Ensemble experiments for global ocean reanalyses and air-sea coupled covariances

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CMCC Contribution to ERA-CLIM2 WP2

 Global ocean ensemble generation for different applications (flow-dependent background-error covariances, initial conditions for seasonal predictions, probabilistic ocean reanalyses)

Sensitivity of global ocean heat content to atmospheric forcing and other uncertainties

• Ensemble-derived coupled covariances for use in strongly coupled data assimilation

Towards idealized strongly coupled data assimilation

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Sensitivity of GOHC to atmospheric forcing

Goals

- Evaluate global ocean heat content changes associated with different atmospheric forcing (from reanalyses) in ensemble systems with (REAENS) and without (CTRENS) ocean data assimilation
- Evaluate sensitivity to the atmospheric forcing in comparison with other • sources of uncertainty in ocean reanalyses
- Provide indications for optimal ensemble generation •

Strategy



Pros:

- Straightforward generation

ERA-MERRA

JRA-55

NCEP

CESR

- Implied assessment of atmospheric reanalysis accuracy from the ocean point of

Cons:

- The ensemble is under-dispersive, ie only the uncertainty in the atmospheric forcing is spanned

Sensitivity of GOHC to atmospheric forcing



Ocean Heat Content Anomaly

Contributions to total heat content tendency (1979-2013, W m⁻²)

	Surface Heat Flux	Assimilation Contribution	Total heat content tendency
CTRENS	-0.50 ± 0.64	-	-0.50 ± 0.64
REAENS	-3.05 ± 0.71	4.53 ± 3.62	1.48 ± 0.18



When Ocean data assimilation is switched on, the impact of atmospheric forcing is marginal (about 10% of the mean signal)

Without Ocean data assimilation, the atmospheric forcing is wrong (cooling) and the spread exceeds the mean signal

Sensitivity of GOHC to uncertainty sources

Comparing sensitivities from perturbed experiments using GOHC norm metrics

ICs: Initial conditions **ASS**: Assimilation configuration (background and observation errors)

OMP: Ocean model physics (small scale physics)

INS: In-situ observations (XBT correction, quality check, sampling)

SST: Sea-surface temperature the reanalysis is nudged to **ICE**: Sea-ice concentration the reanalysis is nudged to **ATM**: Atmospheric forcing



Ocean Heat Content Anomaly Sensitivity [0-bottom]



Ocean Heat Content Anomaly Sensitivity [0–700]

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Coupled covariances and strongly coupled DA

Motivation: Weekly coupled DA proves successful in improving near-surface atmopsheric parameters in simple configurations (especially in the Tropics). Example: Tropical T2m skill scores with/without SST nudging. Does strongly coupled DA lead to further improvements?

Strategy: Use simplified ABL model coupled to NEMO to test the impact of strongly coupled (via ensemble-derived inter-medium covariances and balances)



Coupled covariances and strongly coupled DA

Modeling framework

 NEMO-ORCA05L75 global configuration + CheapAML atmospheric boundary layer model (Deremble et al., 2013):

 $\partial \left(\textbf{T}_{2m}, \textbf{q}_{2m} \right) / \partial t = ADV[\textbf{u}, (\textbf{T}_{2m}, \textbf{q}_{2m})] + DIFF\left[(\textbf{T}_{2m}, \textbf{q}_{2m}) \right] + THDY[\textbf{SST}, \textbf{u}, (\textbf{T}_{2m}, \textbf{q}_{2m}), \textbf{H}_{ABL}]$

- It allows augmenting the ocean state control parameters to include $\rm T_{2M}$ and $\rm Q_{2M},$ now prognostic, in both model and 3DVAR
- Wind is not prognostic and imposed externally (ERA-Interim)

Ensemble System

- Used to derive ensemble airsea cross-covariances
- Ensemble system: Period 2009-2013, 12 members with perturbation of wind (20CRv2d ensemble anomalies) and model physics (EOS, Brankart et al., 2013)



Horizontal average of temperature ensemble spread

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A simplified air-sea balance operator

In order to couple the sea-surface variables with 2m atmospheric variables, balances might be thought either purely statistical, or purely analytical, or mixed (balanced + unbalanced components)

We introduce a balance operator that maps the increments of SST onto those of (**T**_{2m}, **Q**_{2m}) and uses tangent-linear version of CORE bulk formulas (Large & Yeager, 2007)

- $\delta \mathbf{T}_{2m} = \Delta t \left[\delta \mathbf{Q}_{LW} \left(\delta \mathbf{SST} \right) + \delta \mathbf{Q}_{SEN} \left(\delta \mathbf{SST} \right) \right] / \left[\rho_A c \rho_A \mathbf{H}_{ABL} \right]$ (no condensation in ABL)
- $\delta \mathbf{q}_{2m} = \Delta t \left[\delta \mathbf{E} \left(\delta \mathbf{SST} \right) \right] / \left[\rho_A \mathbf{H}_{ABL} \right]$

Where the exchange coefficients (**Ce, Ch** for Evaporation and Sensible heat, respectively) are assumed not to depend on **SST** and taken from the fully non-linear model.

This allows to compare purely statistical balances – cross-covariances of (T,S) and (T_{2m}, Q_{2m}) – and mixed – cross-covariances of (T,S) with the unbalanced components (T_{2m}, Q_{2m}) .

Coupled covariances

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MAIN RESULTS

- Sensitivity of atmospheric forcing in ocean reanalyses is assessed through multi-forcing ensemble reanalysis and compared to other sources of uncertainty: the impact of atmospheric forcing is generally marginal when ocean data assimilation is used; importance of in-situ observation pre-processing and surface nudging
- An ABL model coupled to NEMO is used to derive (full or unbalanced) coupled cross-covariances for use in idealized strongly coupled DA experiments

NEXT STEPS

- Ongoing: extension of 3DVAR control vector to include T2m, Q2m
- Planned: Impact of 3DVAR assimilation of ocean observations on T2m, Q2m skill scores with difference configuration of coupled covariances and introduction of flow-dependence of covariances

Thank you



