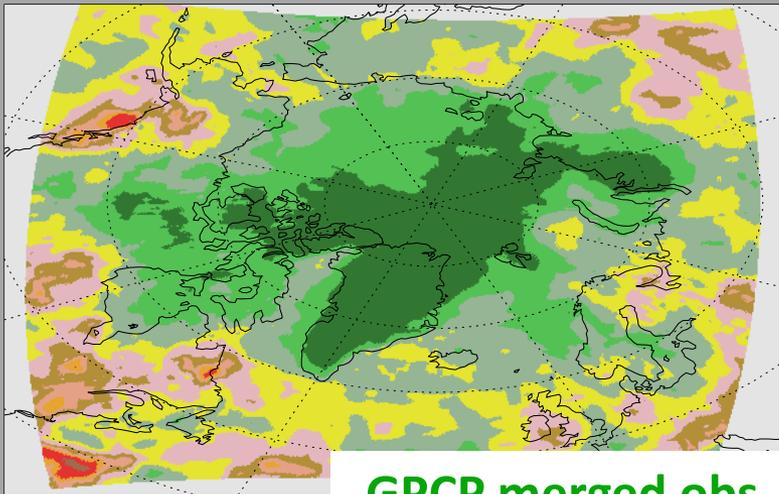
Testing of 2-moment microphysics in
Global GEM variable 15 km over the
Arctic

30 day simulation – July 2008 over Arctic

Polar-GEM:

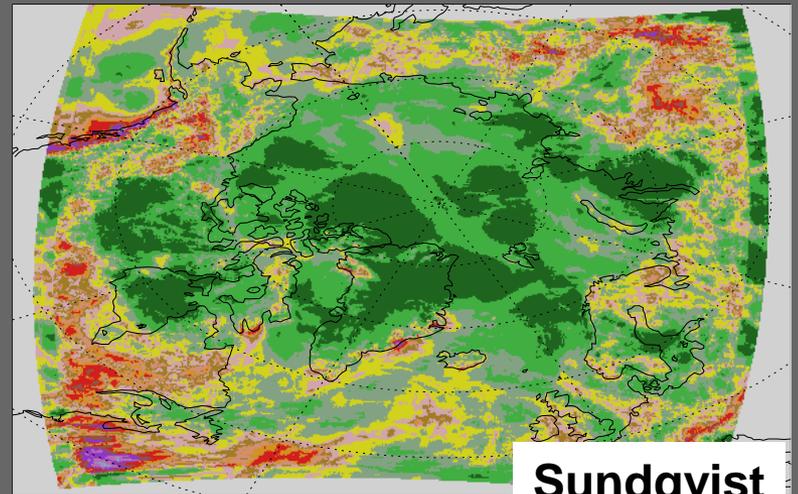
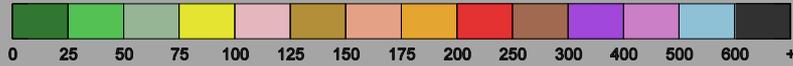


• $\Delta x = 15 \text{ km}$



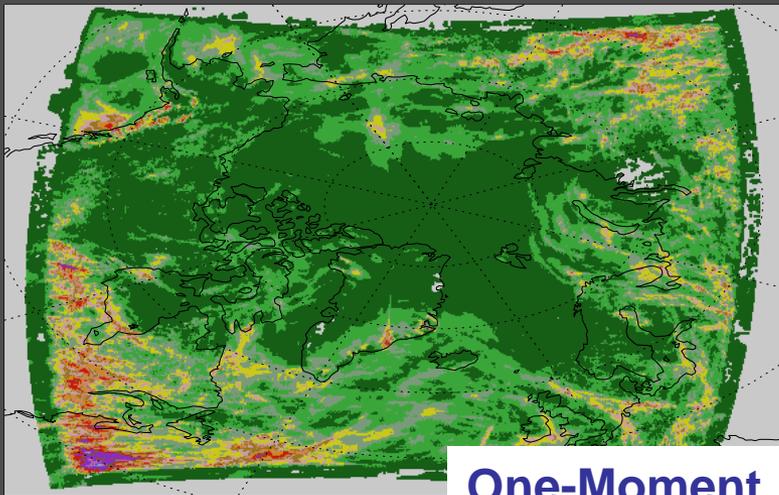
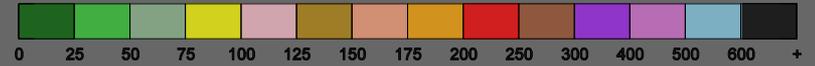
GPCP merged obs

