The assimilation of cloud affected IR radiances

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Outline of talk…

• Issues related to the assimilation of cloudy data

• The scheme developed at ECMWF

• Performance of the cloudy scheme

• Summary and view to the future
Why do we want to use cloud affected data?

1. Only using clear-sky data represents a major under-use of high cost instruments such as AIRS and IASI.

2. It is important to constrain analysis errors in cloudy regions as they are believed to be meteorologically sensitive.
Sensitive areas and cloud cover

Location of sensitive regions
Summer-2001 (no clouds)

sensitivity surviving high cloud cover

sensitivity surviving low cloud cover

monthly mean high cloud cover

monthly mean low cloud cover

From McNally (2002) QJRMS 128
Two potential approaches to handle clouds

1. Use cloud affected radiance observations that have been pre-corrected to remove the cloud signal (i.e. *cloud cleared data*)

2. Extend the NWP analysis to estimate cloud parameters simultaneously with temperature and humidity (*either interacting with model cloud physics or not*)
Two potential approaches to handle clouds

1. Use cloud affected radiance observations that have been pre-corrected to remove the cloud signal (i.e. *cloud cleared data*)

2. Extend the NWP analysis to estimate cloud parameters simultaneously with temperature and humidity (*either interacting with model cloud physics or not*)
Fundamental issues

- The cloud uncertainty in radiance terms may be an order of magnitude larger than the T and Q signal (i.e. 10s of kelvin compared to 0.1s of kelvin).

- The radiance response to cloud changes is highly non-linear (i.e. $H(x) = H_x(x)$).

- Errors in background cloud parameters provided by the NWP system may be too large to provide an accurate linearization point and very difficult to model.

- Trade off between having enough cloud variables for an accurate RT calculation while limiting the number of cloud variables to those that can be uniquely estimated in the analysis from the observations.
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Observed radiance at 11 microns minus radiance calculated in clear sky

Large cold departures indicating cloud contamination in OBS

Clear population (dTb < 0.5K)
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Clear and Cloudy Jacobians

\[ \frac{dR}{dT^{500}} = 0 \quad \text{dR/dT}^* = 1 \]

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full cloud at 500hPa
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Observed radiance at 11 microns minus radiance calculated in clear sky
Observed radiance at 11 microns minus radiance calculated from **NWP cloud background profile**

Many clouds with significant radiance signals are accurately represented by the NWP model and RT modelled!
Observed radiance at 11 microns minus radiance calculated from **NWP cloud** background profile

Many clouds with significant radiance signals are accurately represented by the NWP model and RT modelled!

...but the spread is still very large!
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- Trade off between having enough cloud variables for an accurate RT calculation while limiting the number of cloud variables to those that can be uniquely estimated in the analysis from the observations
A very simple cloud model (e.g. single layer grey cloud amount and pressure) should more readily estimated from the data (independently of T and Q), but will make the forward RT calculation very inaccurate in many cloud conditions.

A more complex cloud model (e.g. cloud liquid and ice at each model level) will allow a more accurate forward RT calculation, but the parameters may be difficult to estimate independently of each other and T and Q and may alias into erroneous increments.
Extension of the ECMWF 4D-Var to assimilate cloud affected infrared radiances

Described in McNally 2009 QJRMS
Key features of the cloudy scheme

- Only cloudy IR radiances from completely overcast scenes are used
- One additional variable (local) added to 4D-Var control vector ($P_{CTOP}$)
- Background values estimated from the observations (not NWP model)
- QC rejection of marine inversion / physically unreasonable clouds
- All IR sensors treated identically (AIRS / IASI / HIRS)
- Cloud information not fed back to NWP model
Why overcast scenes...?
Why use cloudy radiances only in overcast conditions?

- Overcast conditions are least ambiguous in the radiance data\(^*\).
- Cloud control vector collapses to a single number \(P_{\text{CTOP}}\).
- Problems with cloud overlap assumptions vanish.
- Termination of jacobians at cloud top provides new information\(^*\).
- We can measure temperature above clouds better than in clear sky.
- No cross-talk between cloud and surface skin sink variables.
Error in estimation of cloud top pressure

Error decreases as cloud fraction increases

Error decreases for higher clouds

Thanks Andrew…
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- **Termination of jacobians at cloud top provides new information**

- **We can measure temperature above clouds better than in clear sky**

- No cross-talk between cloud and surface skin sink variables
Enhanced temperature estimation at the cloud top

$dR/dT_{500} = 0$

$dR/dT^* = 1$

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full cloud at 500hPa
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Background cloud parameters...
Why not use the NWP model for background cloud parameters?

