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Generalised Nonhydrostatic Model

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Mass, entropy and momentum conservation laws:

Smolarkiewicz et. al. JCP 2016

$$\frac{\partial \mathcal{G}\varrho}{\partial t} + \nabla \cdot (\mathcal{G}\varrho \boldsymbol{v}) = 0$$

$$\frac{\partial \mathcal{G}\varrho\theta'}{\partial t} + \nabla \cdot \left(\mathcal{G}\varrho\boldsymbol{v}\theta'\right) = -\mathcal{G}\varrho\left(\tilde{G}^T\boldsymbol{u}\cdot\nabla\theta_a - \mathcal{H}\right)$$

On a rotating sphere, systems and 3D radaptivity.

$$\frac{\partial \mathcal{G} \varrho \boldsymbol{u}}{\partial t} + \nabla \cdot (\mathcal{G} \varrho \boldsymbol{v} \otimes \boldsymbol{u}) = -\mathcal{G} \varrho \left(\Theta \tilde{G} \nabla \varphi + g \Upsilon_B \frac{\theta'}{\theta_b} + \boldsymbol{f} \times (\boldsymbol{u} - \Upsilon_C \boldsymbol{u}_a) - \mathcal{M}'(\boldsymbol{u}, \boldsymbol{u}, \Upsilon_C) - \boldsymbol{\mathcal{D}} \right)$$

$$\varphi := \left[\rho(\boldsymbol{x}, t), \frac{\rho_b(z)\theta_b(z)}{\theta(\boldsymbol{x}, t)}, \rho_b(z) \right] \quad \varphi := \left[c_p \theta_0 \pi', c_p \theta_0 \pi, c_p \theta_b \pi \right] \quad \Theta := \left[\frac{\theta}{\theta_0}, \frac{\theta}{\theta_0}, 1 \right] \quad \Upsilon_B := \left[\frac{\theta_b(z)}{\theta_a(\boldsymbol{x})}, \frac{\theta_b(z)}{\theta_a(\boldsymbol{x})}, 1 \right] \quad \Upsilon_C := \left[\frac{\theta}{\theta_a(\boldsymbol{x})}, \frac{\theta}{\theta_a(\boldsymbol{x})}, 1 \right]$$

compressible, pseudo-incompressible, anelastic

cast in generalised timedependent curvilinear coordinates, to enable a range of coordinate



NFT Integrator



Advection: MPDATA

Multidimensional Positive Definite Advection Transport Algorithm

$$\frac{\partial G\Phi}{\partial t} + \nabla \cdot (\boldsymbol{V}\Phi) = G\mathcal{R}$$

$$\Phi_i^{n+1} = \mathcal{A}_i \left(\Phi^n + \frac{1}{2} \delta t \mathcal{R}, \boldsymbol{V}^{n+\frac{1}{2}}, G^n, G^{n+1} \right) + \frac{1}{2} \delta t \mathcal{R}^{n+1}$$

Smolarkiewicz & Szmelter, JCP 2009, Acta Geo 2011

Heimholtz: from evolutionary form of the gas law $\frac{1}{\mathcal{G}}\nabla\cdot(\mathcal{G}\boldsymbol{v}^{\nu}) - \frac{1}{\xi}\frac{\phi_{a}}{\phi^{\nu-1}}\left(\frac{1}{\varrho^{*}\phi_{a}}\nabla\cdot(\varrho^{*}\phi_{a}\boldsymbol{v}^{\nu}) - \frac{1}{\varrho^{*}}\nabla\cdot(\varrho^{*}\boldsymbol{v}^{\nu})\right) \\ - \frac{1}{\delta t\xi\phi^{\nu-1}}\left(\varphi^{\nu} - \hat{\varphi}\right) = 0$

Smolarkiewicz et. al. JCP 2016

$$\Rightarrow \frac{1}{\zeta} \nabla \cdot \zeta \left(\check{\check{\boldsymbol{v}}} - \tilde{G}^T C \nabla \varphi \right)$$

Implicit Integrator: GCR(k)

non-symmetric, preconditioned Krylov-subspace elliptic solver

For any initial guess, ϕ^0 , set $r^0 = \mathcal{L}(\phi^0) - \mathcal{R}$, $p^0 = \mathcal{P}^{-1}(r^0)$; then iterate:

For $n = 1, 2, \dots$ until convergence

for
$$\nu = 0, ..., k - 1$$

$$\beta = \frac{\langle r^{\nu} \mathcal{L} (p^{\nu}) \rangle}{\langle \mathcal{L} (p^{\nu}) \mathcal{L} (p^{\nu}) \rangle},$$

$$\phi^{\nu+1} = \phi^{\nu} + \beta p^{\nu},$$

$$r^{\nu+1} = r^{\nu} + \beta \mathcal{L} (p^{\nu}),$$
exit if $||r^{\nu+1}|| \le \epsilon,$

$$e = \mathcal{P}^{-1} (r^{\nu+1}),$$
evaluate $\mathcal{L}(e) = \frac{1}{\rho^*} \nabla \cdot C \nabla e,$
for $l = 0, ..\nu$

$$\alpha_l = \frac{\langle \mathcal{L}(e) \mathcal{L} (p^l) \rangle}{\langle \mathcal{L} (p^l) \mathcal{L} (p^l) \rangle},$$

$$p^{\nu+1} = e + \sum_l^{\nu} \alpha_l p^l,$$

$$\mathcal{L} (p^{\nu+1}) = \mathcal{L}(e) + \sum_l^{\nu} \alpha_l \mathcal{L} (e^l),$$
reset $[\phi, r, e, \mathcal{L}(e)]^k$ to $[\phi, r, e, \mathcal{L}(e)]^0$

