Recent developments in the vertical discretization of the ECMWF model with impact on the stratosphere and tropopause

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Outline

- Finite-element discretization for the vertical
- Cubic spline interpolation in the vertical semi-Lagrangian advection
- Numerical instability during sudden stratospheric warming events at T_L511
- Increase in vertical resolution (60 levels \rightarrow 90 levels)



Finite-element (FE) discretization for the vertical

- We use cubic B-splines as basis functions with compact support (finite elements).
- No staggering of variables used. All (including pressure) are held on the same set of levels (full levels). (Good for semi-Lagrangian advection.)
- Only non-local operations are evaluated in FE space, products of variables are evaluated in physical space. (Similar to spectral transform method in the horizontal.)
- In the semi-Lagrangian version of the ECMWF model, the only non-local operations in the vertical are <u>integrations</u> (no derivatives). Therefore, we have derived the FE form only for the integration operator.



FE scheme: Integral operator in finite-element form

$$F(x) = \int_{0}^{x} f(y) dy$$

Expanding f and F in terms of sets of linearly independent functions with compact support {e_i} and {d_i}, respectively: $\sum_{i=1}^{M} C_i d_i(x) = \sum_{i=1}^{N} c_i \int_{0}^{x} e_i(y) dy$

Using the Galerkin method with $\{d_i\}$ as test functions

$$\sum_{i=1}^{M} C_{i} \int_{0}^{1} d_{j}(x) d_{i}(x) dx = \sum_{i=1}^{N} c_{i} \int_{0}^{1} [d_{j}(x) \int_{0}^{x} e_{i}(y) dy] dx, \quad j = 1, ..., M$$

In matrix form:
$$\underline{\underline{AC}} = \underline{\underline{BC}} \iff \underline{\underline{C}} = \underline{\underline{A}}^{-1} \underline{\underline{BC}} \quad \text{(integral in FE space)}$$

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Incorporating the transformation to finite-element space and back into the Intergal operator, i.e. $\underline{C} = \underline{\underline{S}}^{-1} \underline{f}$ & $\underline{F} = \underline{\underline{S}} \underline{\underline{C}}$ $\Rightarrow \underline{F} = \widetilde{\underline{S}} \underline{\underline{A}}^{-1} \underline{\underline{B}} \underline{\underline{S}}^{-1} \underline{f}$

FE scheme: Cubic B-splines as basis functions



No staggering of basis set {d_i} with respect to set {e_i} (good for semi-Lagrangian adv.)

Condition F(0)=0 enforced by incorporation into basis functions, i.e. $d_i(0)=0$ for all i. Basis functions d_0 , $d_1 \& d_2$ computed by linear combination of e_{-1} with e_0 , $e_1 \& e_2$, respectively.

Not restricted to regular spacing of nodes.



FE scheme: Hat-functions (linear splines) as basis functions



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FE scheme: Accuracy

<u>Test</u>: numerical integration of $\sin(6\pi x)$, $x \in [0,1]$ for different resolutions with N equidistantly spaced nodes <u>Reference</u>=analytical integral I_A. <u>Error</u> =max{(I_N-I_A)/I_A} in %

	FD scheme	Linear FE	Cubic FE	Cubic collocation
N=60	0.82e+0	0.14e-2	0.90e-8	0.14e-2
N=120	0.21e+0	0.85e-4	0.31e-10	0.85e-4
estim. order	2	4	8	4

On nodes $O(h^{2(k+1)})$ where k is the degree of the basis functions Superconvergence and h the distance between nodes



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Benefits from the FE scheme

- FE scheme improves the treatment of the gravity wave terms and dampens the computational (zigzag) mode in the vertical present in finite-difference schemes with no staggering of winds and temperature (Lorenz grid).
 - => Reduces the amplitude of grid-wave noise in the stratosphere.



Improved vertical integration of the continuity equation leads to a more accurate vertical velocity for semi-Lagrangian advection. => improved tracer conservation





Cubic spline interpolation for the vertical in the semi-Lagrangian advection

Ozone conservation





Numerical noise during sudden stratospheric warming (1)



Noise appears only in integrations at high horizontal resolution ($T_L 511$)

It is highly predictable (up to 8 days ahead) suggesting that it is linked to a specific well-predicted feature of the large-scale flow.

Forecasts don't fail, noise disappears again when flow pattern changes

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Numerical noise during sudden stratospheric warming (2)





Noise during sudden stratospheric warming (3)

Vertical trajectory calculation for semi-Lagrangian advection:

$$\eta_A(t+\Delta t) = \eta_D(t) + \Delta t \frac{\dot{\eta}_A(t) + \left[2\dot{\eta}(t) - \dot{\eta}(t-\Delta t)\right]_D}{2}$$



Smoothing of vertical velocity by least square fit through 4 surrounding points instead of just linear interpolation between 2 points. Done only for vertical velocity used in vertical trajectory calculation.



Numerical noise during sudden stratospheric warming (4)

Vertical cross section of Divergence: 20020925 12h step 24h, T511L60 0.000025 (0.0001) Model Levels 20 24 150°W 100°W 50'W 50°E 0° -65.0S "Divergence at level 13: 20020925 12h step=24h T511L60 120*



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40°W

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Vertical cross section of Divergence: 20020925 12h step 24h, T511L60



Increase in vertical resolution



L90: Fit to Radiosonde Temperatures in the Analysis



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Tropical Cold-Point Tropopause in L90 / L60



Averaged over the deep tropics [10S to 10N]. Analyses and radiosondes averaged in time from 20020601 to 20020615 Forecasts averaged over whole month of June.

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L90: Impact on ozone conservation



Model top raised from 0.1hPa to 0.01hPa



L91

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L91: Comparison with CIRA86 Climatology for July





L91: Comparison with CIRA86 Climatology for January





L91: Reduction in vertical velocity in the tropics



Future work

Based on the L90 or L91 model version, try to

- improve vertical transport in the stratosphere
 - benefit for ozone assimilation and interactive ozone with radiation
- reduce large model errors near the stratopause
 - less problems with assimilation of satellite data
- Continue work on the use of vertical spline interpolation in the semi-Lagrangian advection.
- Improve upper boundary condition (Nils Wedi).