July 2008 GPCP total PR (mm)



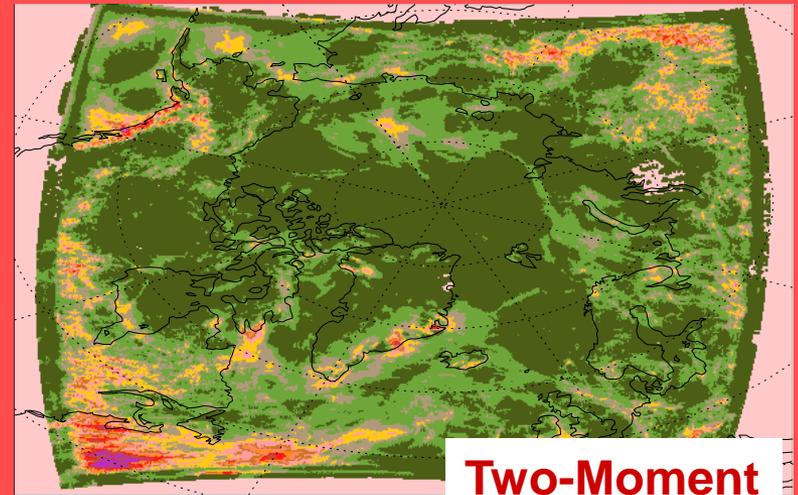
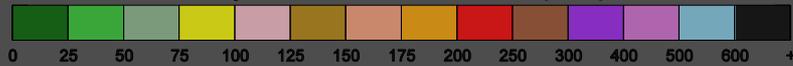
Sundqvist

July 2008 SUND total PR (mm)



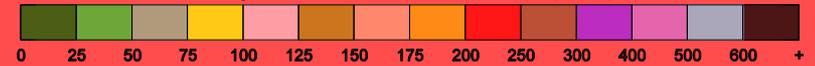
One-Moment

July 2008 MYSM total PR (mm)