Smolarkiewicz & Margolin, 2000

Edge based finite volume discretisation



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Smolarkiewicz & Szmelter, Acta Geo 2011

Convective Planetary Boundary Layer University

Smolarkiewicz et al JCP 2013



Vertical velocity in central xz cross section (top) and the horizontal plane (bottom): instantaneous solution for triangular prismatic mesh after ~13 eddy turnover times

Schmidt & Schumann JFM 1989



Vertical profile of dimensionless resolved heat flux, temperature variance, and vertical velocity; for dimensionless height z/z



Edge-based T Edge-based C EULAG SS LES SS LES Observation

Preconditioning

Elementary preconditioner, slower elliptic solver

Advanced Preconditioner faster elliptic solver

$$I \quad \leftarrow \quad \mathcal{P} \quad \rightarrow \quad \mathcal{L}$$

Choose \mathcal{P} close to \mathcal{L} , but easier to solve.

Take only diagonal terms in C: $\mathcal{P} \approx \mathcal{L}$

Invert \mathcal{P}_{z} implicitly, with \mathcal{P}_{H} lagged:

$$\frac{\partial ee}{\partial \tilde{\tau}} = \mathcal{P}(e) - r \quad \Rightarrow \quad \frac{ee^{\mu+1} - ee^{\mu}}{\Delta \tilde{\tau}} = \mathcal{P}_H(e^{\mu}) + \mathcal{P}_z(e^{\mu+1}) - r$$

$$(I - \Delta \tilde{\tau} \mathcal{P}_z) e^{\mu + 1} = e^{\mu} + \Delta \tilde{\tau} \left(\mathcal{P}_H(e^{\mu}) - r^{\nu + 1} \right)$$

For any initial guess, ϕ^0 , set $r^0 = \mathcal{L}(\phi^0) - R$, $p^0 = \mathcal{P}^{-1}(r^0)$; then iterate:

For
$$n = 1, 2, ...$$
 until convergence

for
$$\nu = 0, ..., k - 1$$

$$\beta = \frac{\langle r^{\nu} \mathcal{L} (p^{\nu}) \rangle}{\langle \mathcal{L} (p^{\nu}) \mathcal{L} (p^{\nu}) \rangle},$$

$$\phi^{\nu+1} = \phi^{\nu} + \beta p^{\nu},$$

$$r^{\nu+1} = r^{\nu} + \beta \mathcal{L} (p^{\nu}),$$
exit if $||r^{\nu+1}|| \le \epsilon,$

$$e = \mathcal{P}^{-1} (r^{\nu+1}),$$
evaluate $\mathcal{L}(e) = \frac{1}{\rho^*} \nabla \cdot C \nabla e,$
for $l = 0, ..\nu$

$$\alpha_l = \frac{\langle \mathcal{L}(e) \mathcal{L} (p^l) \rangle}{\langle \mathcal{L} (p^l) \mathcal{L} (p^l) \rangle},$$

$$p^{\nu+1} = e + \sum_l^{\nu} \alpha_l p^l,$$

$$\mathcal{L} (p^{\nu+1}) = \mathcal{L}(e) + \sum_l^{\nu} \alpha_l \mathcal{L} (e^l)$$
reset $[\phi, r, e, \mathcal{L}(e)]^k$ to $[\phi, r, e, \mathcal{L}(e)]^0$





Smolarkiewicz & Margolin, 2000





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Octahedral mesh has reasonably consistent structure

Latitude n = 1 ... Nhas 20 + 4n longitudes



Remove every other latitude

Half the number of longitudes at the remaining latitudes

Reduces horizontal nodes by a factor of ~ 4

Need to pick starting latitude

Coarse mesh is almost octahedral

All coarse mesh nodes coincide with nodes on the fine mesh







Octahedral 16 mesh:

Remove odd latitudes

computational domain

Single level of mesh coarsening





Octahedral 16 mesh:

Remove odd latitudes

computational domain

Single level of mesh coarsening



Multigrid Preconditioning







Smoother / Solver, varying number of iterations depending requirements

$$(I - \Delta \tilde{\tau} \mathcal{P}_z) e^{\mu + 1} = e^{\mu} + \Delta \tilde{\tau} \left(\mathcal{P}_H(e^{\mu}) - r^{\nu + 1} \right)$$

V-Cycle: Solve on h=1, else cycle (e_h, r_h) finalise cycle (e_h, r_h) : $\mathrm{Smooth}(e_h)$ calculate residule: $rr_h = \mathcal{P}(e_h) - r_h$ restrict $rr_h \to rr_{h-1}$: $rr_{h-1} = \mathcal{I}_h^{h-1} rr_h$ initialise solution error: $ee_{h-1} \equiv e - e'$ $cycle(ee_{h-1}, rr_{h-1})$ Interpolate solution error: $ee_h = \mathcal{I}_{h-1}^h ee_{h-1}$ Add coarse grid error: $e_h = e_h - ee_h$ $\mathrm{Smooth}(e_h)$

Multigrid Preconditioning

Baroclinic Instability – 9 days.

Compressible Nonhydrostatic Equations. Multigrid / implicit vertical preconditioner.

Multigrid: Wall time ~ 4hrs Elliptic solver iterations average – 3:0

Without multigrid: Wall time ~ 23hrs Elliptic solver iterations average – 60:7 Setup: Octahedral 80 grid, time-step 900s, 1/1 task/thread,





Conclusion



• NFT-MPDATA solvers based on an edge based finite volume method for unstructured grids provide a high quality results for a wide rage of mesh types and is applicable for all scales atmospheric flows.

 Substantial efficiency gains have been achieved by introducing a horizontal multigrid preconditioner to a Krylov solver of Helmholtz equations forming a part of a global nonhydrostatic model.