Numerical Weather Prediction at high latitudes

Nils Gustafsson, SMHI
High latitude NWP - outline of talk

• The HIRLAM-A program
• The reference HIRLAM forecasting system
• Verification scores
• The Nordic temperature problem – a new snow scheme
• (Turbulence in stable boundary layers)
• Snow band simulations (very old slides!)
• Data assimilation developments (4D-Var, AMSU-A and AMSU-B over sea ice)
Material for this talk was contributed by:

Tage Andersson, Per Dahlgren, Stefan Gollvik, Xiang-Yu Hans Huang, Lars Meuller, Patrik Samuelsson, Harald Schyberg, Sander Tijm, Vibeke W. Thyness, Frank T. Tveter and Xiaohua Yang
Regional NWP groups in Europe

<table>
<thead>
<tr>
<th>Region</th>
<th>Count</th>
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<tbody>
<tr>
<td>ALADIN</td>
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<tr>
<td>COSMO</td>
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<tr>
<td>HIRLAM</td>
<td>8</td>
</tr>
<tr>
<td>UK</td>
<td>1</td>
</tr>
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</table>
Three recommendations by the international evaluation of HIRLAM - strategy for the HIRLAM-A program

• Continue develop synoptic scale forecasting systems based on a merge of the HIRLAM and ALADIN systems.
• Develop a meso-scale forecasting system in collaboration with ALADIN
• Develop EPS, first for synoptic scale forecasting

(shorter than the original text)
The HIRLAM synoptic scale forecasting system

• The semi-implicit, semi-Lagrangian grid-point model
• The 3-Dimensional Variational data assimilation; 3D-Var $\rightarrow$ 4D-Var
• The incremental Digital Filter initialization (to disappear with 4D-Var)
Physical parameterizations in the HIRLAM gridpoint forecast model

• CBR turbulence (Cuxart et al.)
• STRACO condensation, clouds and convection (Sass); Rasch-Kristjansson + Kain-Fritsch as an option (used by SMHI)
• Savijärvi-Sass-Rontu radiation
• ISBA surface and soil → new scheme including canopy temperature and snow
Use of observations in HIRLAM

- **Operationally**: TEMP, PILOT, SYNOP, SHIP, AIREP, AMDAR, DRIBU, SATOB and AMSU-A radiances
- **Being implemented**: Scatterometer winds
- **Trials with**: AMSU-B radiances, AMSU-A/B over ice and land, groundbased GPS (zenith and slant delays), MODIS winds, radar radial winds
Verification of 2 m temp. forecasts against Swedish SYNOPs
Verification of 2 m temp. forecasts against Swedish SYNOPs (after post-processing with Kalman filter)
Verification of precip. forecasts against Swedish SYNOPs

![Graph showing precipitation mean absolute error and mean error over Sweden from September 2005 to August 2006, comparing forecasts from various models.](image)

- **MAE mm/12 hours**
  - Hirlam11 12-030 MAE = 0.833
  - Hirlam22 12-030 MAE = 0.682
  - Ecmwf 00+042 MAE = 0.723

- **ME mm/12 hours**
  - Hirlam11 12-030 ME = 0.674
  - Hirlam22 12-030 ME = 0.413
  - Ecmwf 00+042 ME = 0.224
Verification of 10 meter wind forecasts against Swedish SYNOPs
The Nordic temperature problem

• A new surface and snow parameterization scheme by Stefan Gollvik and Patrik Samuelsson (similar for HIRLAM and the Rossby Center regional climate model)
• A HIRLAM 1D model study by Sander Tijm
HIRLAM NEW SURFACE SCHEME

- 5 surface energy balances
- 6 surface liquid water storages
- 3 soil temperature levels and 2 soil water levels
- Heat conduction
- Soil-freezing algorithm
- Physical snow description

Three aerodynamic resistances

Soil-freezing algorithm
The Nordic cold winter temperature problem

Graph showing daily temperature variations from 27/01 to 30/01 with specific dates and times.
The Nordic cold winter temperature problem
The Nordic cold winter temperature problem
The Nordic winter cold temp. is less of a problem with new snow scheme

H634_snow_1S40
Spring problem

Daily cycle wrong, min T too high, max T too low
Spring problem – fluxes wrong
Spring problem (6.2.1)
Spring problem (6.2.1)
Spring no longer a problem! (6.3.4snow)
Spring no longer a problem! (6.3.4snow)
Simulation of snow-bands in the Baltic Sea

From

Andersson and Gustafsson, 1994:

Coast of Departure and Coast of Arrival: Two important concepts for formation of convective snowbands over seas and lakes. MWR

• Snowbands in January 1987; Why does it snow so much in Oskarshamn?