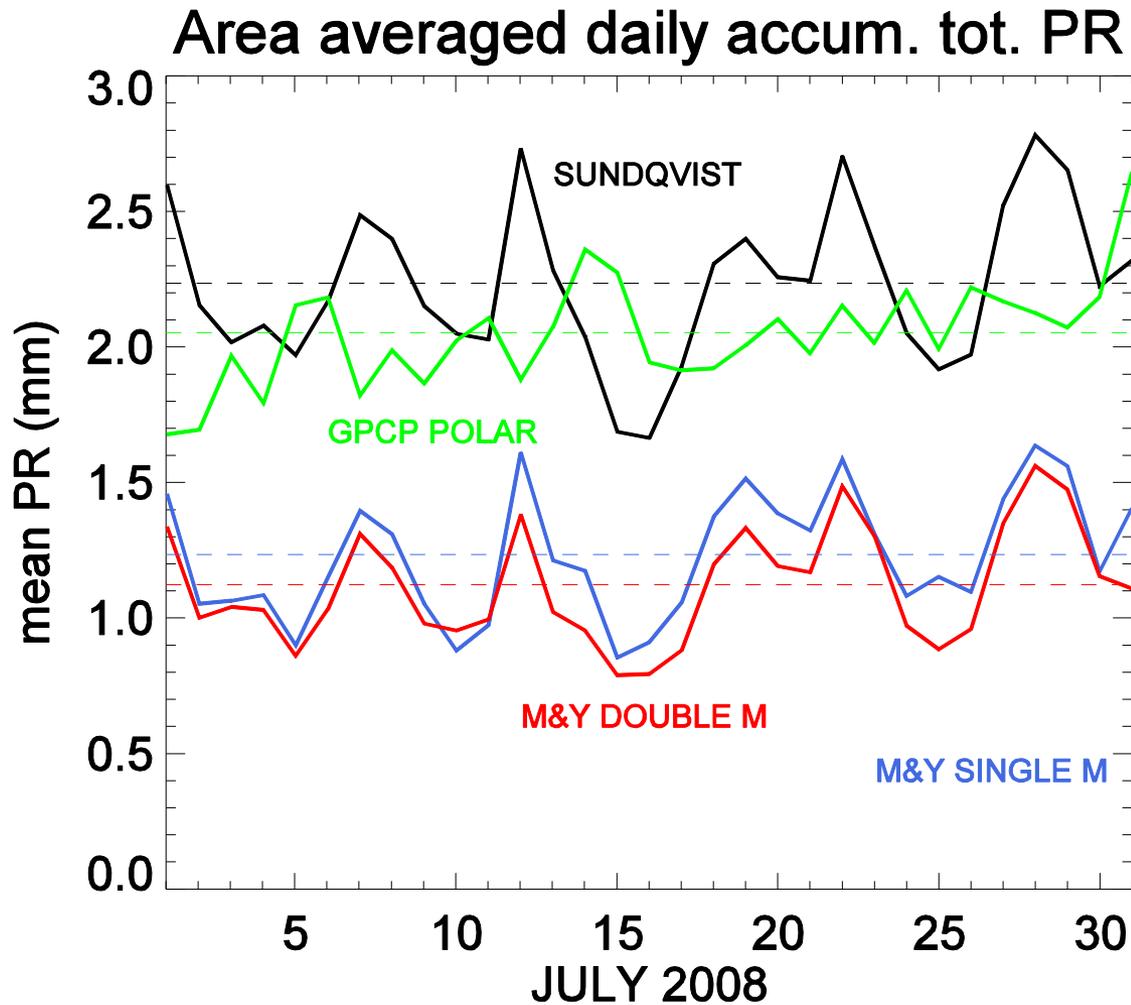


Two-Moment

July 2008 MYDM total PR (mm)



PRECIPITATION



GPCP
merged
obs



Cloud Scheme:

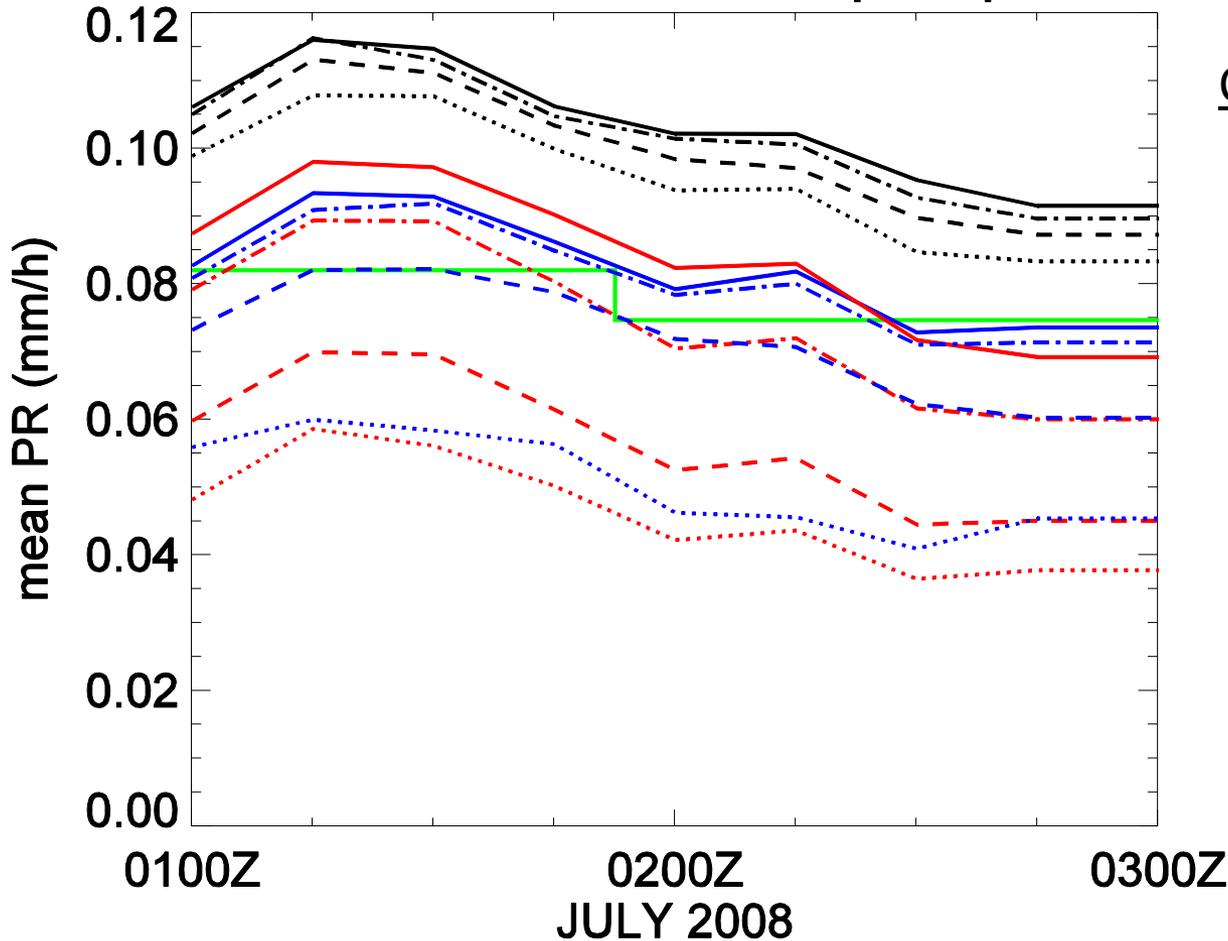
Sundqvist

One-Moment

Two-Moment

SENSITIVITY TO TIME STEP

PR: accum. of total precip.



Cloud Scheme:

Sundqvist

One-Moment

Two-Moment

GPCP
merged
obs



— $\Delta t = 60s$

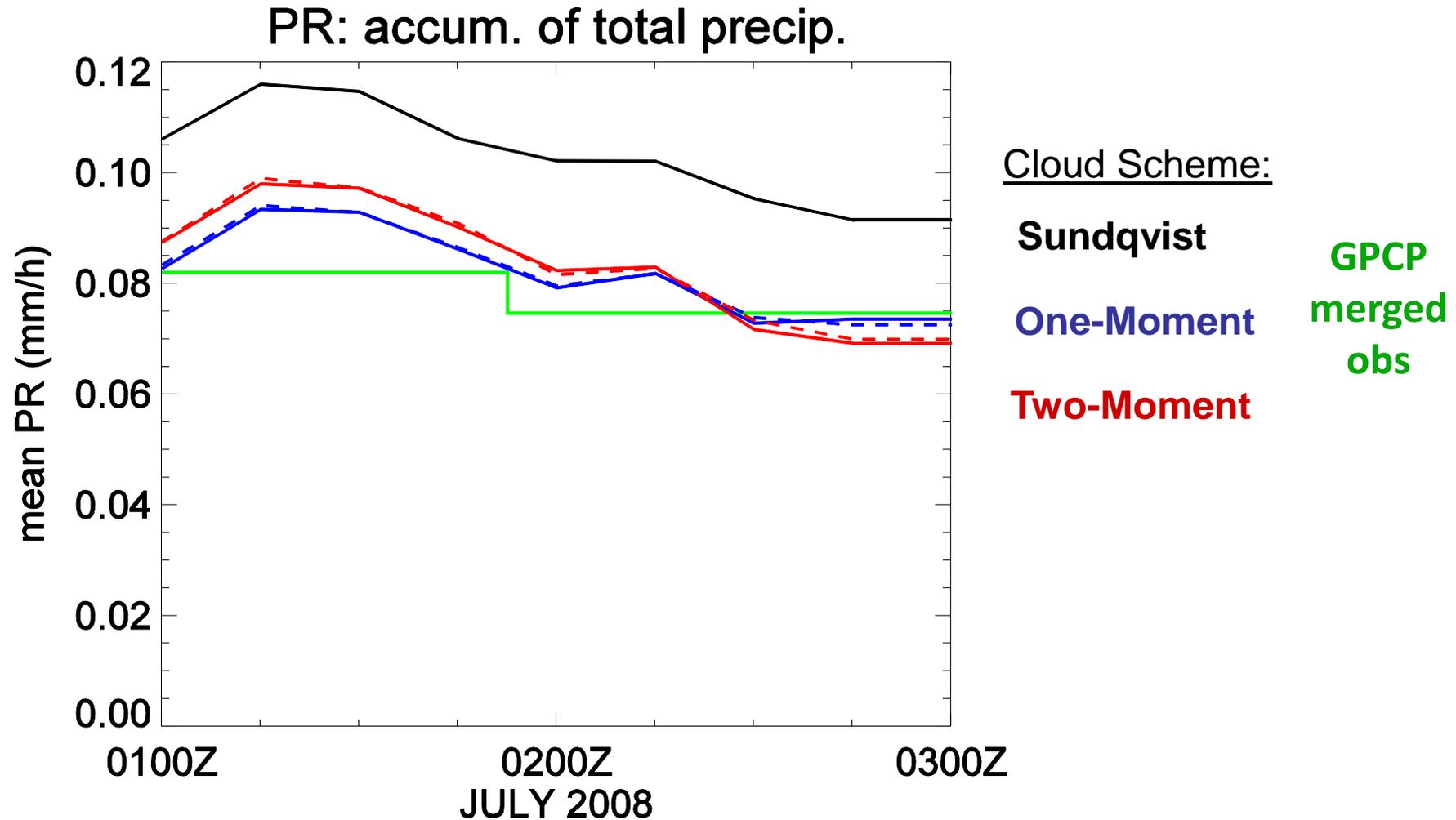
- . - . $\Delta t = 120s$

- - - $\Delta t = 225s$

..... $\Delta t = 450s$

60-h Simulation

SENSITIVITY TO TIME STEP



60-h Simulation ($\Delta t = 60$ s)

SUMMARY

- 1) Multi-moment mixed phase bulk cloud microphysical schemes have been developed and implemented in GEM-LAM and GEM-Global Variable**
- 2) Comparison with in-situ field measurements allows improvements in the scheme**
- 3) Implementation in GEM-Global Uniform is planned but still needs work to address**
 - a) time splitting for microphysics**
 - b) subgrid scale cloud fraction**
 - c) simplification to allow for a mixture of higher and lower moment hydrometeor categories**

SEDIMENTATION: Bulk scheme

$$\left. \frac{\partial \rho q_x}{\partial t} \right|_{SEDI} = \frac{\partial (\rho q_x \bar{V}_{xq})}{\partial z}$$

\bar{V}_{xq} = mass-weighted fall velocity

SM

$$\left. \frac{\partial N_x}{\partial t} \right|_{SEDI} = \frac{\partial (N_x \bar{V}_{xN})}{\partial z}$$

\bar{V}_{xN} = number-weighted fall velocity

DM

$$\left. \frac{\partial Z_x}{\partial t} \right|_{SEDI} = \frac{\partial (Z_x \bar{V}_{xZ})}{\partial z}$$