The disagreement between the OBS and the model is not excessive, but still large enough to often stretch the TL approximation and limit convergence.

There also a difficulty in post-processing the model cloud profile variables to the quantity representative of that seen by the radiance observations.

CTOP: NWP minus 2D least squares

Observed radiance at 11 microns minus radiance calculated from NWP cloud background profile

70hPa bias!
Background cloud parameters
(2D least squares method)

We find \( N \) (cloud fraction) and \( P \) (cloud top pressure) which minimize the squared radiance departures summed over \( J \) (currently \( J=3 \)) channels:

\[
\sum_j \delta^2 = \sum_j [(R_j^{obs} - R_j^\circ) - N(R_j^\ast(p) - R_j^\circ)]^2
\]

Analytically solving for \( N \):

\[
N = \frac{\sum_j [(R_j^{obs} - R_j^\circ)(R_j^\ast(p) - R_j^\circ)]}{\sum_j [R_j^\ast(p) - R_j^\circ]^2}
\]

and numerically finding the value of \( P \) that gives the overall minimum departure.
Qualitatively – the location and altitude of overcast locations seems reasonable when compared to MODIS equivalent products.
Background 2D cloud parameters (comparison with AVHRR)

AVHRR cluster analysis based on imager pixels within the IASI field of view – one week of data 2008-08-07 to 2008-08-14

IASI data identified as overcast have very low AVHRR variance
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• All IR sensors treated identically (AIRS / IASI / HIRS)

• Cloud information not fed back to model
Quality Control...
Problem in MSC regions / inversions

Model cloud cover

Strong inversions confuse the CTP estimation which puts the cloud too high ...thus leaving a positive residual in sounding channels...

Satellite puts cloud here

Temperature profiles

The inversion means that both of these cloud have the same emission ...

Temperature increments

Note: there is some LIDAR evidence to suggest the model clouds are too low in the (SH) MSC regions and thus the associated model temperature / humidity profile (from which initial cloud parameters are computed) is unlikely to be correct!
Problem in MSC regions / inversions

Model cloud cover

Vertical profile of cloud cov. 20060111 02 step 24 Envizor 0201 point (-30.0,105.0)

Strong inversions confuse the CTP estimation which puts the cloud too high ...thus leaving a positive residual in sounding channels...

Satellite puts cloud here

Temperature profiles

The inversion means that both of these must have the same emission ...

Temperature increments

Vertical profile of temp. 20060111 2100 step 0 Envisat eci point (+50.0,105.0)

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Non-physical cloud solutions

Initial cloud fraction estimates are sometimes found to stray outside **physical bounds** (i.e. $0 < N_e < 1$). These are removed as a QC step from further assimilation (as they may indicate multi-level cloud situations and show a degraded fit to the observations)

IASI Tb  S.Hemis
used Tb METOP-A IASI
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used Tb METOP-A IASI

STD. DEV
Cloudy assimilation system applied to combined HIRS / AIRS / IASI
Experiment design

Period = 3 months in January/February/March 2008

Resolution = T255

HIRS radiances from METOP-A and NOAA-17 used (LW)
AIRS radiances from AQUA used (LW/WB/SW)
IASI radiances from METOP-A (LW)

CNTRL = ECMWF operations (clear channels from HIRS / AIRS / IASI)

EXPT = CNTRL + HIRS / AIRS / IASI in overcast locations

Background cloud conditions from 2D least squares fit to 4 channels

Background errors CTOP = 5hPa and CFRAC = 0 (local sink variables)

QC applied rejecting low clouds and “bad” 2D solutions
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Where are the extra data?