• Simulations with HIRLAM 22 km, 16 levels

• Sensitivity to changes in coastlines, ice borders, orography heights and surface roughness
Reference experiment: No changes to coastlines etc.
Experiment 1:
Removal of the Bay of Finland
(Coast of departure)
Experiment 2:
Removal of the Stockholm peninsula
(Coast of arrival)
Experiment 3:
Removal of the Bay of Finland and the Stockholm Peninsula
Experiment 4:
Widening of the
“Baltic Sea”
Experiment 5:
Flat orography in Southern Scandinavia
Experiment 5:
Flat orography in Southern Scandinavia
Experiment 6:
Flat orography + constant surface roughness length in Southern Scandinavia
Summary

• High resolution observations needed for regional NWP
• Wind and humidity observations important
• Thermodynamic imbalances problematic
• A lot of suitable observations available

4-dimensional variational data assimilation
Development of HIRLAM 4D-Var.


2000: First experiments with ”non-incremental” 4D-Var.


2003-2004: Semi-Lagrangian scheme (SETTLS), outer loops (spectral or gridpoint HIRLAM) and multi-incremental minimization.

Single observation experiment with HIRLAM 4D-Var;

What is the effect of a single surface pressure observation increment of -5 hPa at + 5 hours in the assimilation window?

1: In the center of a developing low

2. In a less dynamically active area
Surface pressure increments for the Danish storm

4D-Var, spectral TL prop. of incr.

3D-Var

4D-Var; gp model prop. of incr.
Effects of a -5 hPa surface pressure observation increment at +5 h on the initial wind and temperature increments

Winds at model level 20 (500 hPa) and temperatures at level 30 (below)

NW-SE cross section with temperatures and normal winds
Effects of a -5 hPa surface pressure observation increment at +5 h in a less dynamically active area

Surface pressure assimilation increment at +0 h

Difference between non-linear forecasts at +6 h with and without the 4D-Var assimilation increment
Recent HIRLAM 4D-Var tests

• The SMHI 22 km area (306x306x40 gridpoints)
• SMHI operational observations (including AMSU-A and "extra" AMDAR observations)
• 6 h assimilation cycle; 3D-Var with FGAT; 6 h assimilation window in 4D-Var; 1 h observation windows
• 66 km assimilation increments in 4D-Var (linear grid); 44 km assimilation increments in 3D-Var (quadratic grid)
• Non-linear propagation of assimilation increments
• 3 months of data (January 2005, June 2005, January 2006)
Average upper air forecast verification scores – January 2006

\( \text{\( o3d = 3D-Var \)} \quad \text{\( o4d = 4D-Var \)} \)

Statistics for 62 stations
Wind speed
Period: 20060102-20060128
Solid RMS; Dotted STDV; Dashed BIAS; Number of cases 12254
Forecast lengths used: 12 24 36 48

Statistics for 62 stations
Temperature
Period: 20060102-20060128
Solid RMS; Dotted STDV; Dashed BIAS; Number of cases 12502
Forecast lengths used: 12 24 36 48

Wind speed
Temperature
Mean sea level pressure forecast verification scores – January 2006

\[ o_{3d} = 3D-\text{Var} \quad o_{4d} = 4D-\text{Var} \]
Time series of mean sea level pressure verification scores – January 2006

\[ o3d = 3dvar \quad o4d = 4D-Var \]
25 January 2006 case

3D-Var
+0 h

4D-Var
+0 h

3D-Var
+24 h

4D-Var
+24 h
Assimilating AMSU-A over sea ice in HIRLAM 3D-Var

From Harald Schyberg, Vibeke W. Thyness and Frank T. Tveter
Norwegian HIRLAM 20km model domain
Impact of added AMSU-A over sea ice experiment setup

- Experiment period 26 February to 30 April 2005
- HIRLAM 20 km resolution, 3D-Var assimilation,
- Conventional observations and direct use of AMSU with forward model RTTOV

Parallel suites with 6 hr cycling, forecasts up to 48 hrs, to highlight effect of observations over sea ice:

- Reference: No AMSU
- Experiment 1: Upper AMSU-A channels over sea ice only (Ch 1 to 5 only in “passive mode”)
- Experiment 2: Both upper channels and surface sensitive channels over sea ice, using simple emissivity estimate (not completed)
Ref-Exp1 increments, analysis
(RMS diff, Z 500 hPa)

Z0500, Sq[Var[Exp(+00) - Ref(+00)]]

Maximum value = 31.793
Verification statistics against observations on EWGLAM list (List of SYNOPs considered reliable over Europe)

- Problem: Scarcity of Arctic conventional observations for verification

- Overall statistics: Positive impact of adding AMSU-A on EWGLAM verification, but impact highly situation dependent
14 March 2005 00Z (+24 hrs exp solid and ref dashed)
14 March 2005 00Z (analysis exp solid and ref dashed)
Timeseries of pressure obs (blue dots) and ref (red) /exp (blue) forecasts (18-36hrs) at Bjørnøya
Conclusions from these impact and case studies

- Generally difficult to measure the impact over the sea ice regions
- Impact can be kept in model for long periods in data-sparse regions.
- Single observations sometimes not sufficient for bringing the model “back on track” – rejection by quality control checks
- Verifying with the conventional observation network over Europe, we find that the impact varies with circulation pattern. With general upper flow from the sea ice towards the North Atlantic: cases of significant impact identified.
- Inclusion in operational model runs with HIRLAM being implemented
Use of lower tropospheric channels

Large part of the Arctic ice cap comprised of closed, near 100% concentration, sea ice

Sea ice emissivity is usually more stable in time than typical weather systems

Can we take advantage of daily maps of sea ice properties (from SSMI) and use first-year (FY) and multi-year (MY) concentrations as predictors for AMSU channel surface emissivity over the inner Arctic?