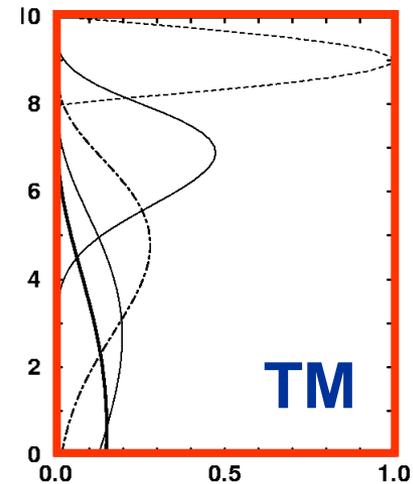
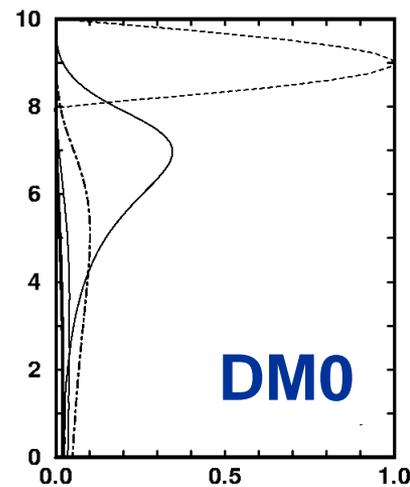
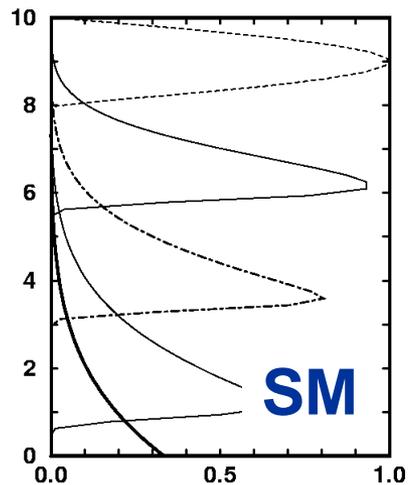
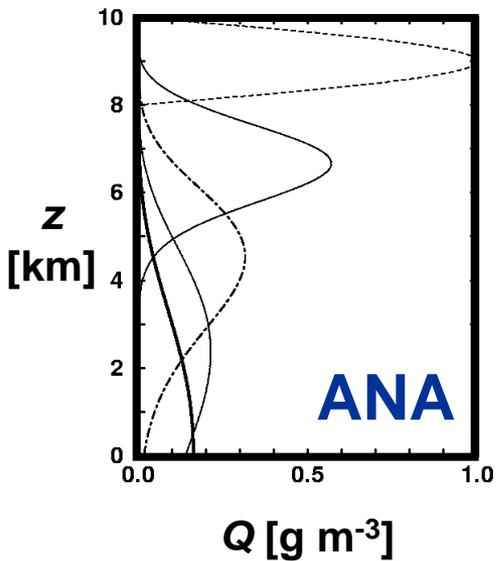
\bar{V}_{xZ} = reflectivity-weighted fall velocity

TM

For a given size distribution, $\bar{V}_{xZ} > \bar{V}_{xq} > \bar{V}_{xN}$

Effects on sedimentation terms

$$(Q = \rho q)$$



TM better than DM0 better than SM

DIFFERENCE RELATED TO SIZE SORTING

Disadvantages of 1-moment scheme

a) Inconsistency in modeling physical processes

From closure relation, N_T and q vary monotonically $\rightarrow N_T$ increases or decreases with q , but

in breakup, N_T increases but $q = \text{constant}$, and
in diffusional growth, q increases but $N_T = \text{constant}$.

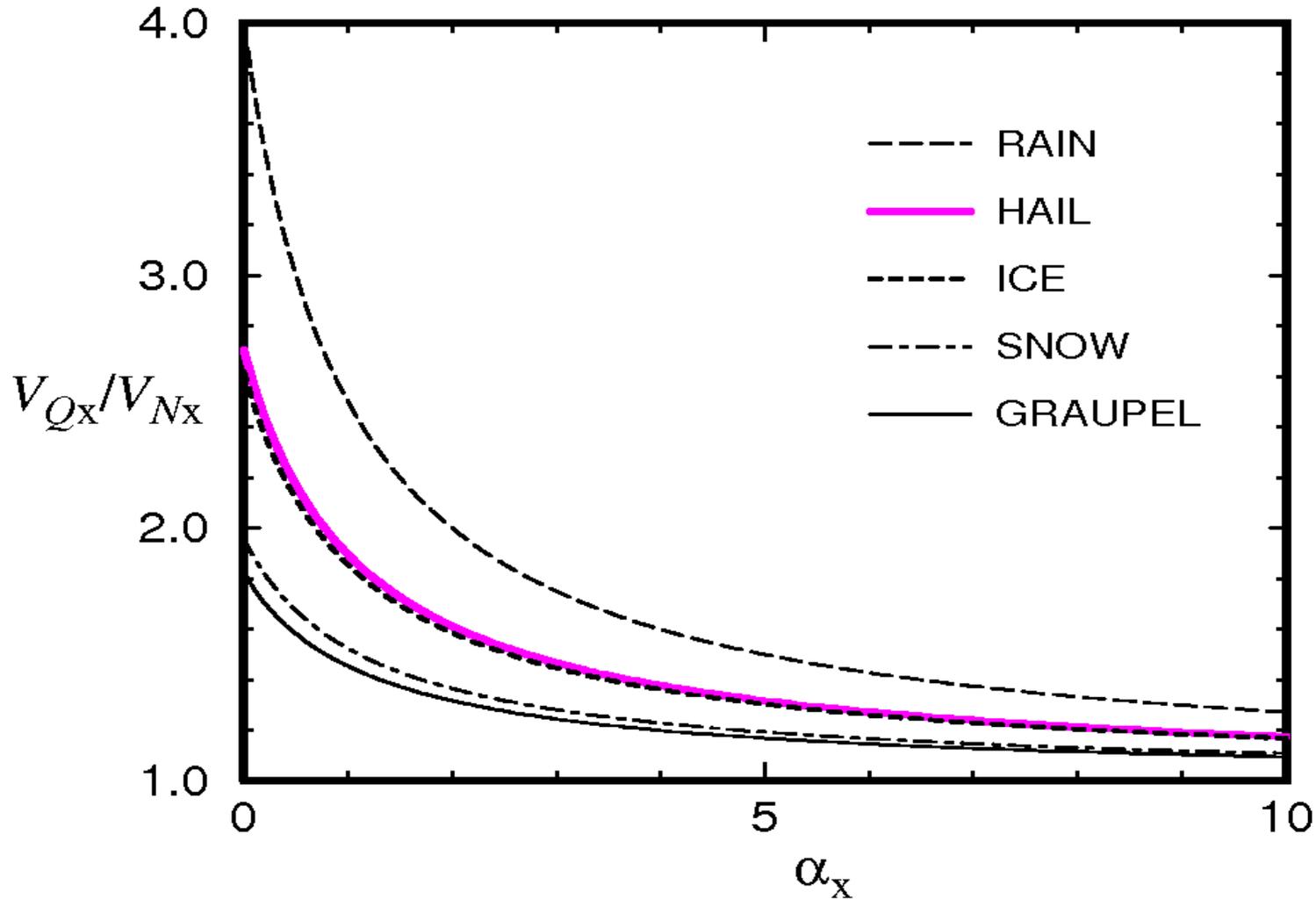
c) Inconsistency in modeling size sorting in sedimentation