Combined clear data coverage of mid/lower tropospheric sounding radiances:

- IASI channel 434 (METOP-A)
- AIRS channel 355 (AQUA)
- HIRS channel 7 (NOAA-17 / METOP-A)

Additional overcast locations where cloudy radiance analysis fills gaps due to cloud detection rejections:

- IASI channel 434 (METOP-A)
- AIRS channel 355 (AQUA)
- HIRS channel 7 (NOAA-17 / METOP-A)

Typically the overcast locations only provide an extra 10% to the total data.
Impact of overcast data on the analysis...
Temperature increments above low clouds

Overcast data coverage and Cloud top pressure

Analysis temperature increments at 700hPa
Temperature increments above high clouds

Overcast data coverage and Cloud top pressure

Analysis temperature increments at 250hPa
Reduced temperature increments at isolated observation locations

Monthly averaged RMS temperature increment difference (CLOUDY minus CTRL). Shaded areas indicate a reduction in increments in excess of 0.1K when the cloud radiances are assimilated.
...remember this ...

\[ \frac{dR}{dT^{500}} = 0 \]

\[ \frac{dR}{dT^*} = 1 \]

full cloud at 500hPa

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Temperature increments at the cloud top

Cell of very high overcast clouds off the coast of PNG seen by IASI

All IASI channels collapse to near delta-functions at the cloud top giving very high vertical resolution temperature increments just above the diagnosed cloud.
Impact of overcast data on forecasts …
Generally the impact of the extra overcast data on the hemispheric forecast error scores is neutral or slightly positive (with no statistical significance).

However, statistically significant forecast impacts are obtained in the Tropics where temperature forecasts are improved at all ranges.
Cloud obscured singular vector?

In this case the use of overcast observations resulted in analysis differences in an area suggested to be sensitive by the singular vector locations.

Extra overcast data used compared to CTRL

500hPa temperature analysis difference (K)

Location of leading 500hPa singular vectors

SH 500hPa Z

CNTRL CLOUDY
Summary

- The ECMWF 4D-Var has successfully been extended to make additional use of overcast radiance data.

- The restriction to overcast scenes and the strict QC currently yields < 10% extra radiance data.

- The small amount of additional data do not significantly influence the bulk characteristics of the analysis or departure statistics – although some isolated reduction of increments is observed.

- At locations where there are extra radiance observations - high vertical resolution increments (above overcast cloud top) look reasonable, but need further detailed validation.

- No statistically significant impact on forecast performance apart from improved Tropical temperature scores.
Next Steps

- Use imager data (MODIS/AVHRR) to validate 2DLS background cloud estimates and investigate the possibility of using imager identification of overcast scenes for data selection / QC rejection

- Use CLOUSAT data to validate the 2DLS background cloud top estimates in overcast conditions (particularly MSC)

- Continue to search for individual cases of forecast impact – possibly using singular vectors or adjoint sensitivity diagnostics

- Investigate use of a post-processed NWP cloud background for the cloudy IR analysis to replace the 2DLS

- *Investigate the options for feeding the cloud information back to the model physics* (e.g. via cloud fraction ?)

- *Understand how this approach to using cloudy data blends (or not) with other future developments* (rainy radiances)
End
Some questions?

- What are the implications of channels used in the 2DLS and then in the 4DVAR (potentially all T/Q information could be removed by inserting a cloud. Is the problem biggest for HIRS and does the overcast limitation help?

- Can we make better use of post-processed NWP cloud parameters to provide independent background for cloud analysis?

- Can we make use of imager cloud information – either as a background of to at least verify other background cloud parameters (2DLS or NWP) – or as a QC mechanism

- Must study the (O-B) stats for unambiguously clear data and cloudy data with the 2DLS estimated cloud signal removed. If the latter is very small it suggest that a lot of T/Q signal is being dumped into cloud in 2DLS.

- The neutral forecast impact – is it a mix of good and bad or just small?

- Two possible sources of improvement in analyses and forecasts?
  - dumping erroneous signal from cloud detection into cloud sink variable
  - real new useful T/Q information above overcast clouds.
The effect of T,Q error on the estimation of cloud top pressure

Realistic errors placed on T,Q from B for the simultaneous estimation of cloud top pressure from AIRS / IASI

However, the cloud top pressure estimate is not significantly affected by these T,Q errors and the accuracy is similar even when the T,Q are known perfectly
Experiments with overcast HIRS data only

Monthly mean (model) low cloud cover

Temperature increments

Improved fit to isolated TEMP data (averaged over 1 month)
Experiments with overcast HIRS data only

Forecast scores averaged over 1 month generally neutral – but some improvement in S. Hemisphere short-range forecasts at the 95% significance level.
Extra IR data from overcast locations
(after QC typically < 10% shown in red)

Note that the extra overcast radiances do not alter the overall (O – B) statistics
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surface

Choice of cloud parameters and ambiguity with T and Q

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... note that fundamentally the estimation problem is the same ...