Method implemented in IOMASA (starting point for further work in EU project DAMOCLES and possibly EUMETSAT OSI SAF)
Emissivities: A first simple approach

Output from the EUMETSAT OSI SAF chain. Ice concentrations are used as input to the emissivity calculations.
Estimation of typical FY and MY sea ice emissivities (Toudal, 2005)

Simplified theory for microwave radiative transfer, where the main assumptions are that

- The atmospheric attenuation can be reasonably approximated by an absorption coefficient and an effective atmospheric temperature $T_a$
- The water vapour load is minimal so the main contribution to the absorption is from oxygen
- Then the surface emissivity can be estimated from the measured brightness temperature $T_b$
Example timeseries of AMSU-A emissivities for a limited area of Multi Year ice north of Greenland (from Toudal)
Emissivities

Use OSI SAF FY and MY ice concentrations with typical values (Toudal) of AMSU emissivities for these surfaces:

\[ \epsilon = c_W \epsilon_W + c_F \epsilon_F + c_M \epsilon_M, \]
\[ c_W + c_F + c_M = 1. \]

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<th>First year ice</th>
<th>Multi year ice</th>
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<tr>
<td>1</td>
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<td>0.874</td>
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<tr>
<td>2</td>
<td>0.970</td>
<td>0.829</td>
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<tr>
<td>3</td>
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<td>0.744</td>
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Comparison with constant emissivity, channel 2

Channel 2 NO EMIS–MODEL
amsu–a on noaa15
R^2= 0.01  a0= 238.643  a1= 0.037

Mean difference
Nobs= 62402 NO EMIS model
mean diff= −30.396 , std= 18.809

Channel 2 WITH EMIS–MODEL
amsu–a on noaa15
R^2= 0.72  a0= 77.573  a1= 0.681

Mean difference. Nobs= 55619
mean diff= −7.798 , std= 10.687
Comparison with constant emissivity, channel 5

Channel 5 NO EMIS–MODEL
amsu–a on noaa15
R^2 = 0.94  a0 = 16.857  a1 = 0.925

Mean difference
Nobs = 62402  NO EMIS model
mean diff = 0.458 , std = 0.862

Channel 5 WITH EMIS–MODEL
amsu–a on noaa15
R^2 = 0.98  a0 = 9.7  a1 = 0.956

Mean difference
Nobs = 55619  mean diff = 0.476 , std = 0.621
Possible further developments on emissivities

Could probably improve present method with
- Further tuning and adjustment of emissivities using background departure statistics
- Add regional/seasonal and incidence angle dependence to pure FY and MY AMSU emissivities

But the long-term strategy should probably be a combination of a statistical approach and a surface microphysical emissivity model
- Better predictors for emissivity than SSMI-based FY and MY concentration retrievals
- Correlations of emissivities between channels exploited more directly
- Emissivity in control variable (implemented at SMHI) with a first guess estimate
- Feedback of obs departures to emissivity predictors
- Include meteorological history in a microphysical model of the sea ice surface (snow, freeze, melt, …)
Assimilation of AMSU-B moisture retrievals or AMSU-B radiances over sea ice surfaces?

Work done by
Per Dahlgren, SMHI within
EU-IOMASA.

• Only AMSU-B moisture retrievals have been assimilated so far
• Illustrates the verification problem
• REF: reference experiment, EXP: with assimilation of moisture retrievals
Impact of the AMSU-B moisture retrieval at different forecast lengths (REF-EXP):

+00h, +24h and +48h
Differences (EXP-REF) between monthly means of water vapor at model 30: +00h, +24h and +48h
Differences (EXP-REF) between monthly means of cloudiness at +12h: Low clouds, middle clouds and high clouds
Differences (EXP-REF) between monthly means of 2 meter temp. forecasts (12h)
Verification of forecasts against Arctic SYNOPs
Concluding remarks

• The new HIRLAM surface/soil/snow scheme seems to “solve” important Nordic temperature forecasting problems (without degrading other features?)

• Convective snow band in the Baltic Sea in January 1987 were sensitive to distribution of land, sea and sea ice.

• HIRLAM 4D-Var seems to provide flow-dependent influence of observations and consistently improved forecast scores.

• More upper air data are needed in the Arctic (for example satellite radiances over sea ice). We need to improve the lower boundary conditions for the assimilation.

• Data impact studies may sometimes be confused by model (and observation) biases.
Thank You!