\rightarrow mean size increases with decreasing height, but not necessarily true in 1-moment as mean diameter is

$$D_m = \left[\frac{\rho q}{c N_T} \right]^{\frac{1}{3}}$$

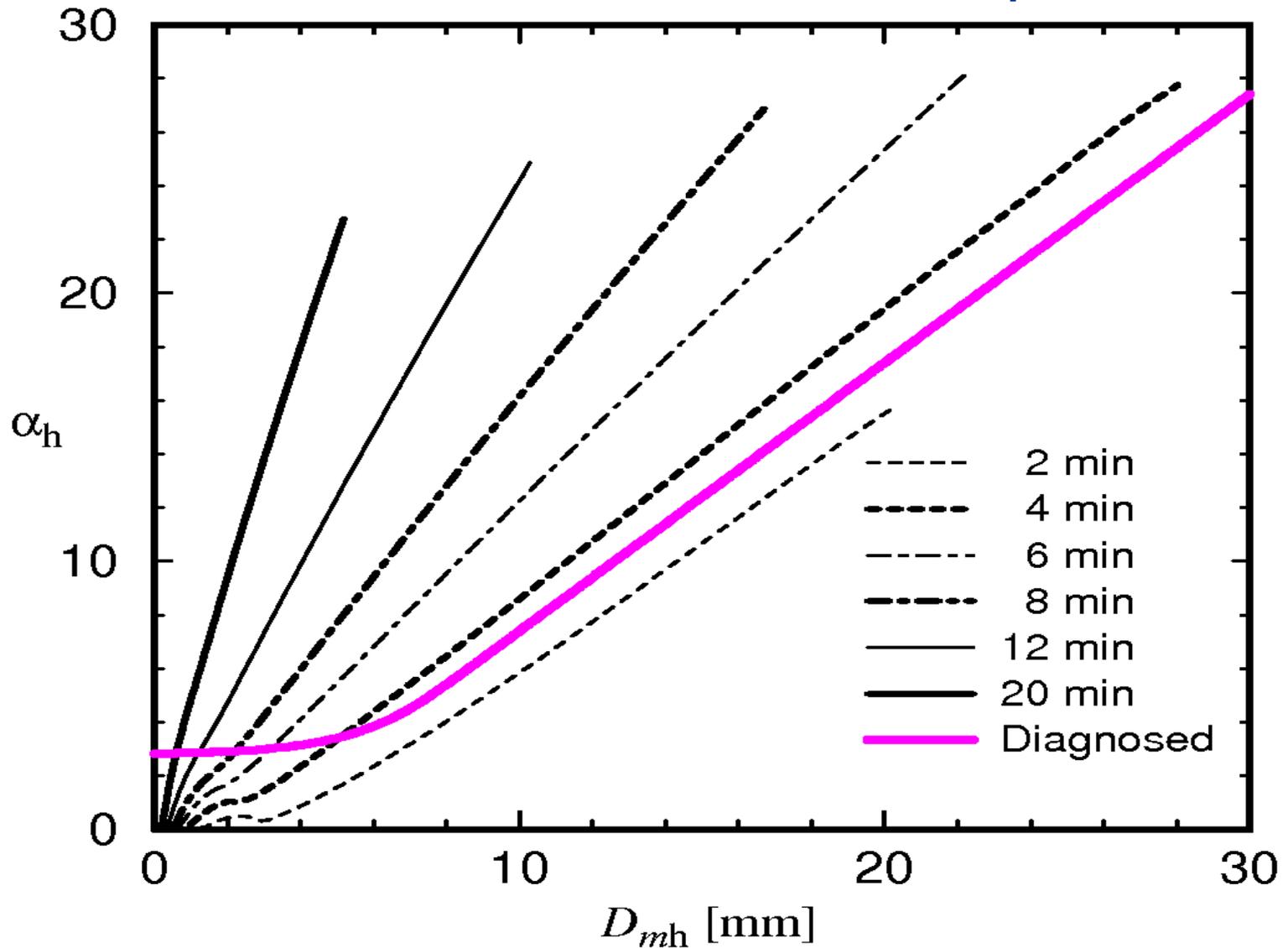
Disadvantages of 2-moment fixed α scheme in sedimentation

Rate of change of D_{mx} (size sorting)
proportional to fallspeed ratio

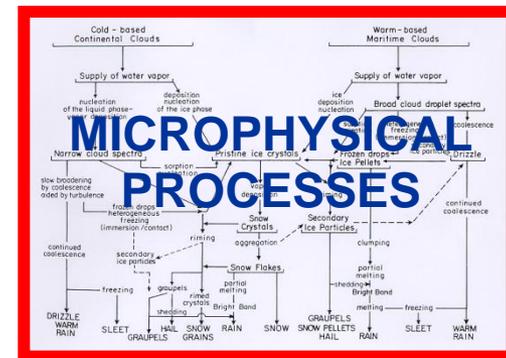


Diagnosed $\alpha \rightarrow$ sedimentation results in larger mean size (larger D_m) but narrower spectrum (larger α)

From TRIPLE-MOMENT sedimentation profiles:



How well do the various bulk scheme predict sources/sinks?



$$\left. \frac{dq_x}{dt} \right|_S = \left. \frac{dq_x}{dt} \right|_{prod} + \left. \frac{dq_x}{dt} \right|_{proc2} + \dots$$

e.g.

$$\left. \frac{dq_x}{dt} \right|_{CL} = \int_0^{\infty} \left. \frac{dm(D)}{dt} \right|_{CL} N(D) dD$$

CONTINUOUS COLLECTION OF CLOUD WATER

$$\left. \frac{dm(D)}{dt} \right|_{CL} = \frac{\pi D^2}{4} V(D) E_{xc} \rho q_c = \left(\frac{\pi}{4} E_{xc} \rho q_c \right) D^{2+b_x}$$

$$\left. \frac{dq_x}{dt} \right|_{CL} = \left(\frac{\pi}{4} E_{xc} \rho q_c \right) \int_0^{\infty} D^{2+b_x} N(D) dD$$

p^{th} moment:

$$M_x(p) \equiv \int_0^{\infty} D^p N_x(D) dD$$

$$\left. \frac{dq_x}{dt} \right|_{CL} \propto M_x(2 + b_x)$$

$$V_x(D) = \gamma a_x D^{b_x}$$

How well do the various bulk scheme predict sedimentation and sources/sinks?

TM and DIAG DM schemes

better than

SM AND FIXED DM schemes